

What are the Kuiper Belt objects telling us about planet formation?

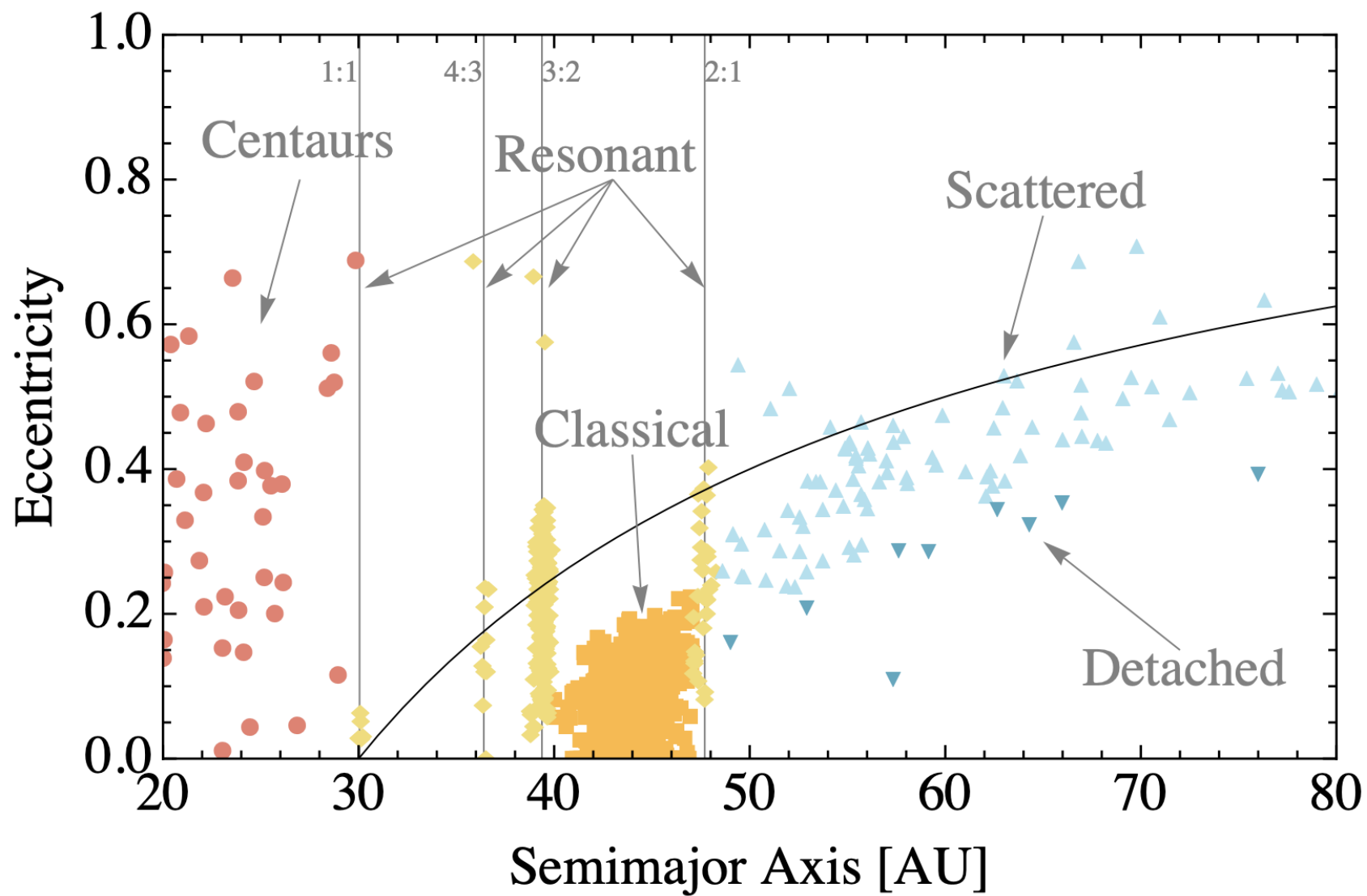


Wladimir Lyra
New Mexico State University

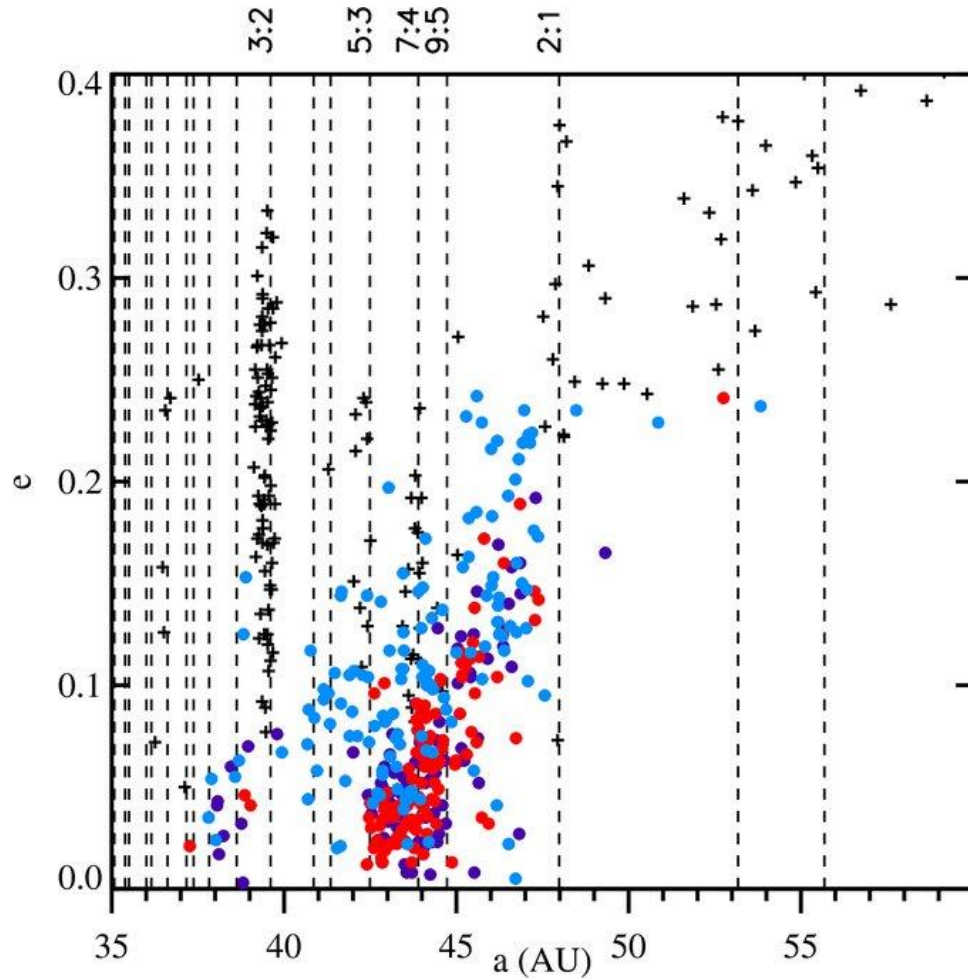


***Planet Formation in the Southwest (PFITS+)
Collaboration***

Structure of the Kuiper Belt



Structure of the Kuiper Belt



+ Resonant and Scattered

Cold Classical $i < 2^\circ$

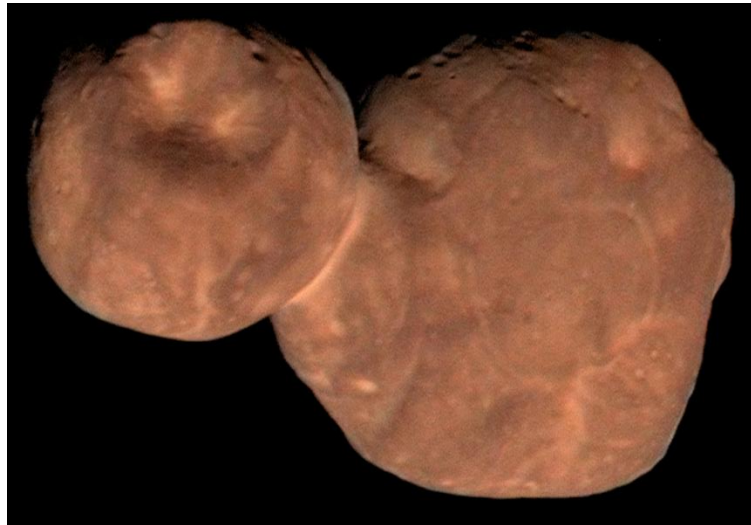
"Ambiguous" $2^\circ < i < 6^\circ$

Hot Classicals $i > 6^\circ$

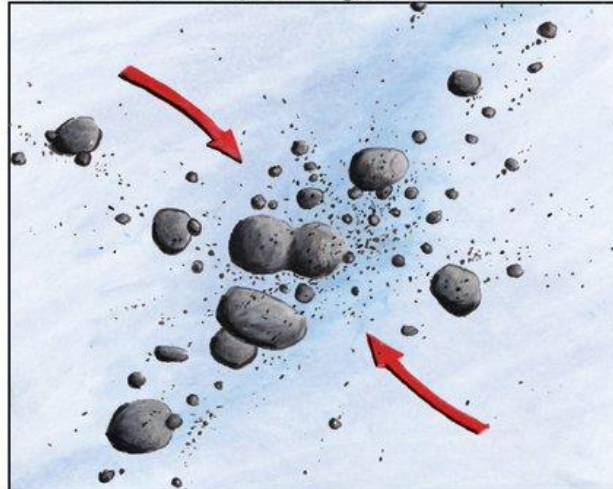
Cold Classicals: Presumably
pristine planetesimals

Arrokoth (MU₆₉)

The Formation of 2014 MU69

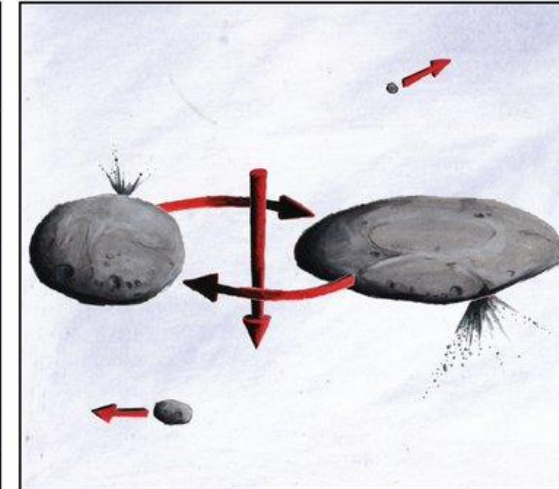


About 4.5 billion years ago...



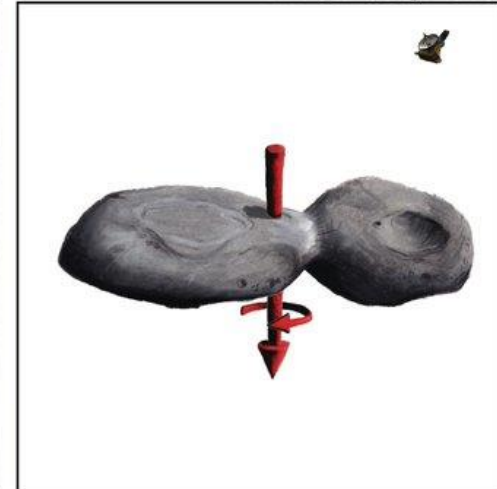
A rotating cloud of small, icy bodies starts to coalesce in the outer solar system.

 New Horizons / NASA / JHUAPL / SwRI / James Tuttle Keane



Eventually two larger bodies remain.

...1 January 2019.



The two bodies slowly spiral closer until they touch, forming the bi-lobed object we see today.

New Horizons Flyby, Jan 2019

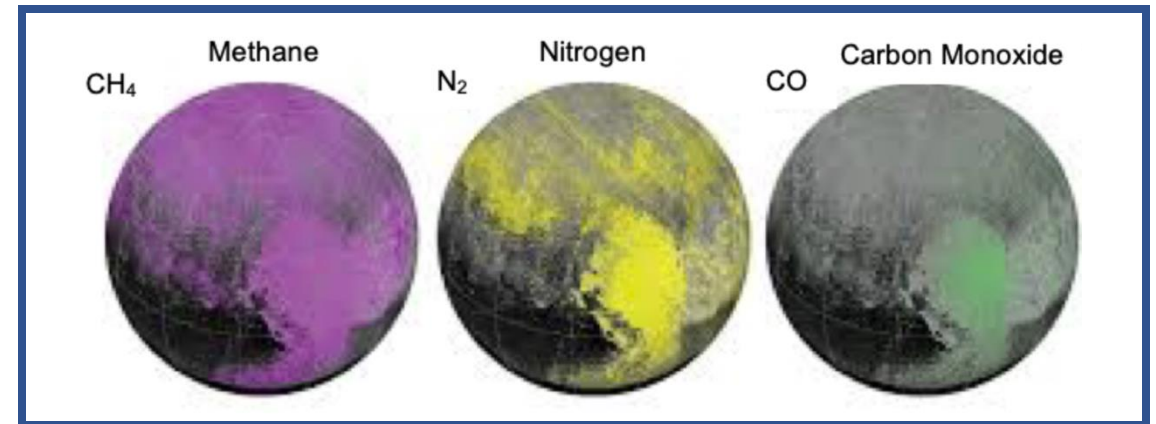
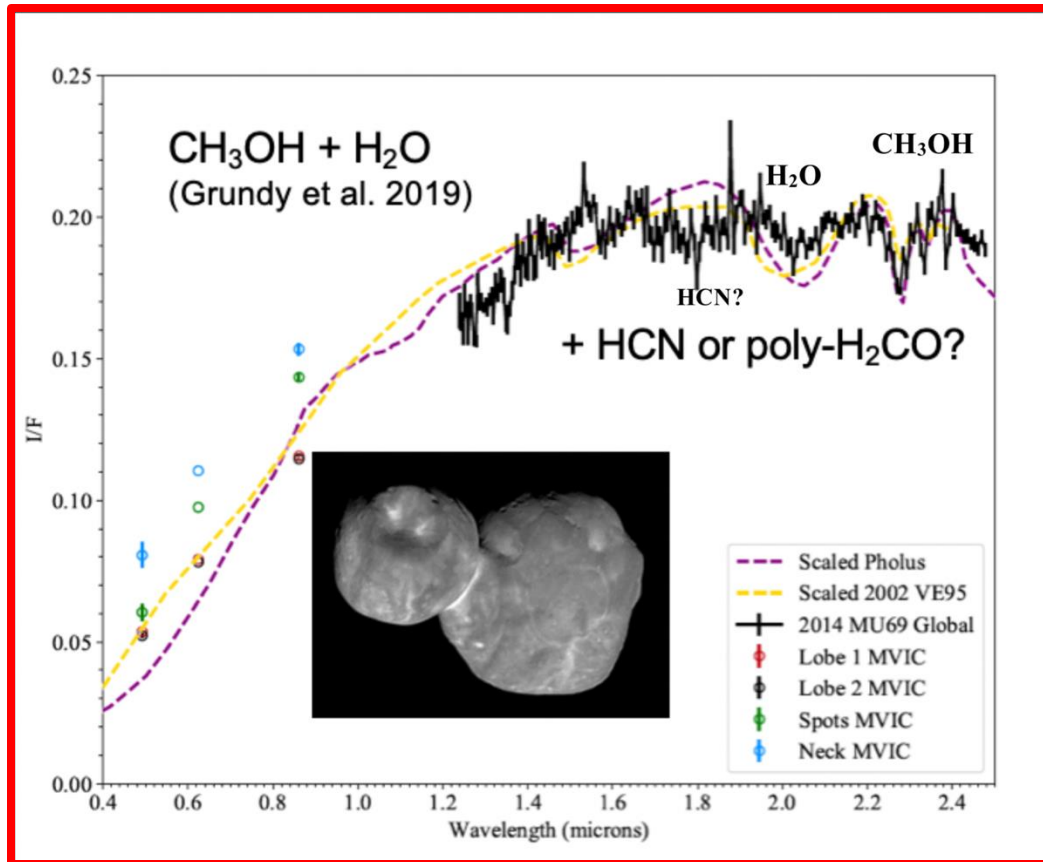
Sketch by J.T. Keane



Arrokoth and Pluto ices are different

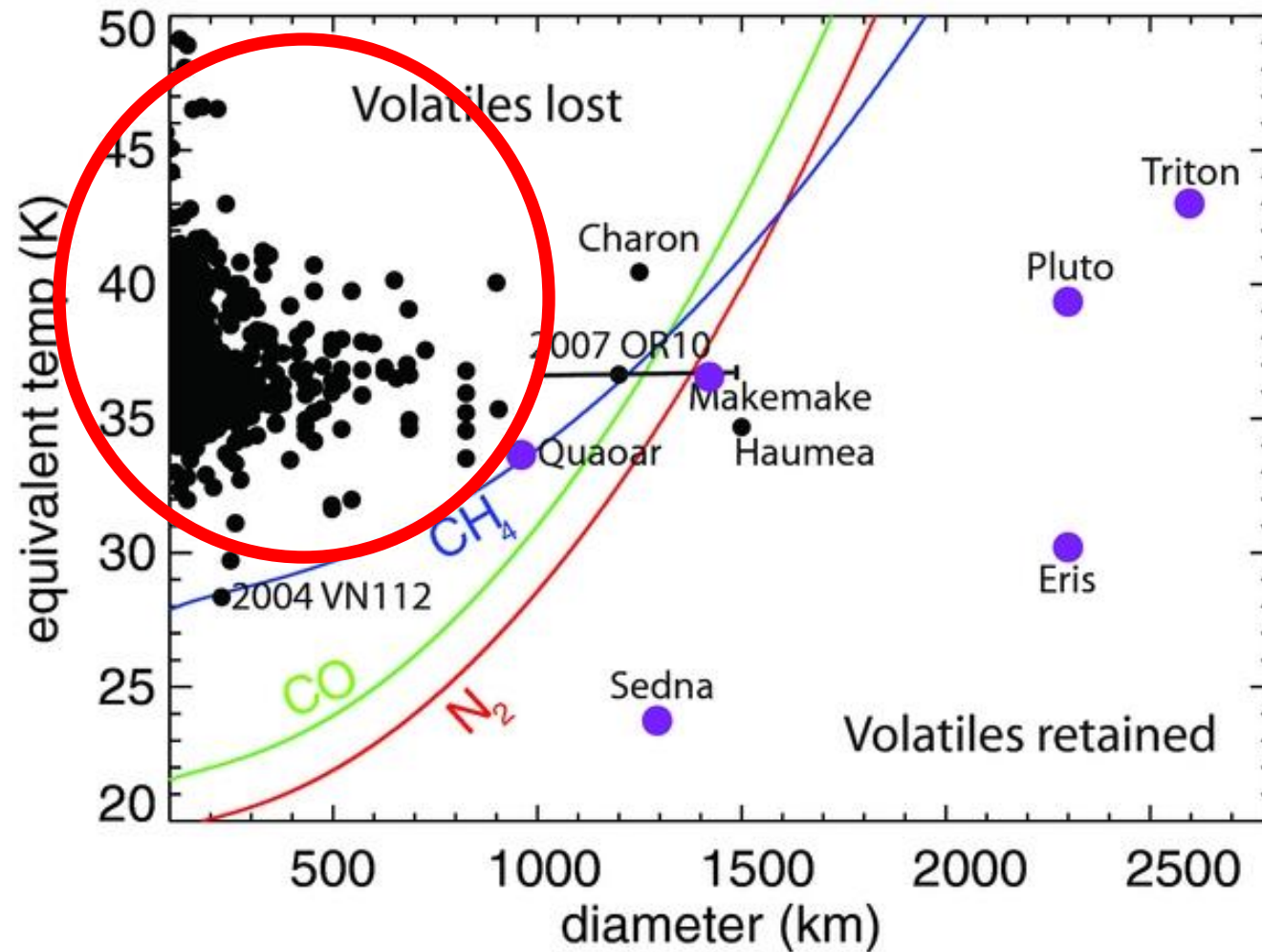
Arrokoth : Methanol, H₂O, HCN

Pluto : CH₄, N₂, CO



Retention of volatiles

Arrokoth formed with volatiles but lost them



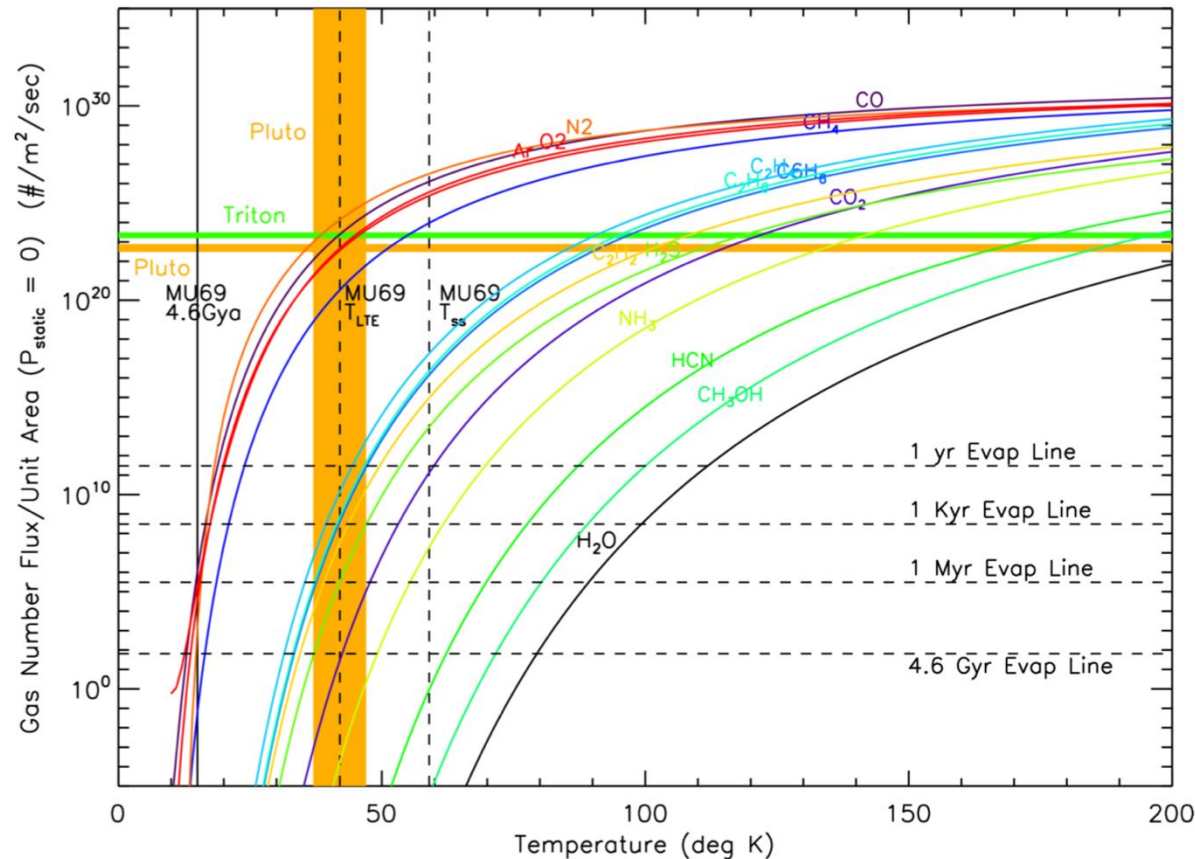
Retention of volatiles

Pluto cannot have been formed by accumulation of bodies like Arrokoth

Hypervolatiles (CH_4 / CO / N_2)

Lost under vacuum pressure and microgravity in ~ 1 Myr for 40 K

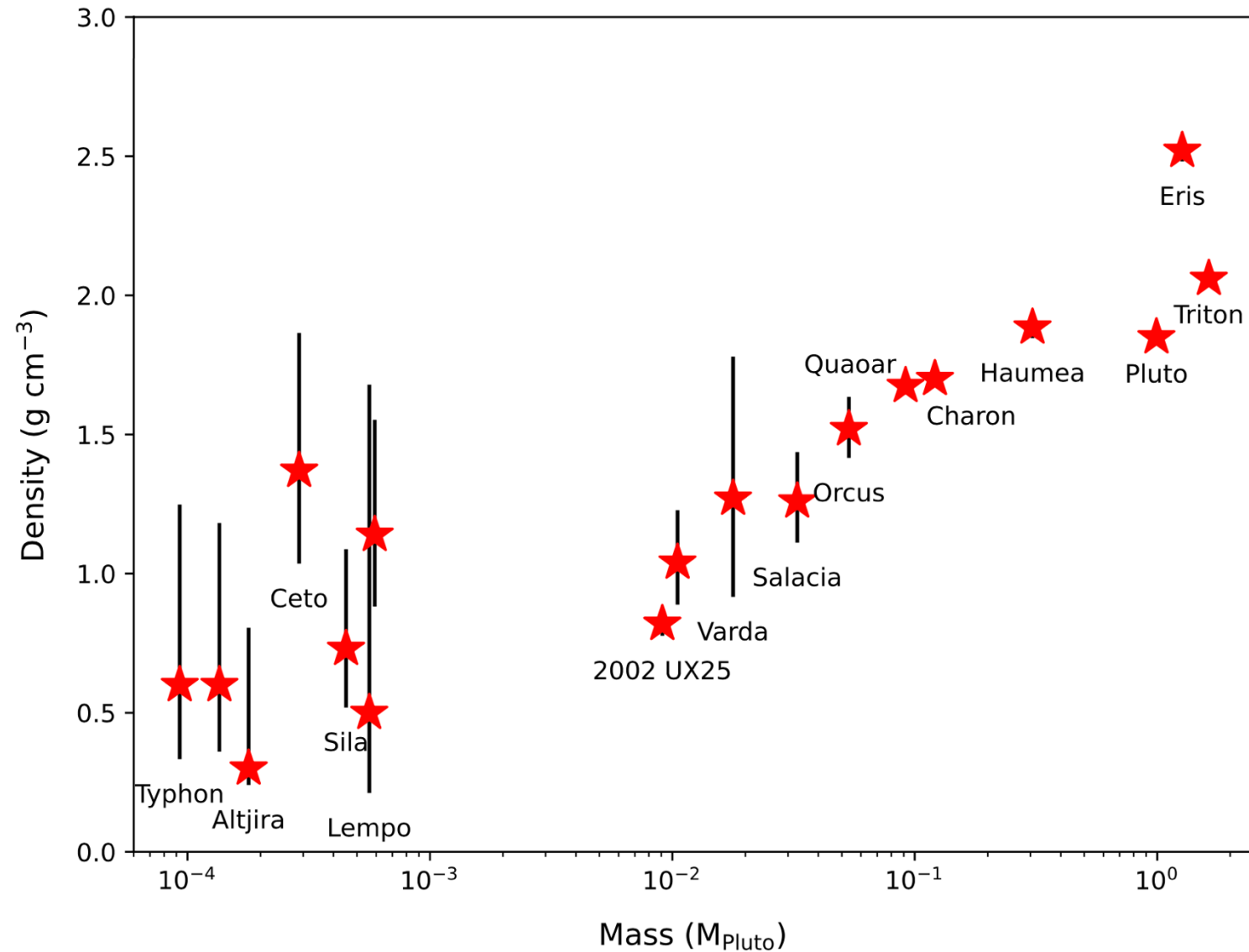
Retained for long times only for $T < 20\text{K}$



Arrokoth hypervolatiles lost when disk dissipates and temperature rises to 40K.

Hypervolatile-rich **Pluto** cannot form via planetesimal accretion

The size-density relationship of Kuiper Belt objects



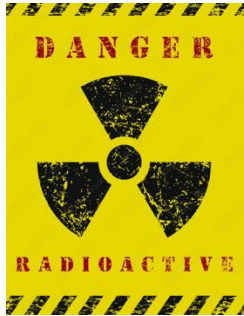
Brown (2013), McKinnon et al. (2017), Noll et al. (2020), Cañas & Lyra + (2024)

Data; Thomas (2000), Stansberry et al. (2006), Grundy et al. (2007), Brown et al. (2011), Stansberry et al. (2012), Brown (2013), Fomasier et al. (2013), Vilenius, et al. (2014), Nimmo et al. (2016), Ortiz et al. (2017), Brown and Butler (2017), Grundy et al. (2019), Morgado et al. (2023), Pereira et al. (2023).

Previous best bet: Porosity removal by gravitational compaction

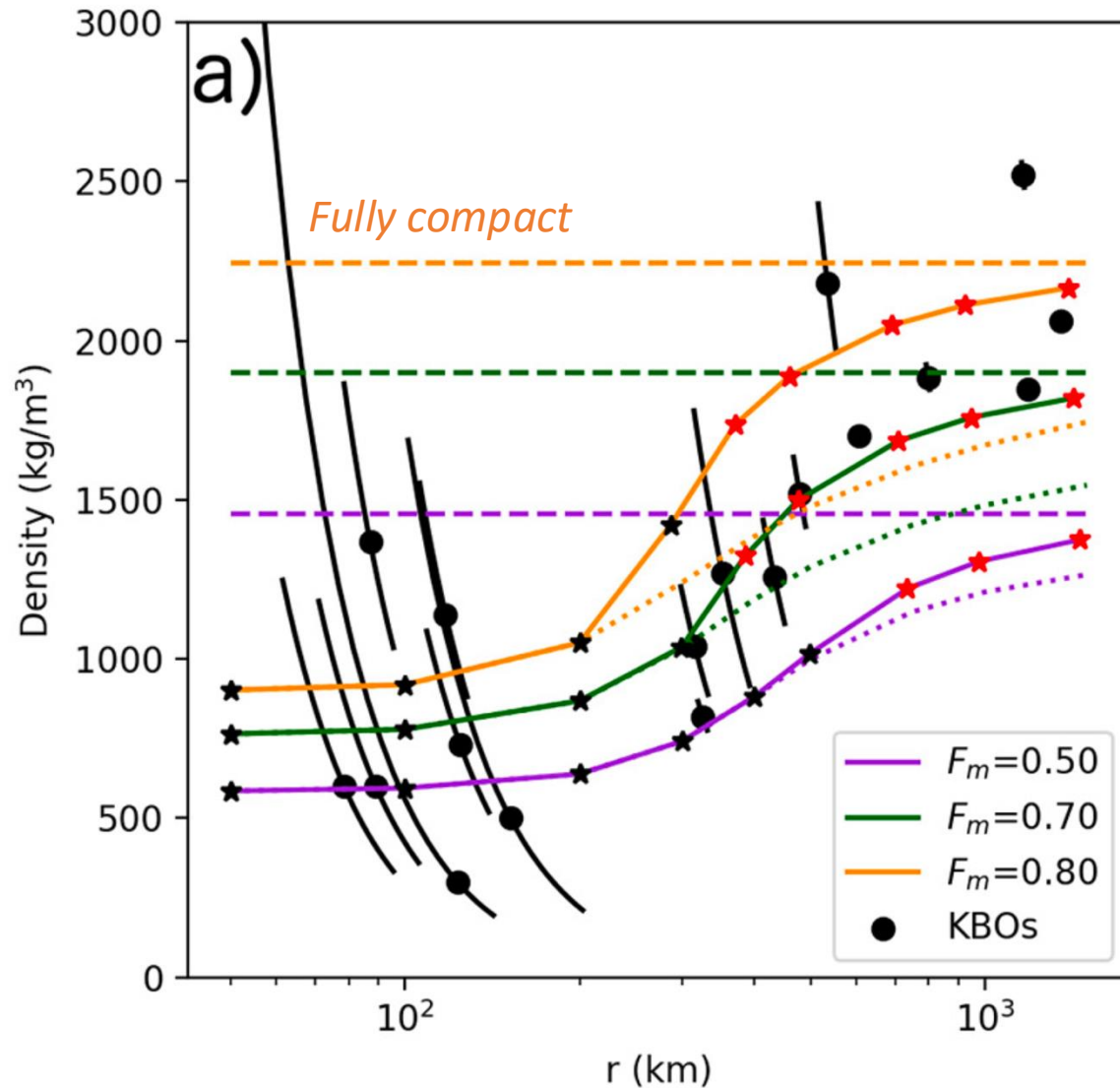
Problem

- Timing! ^{26}Al would melt if formed within 4 Myr



Assumptions

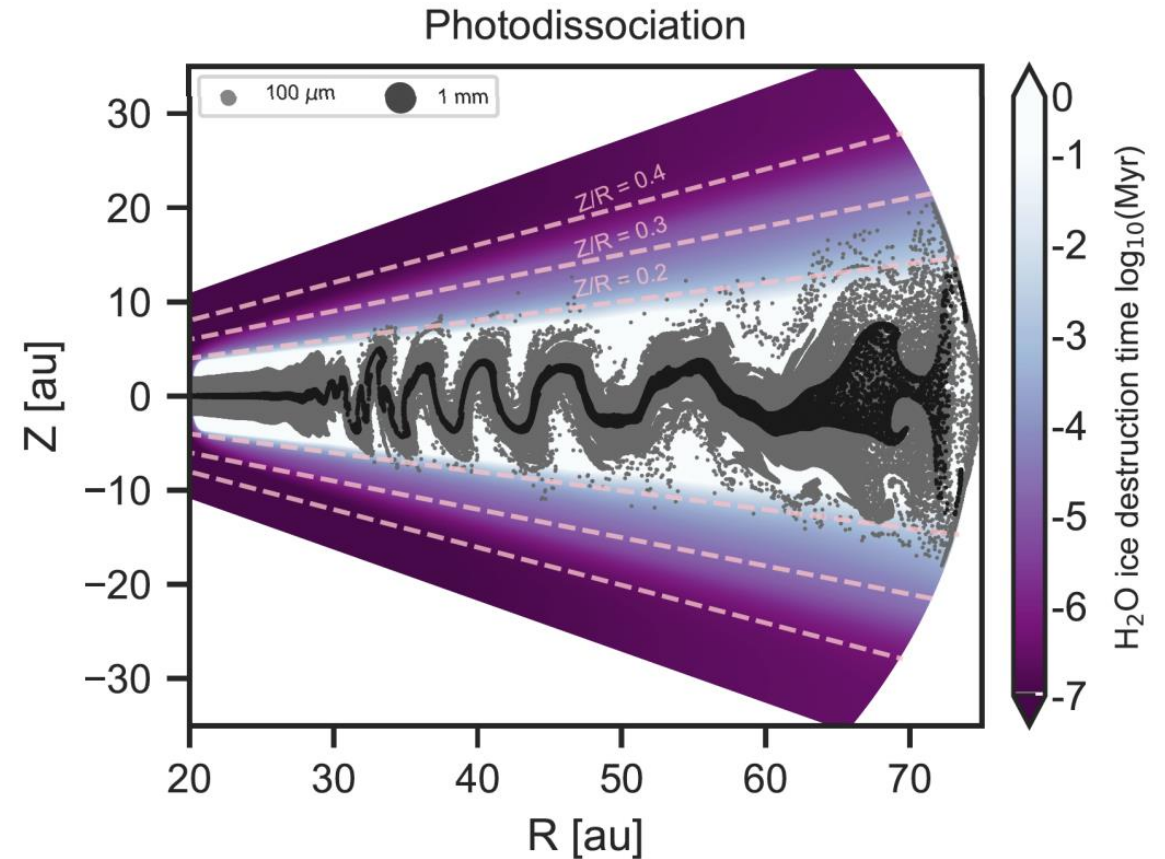
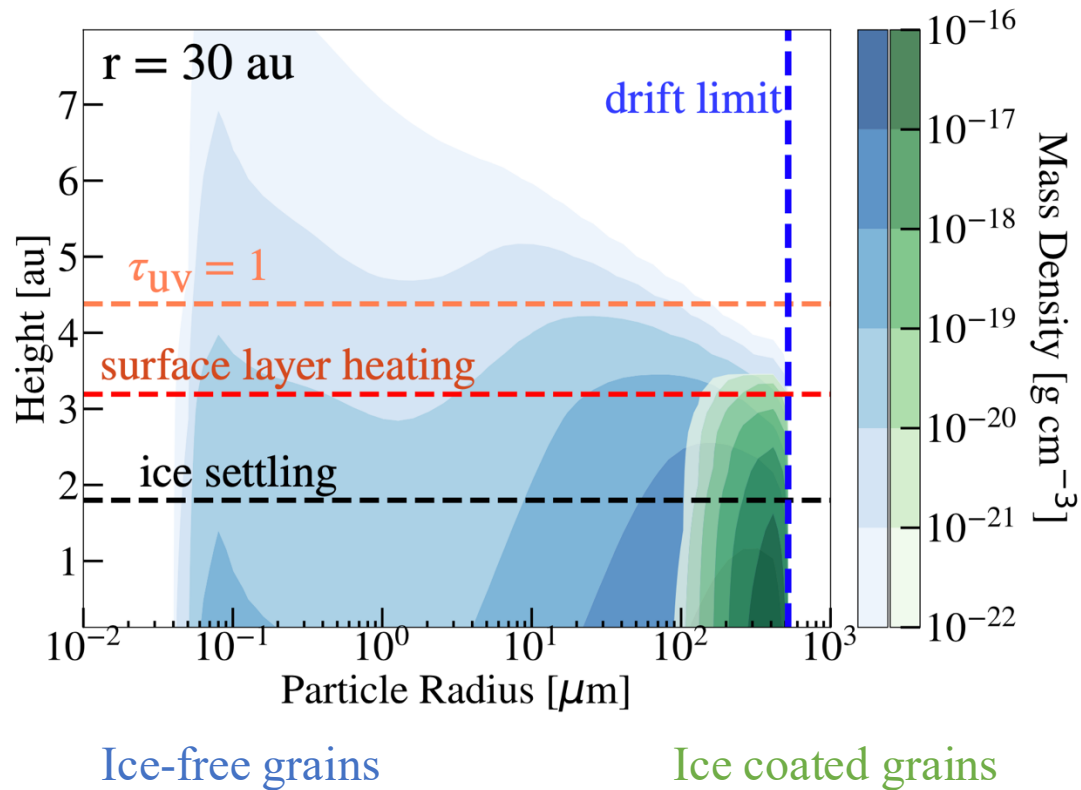
- ~~Constant composition at birth and growth~~
- ~~Growth by planetesimal accretion~~



F_m = rock mass fraction

Thermal and photo-desorption

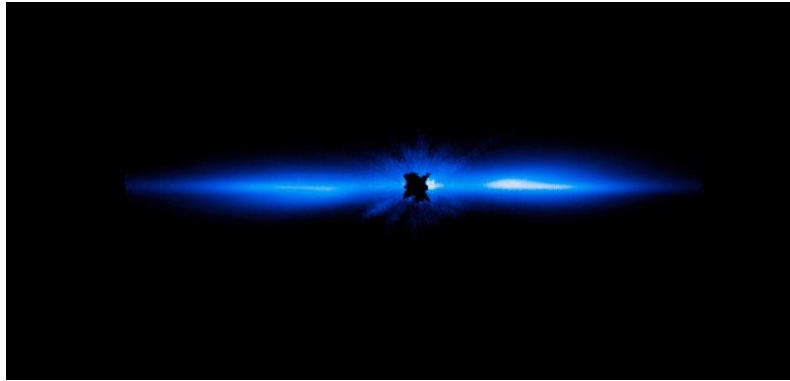
Heating and UV irradiation remove ice on Myr timescales (Harrison & Schoen 1967).



Thermal and photo-desorption

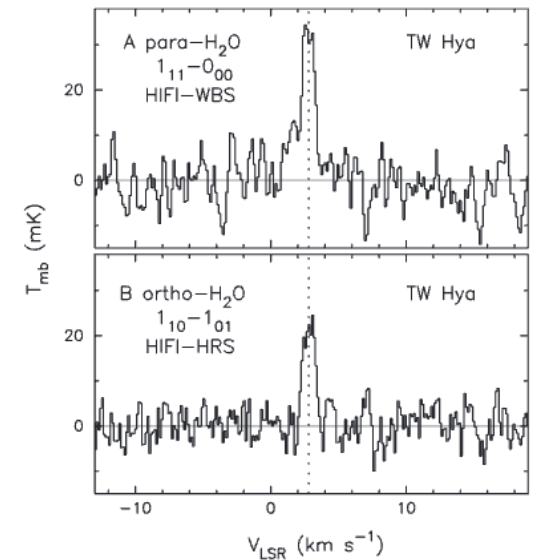
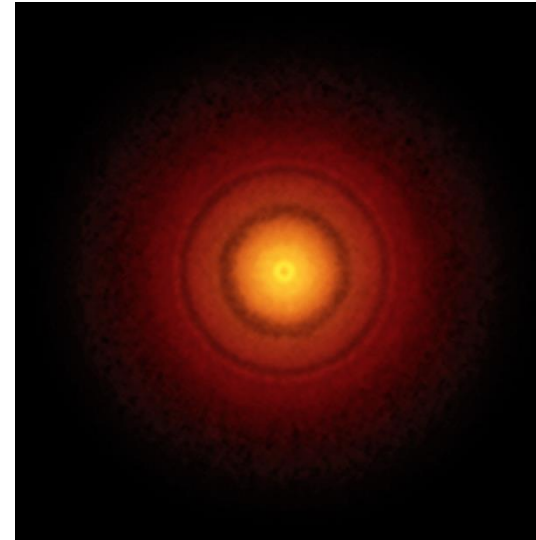
Evidence in disks

Debris disks



Only upper limits for ice in β Pic
(Ballering et al. 2016, Cavallius et al 2019)

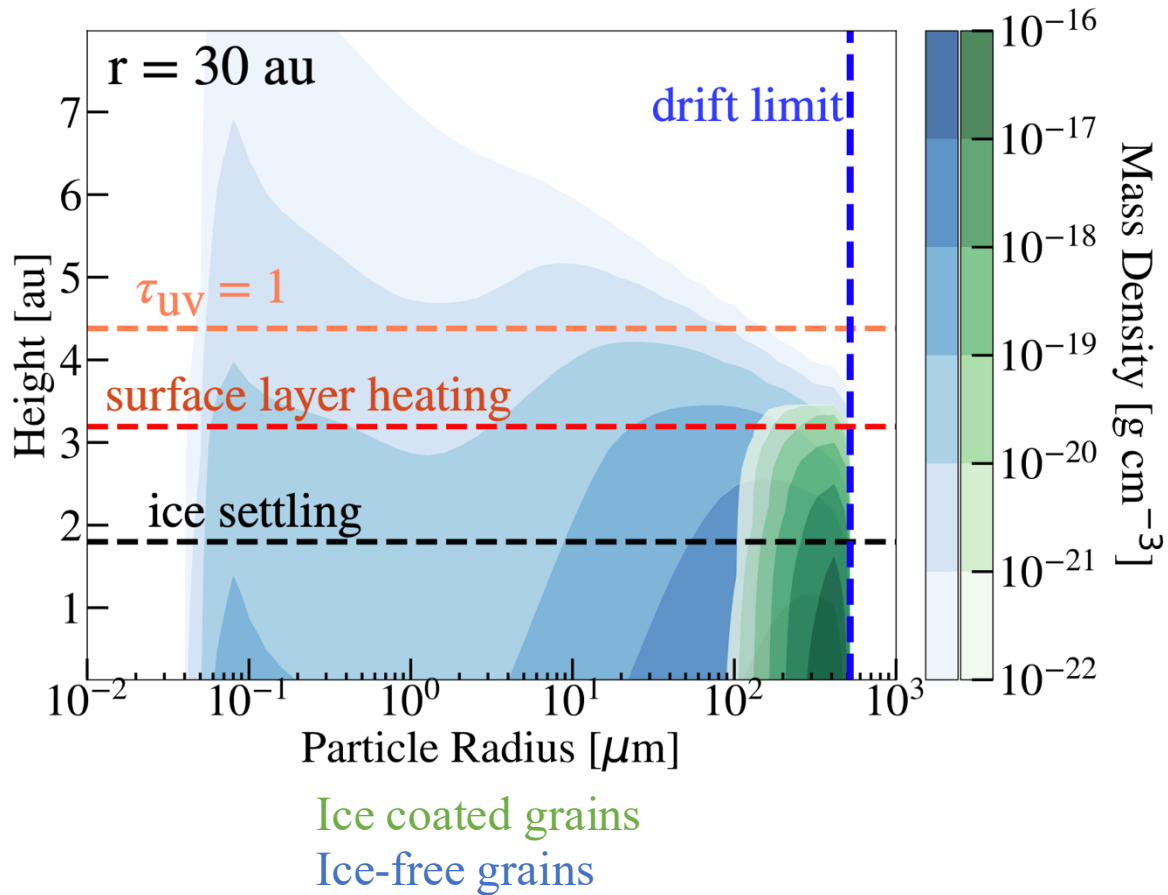
Primordial disks



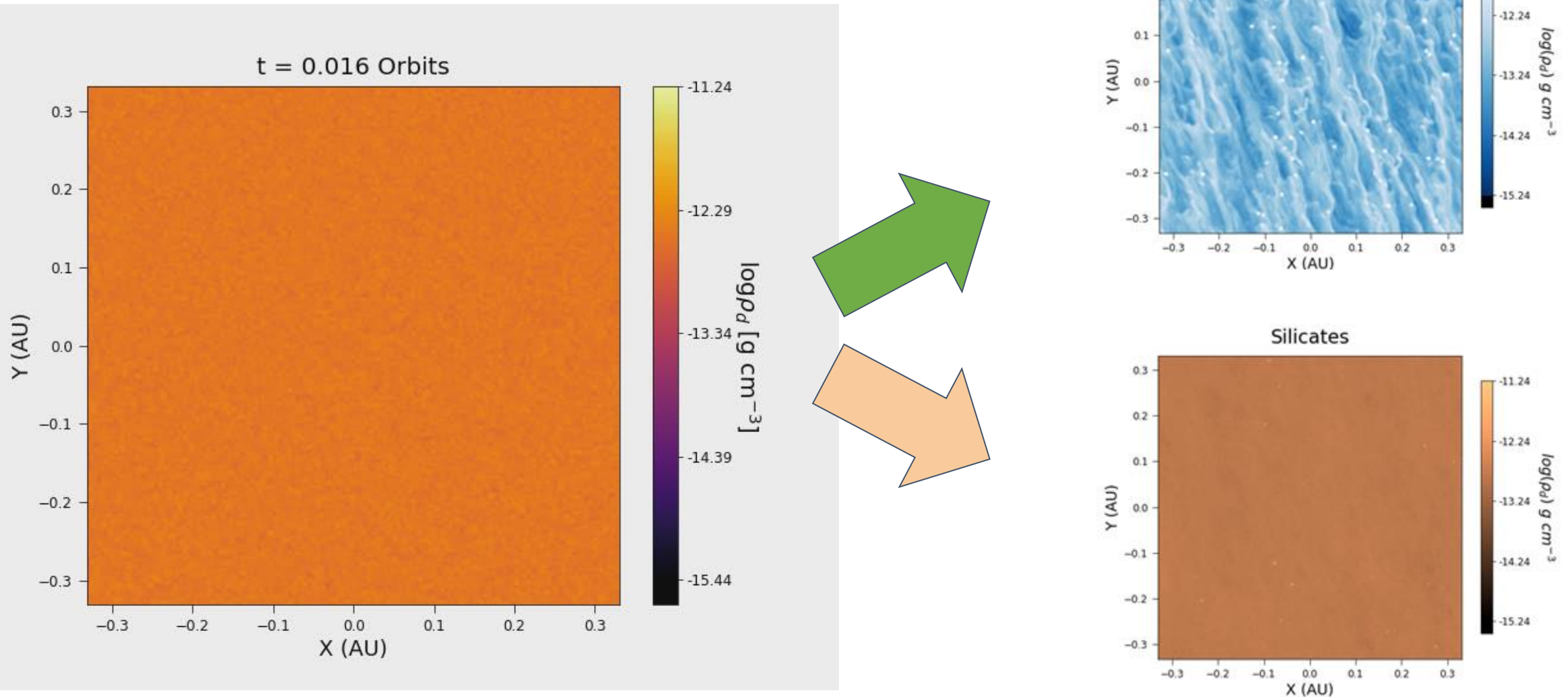
Water vapor observed at large distances for TW Hya
(Hogerheijde et al. 2011, 2012)

Toy model

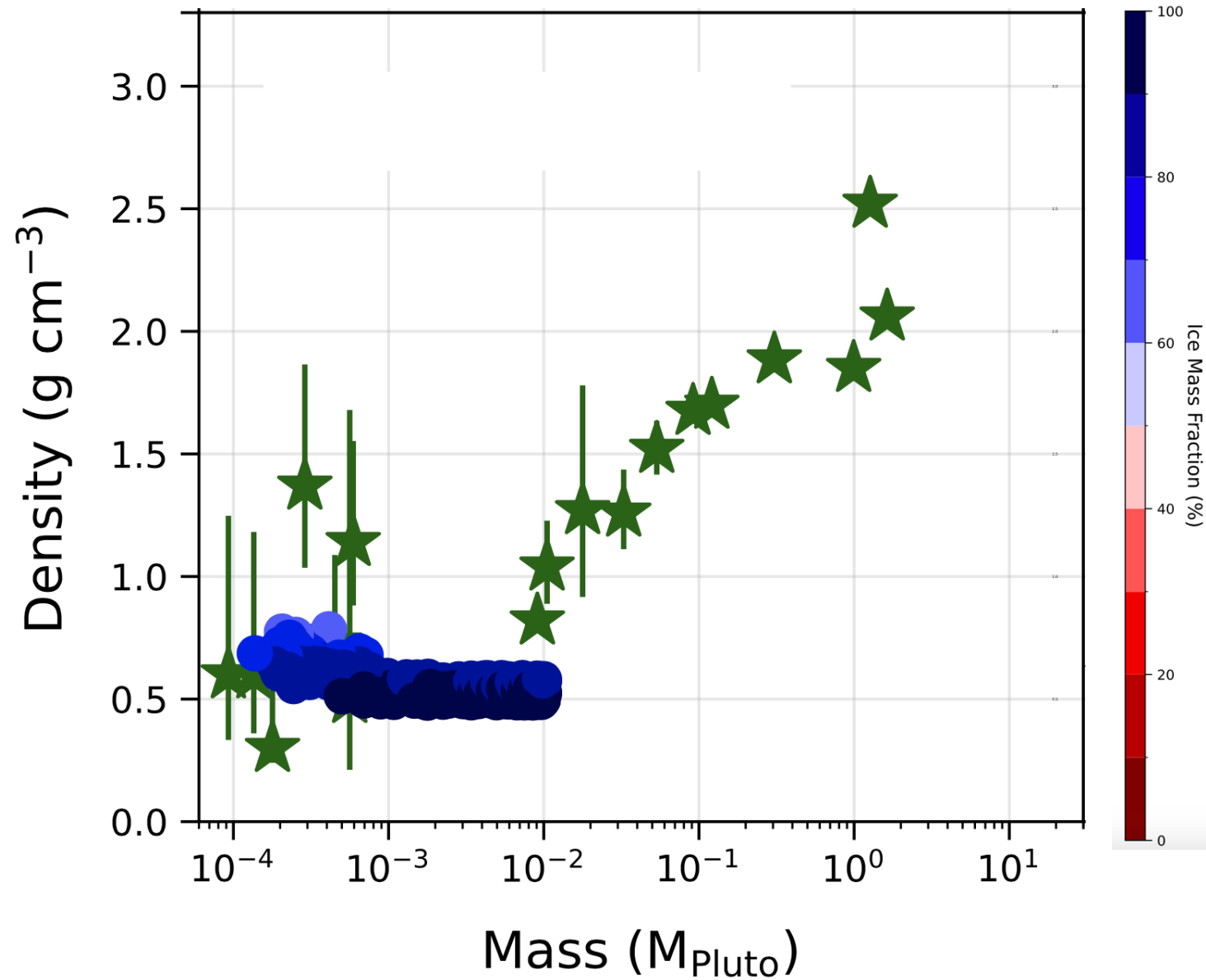
Small grains lofted in the atmosphere lose ice
Big grains are shielded and remain icy.



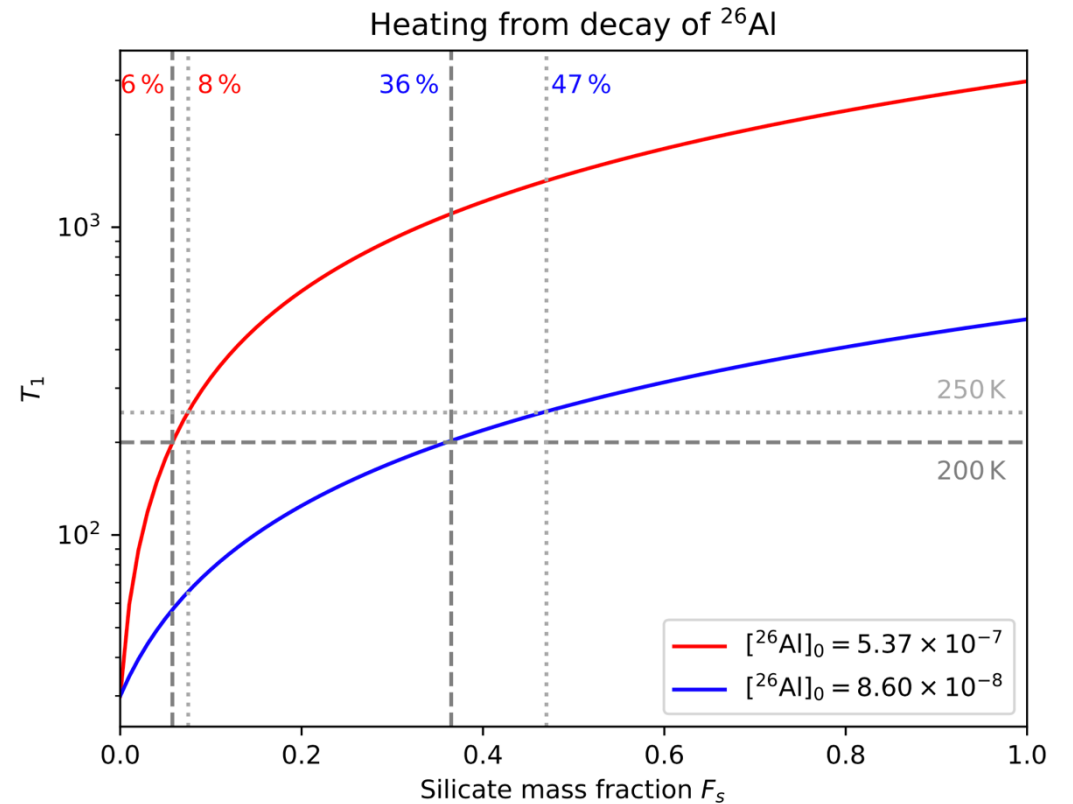
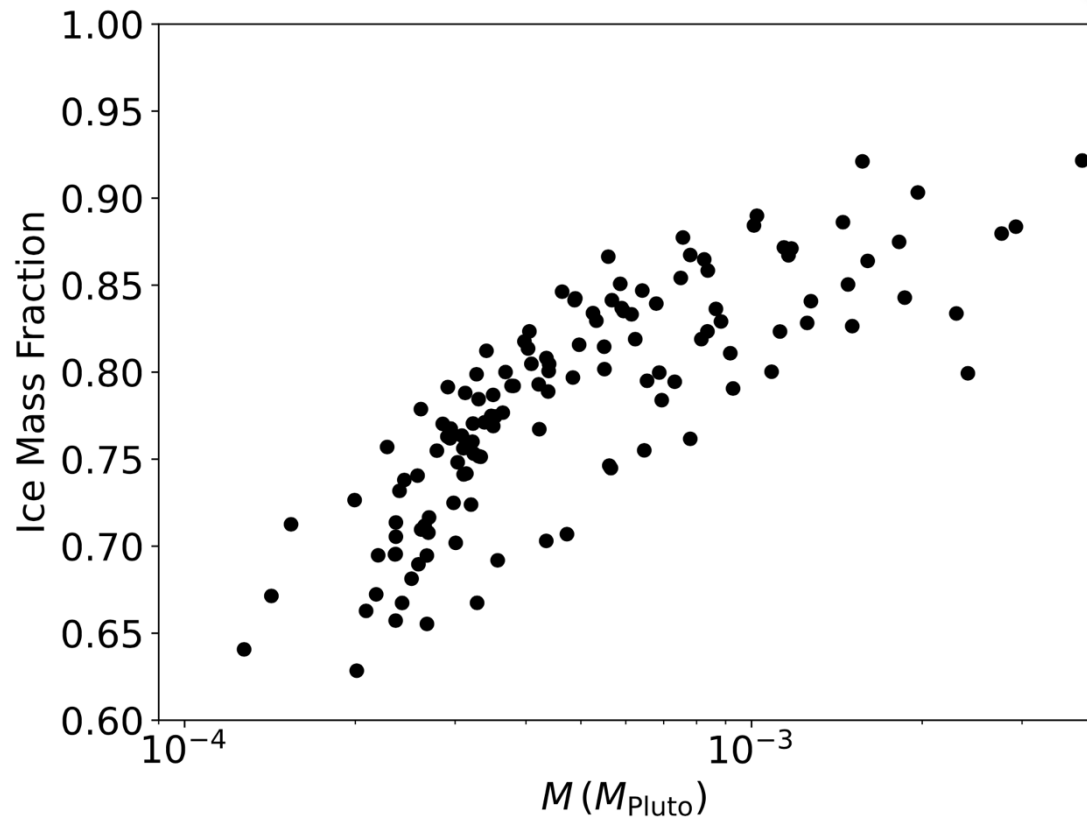
Split into icy and silicate pebbles



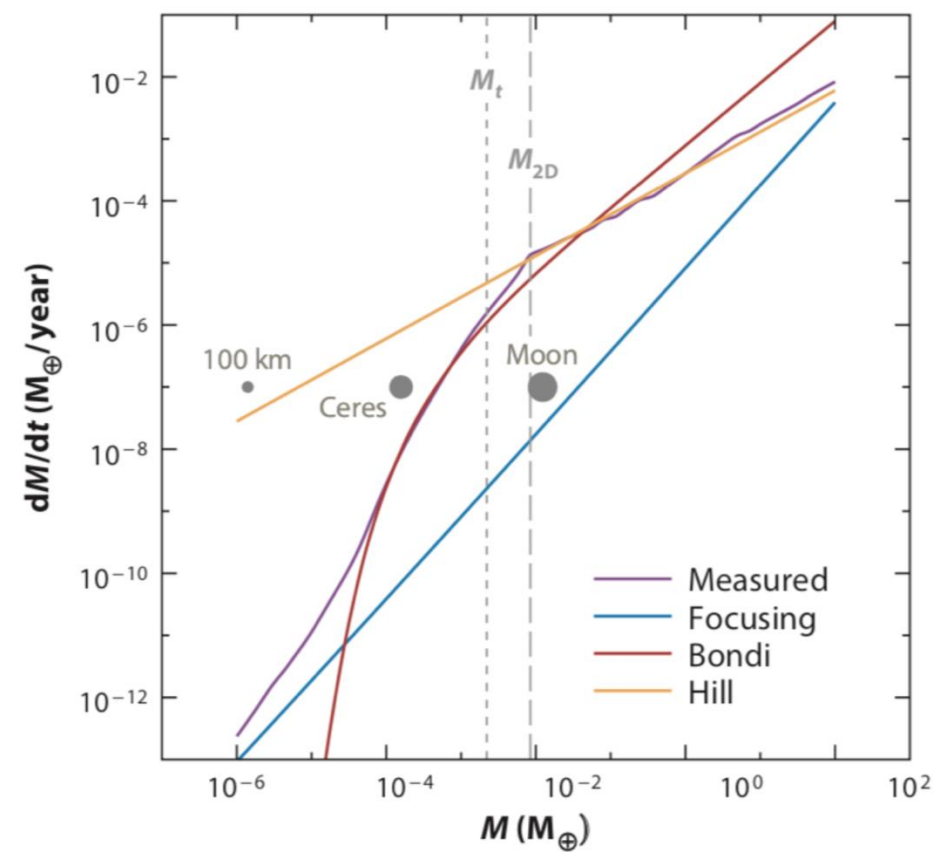
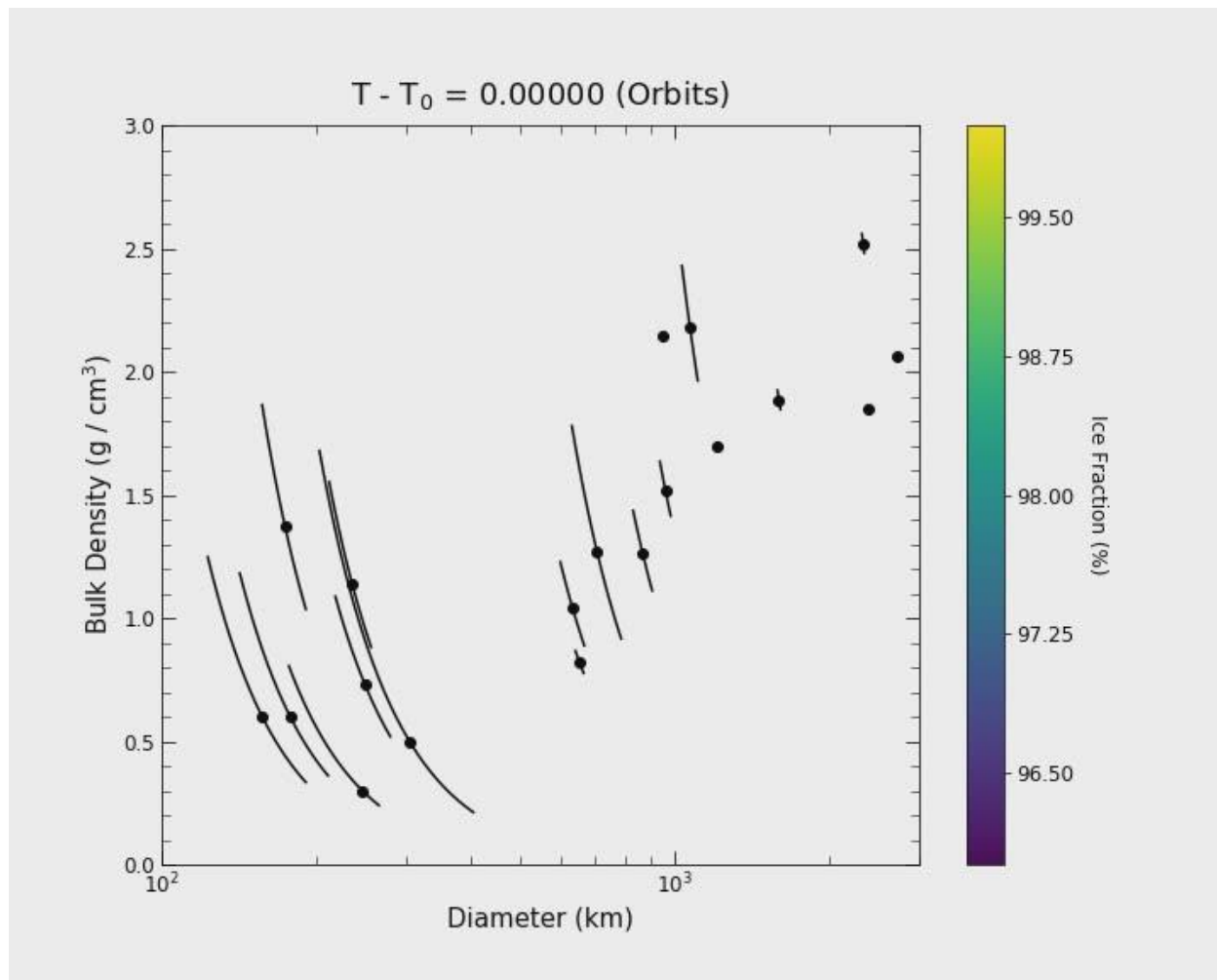
The first planetesimals are icy



^{26}Al binds to silicates: the first planetesimals won't melt



Integrate pebble accretion

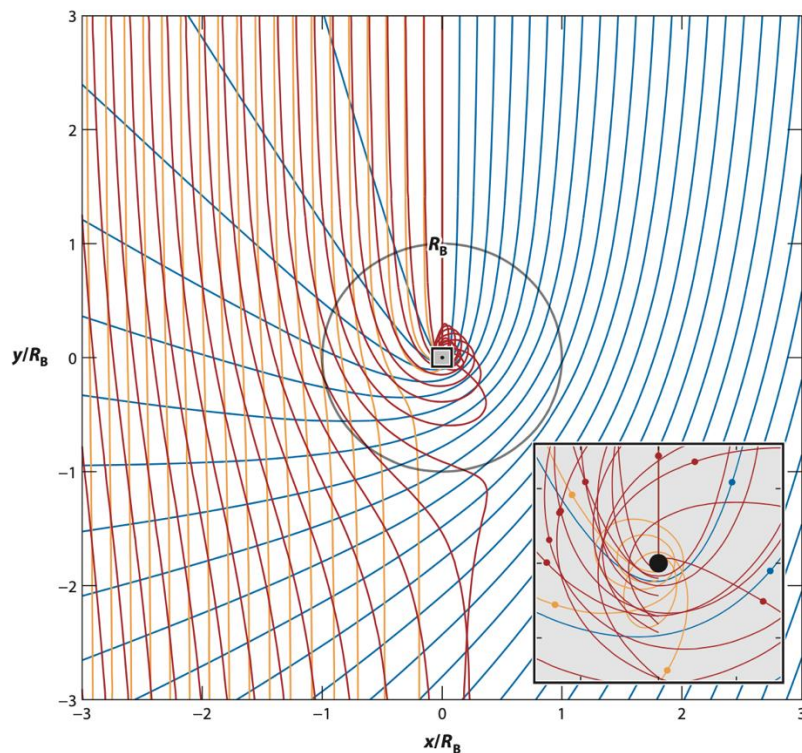


Pebble Accretion: Pebbles of different size accrete differently

"Goldilocks effect" in the Bondi regime

- Large
- Medium
- Small

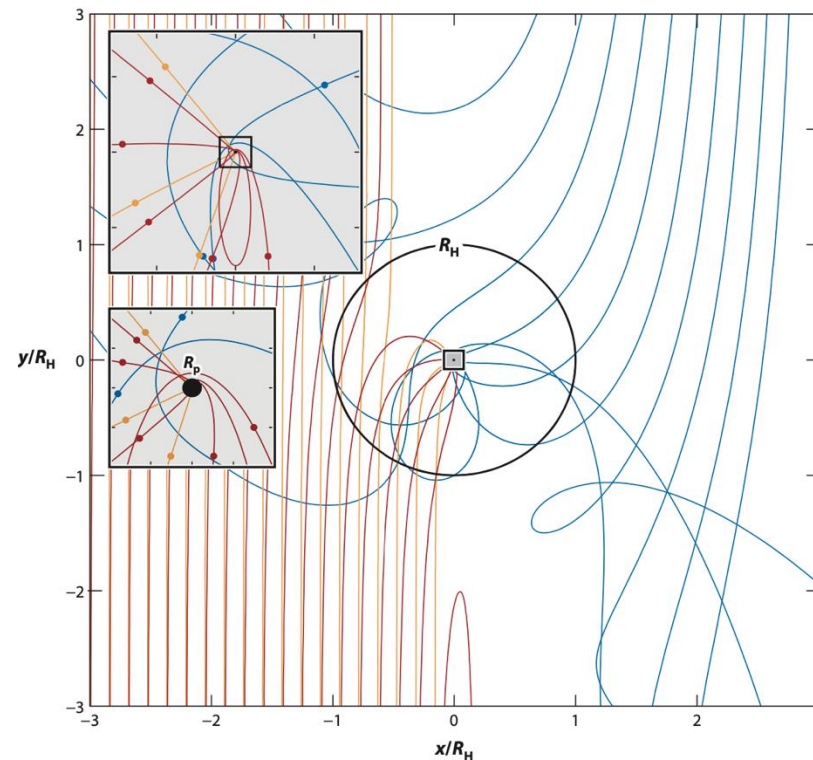
Bondi Regime



Best accreted pebble

Drag time \sim Bondi Time

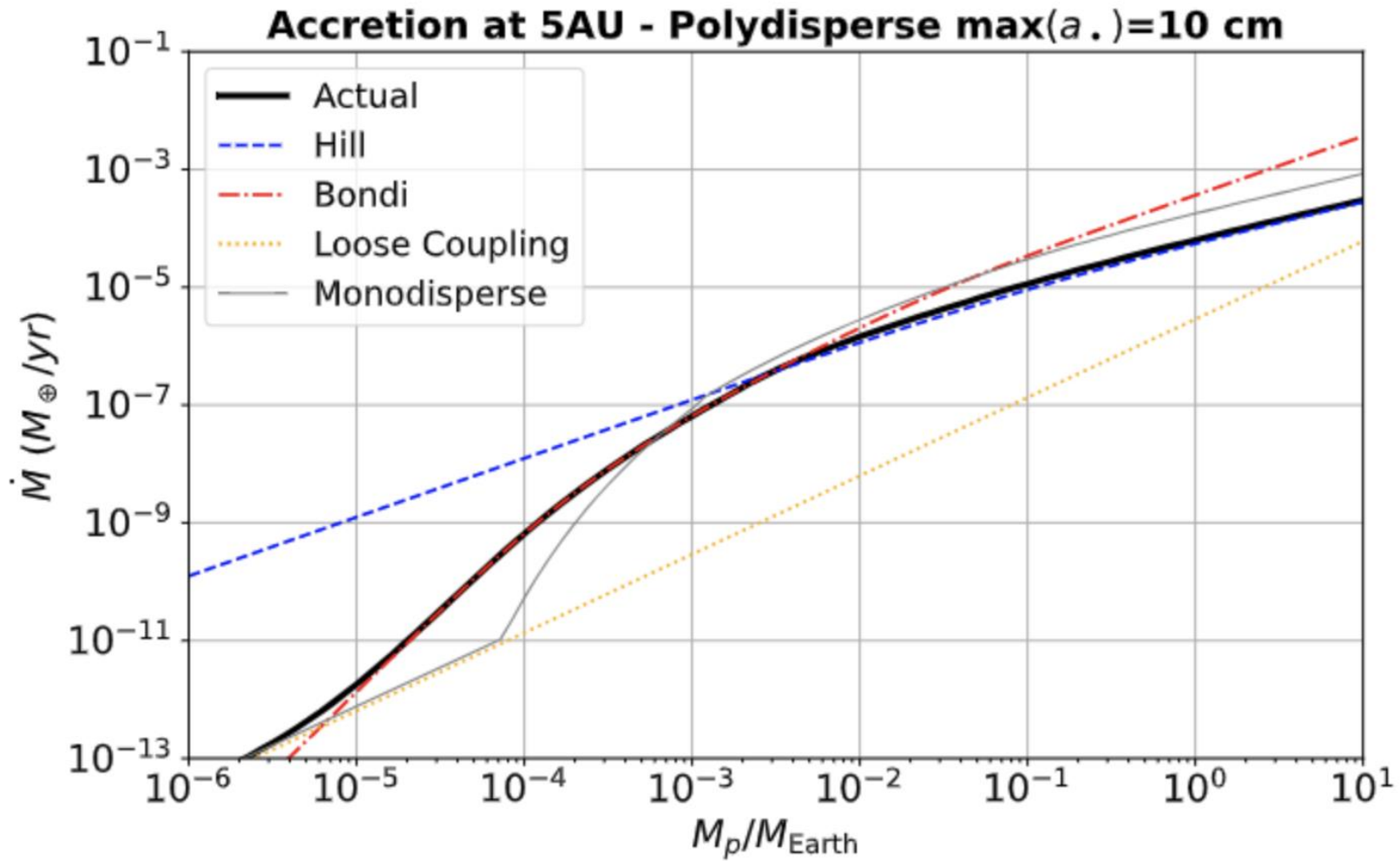
Hill Regime



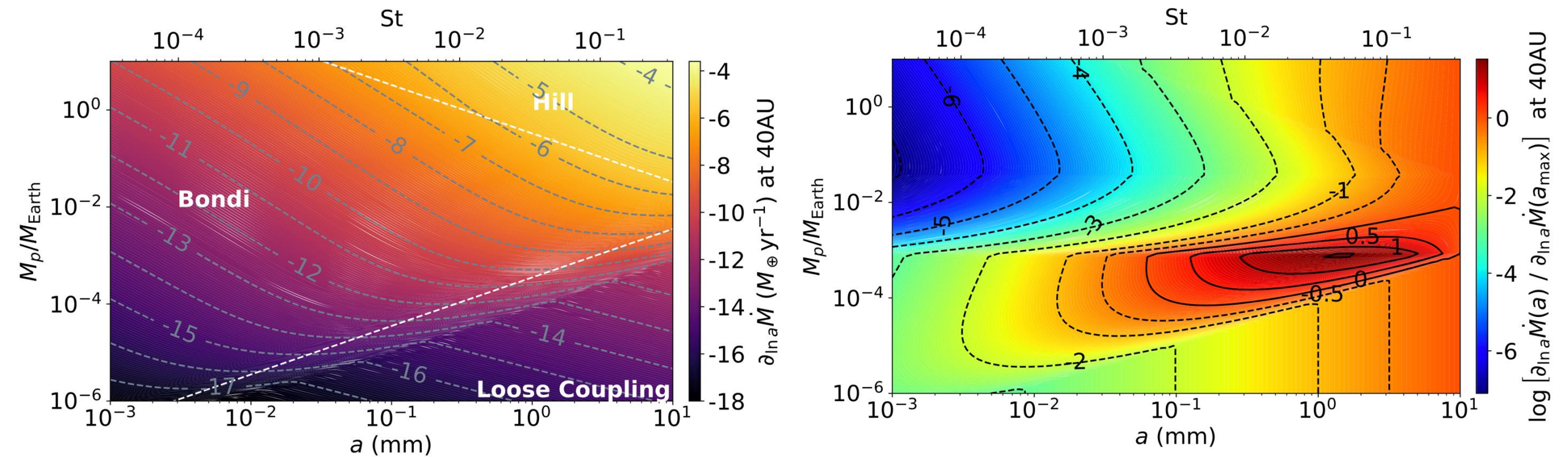
Best accreted pebble

Drag time \sim Orbital Time

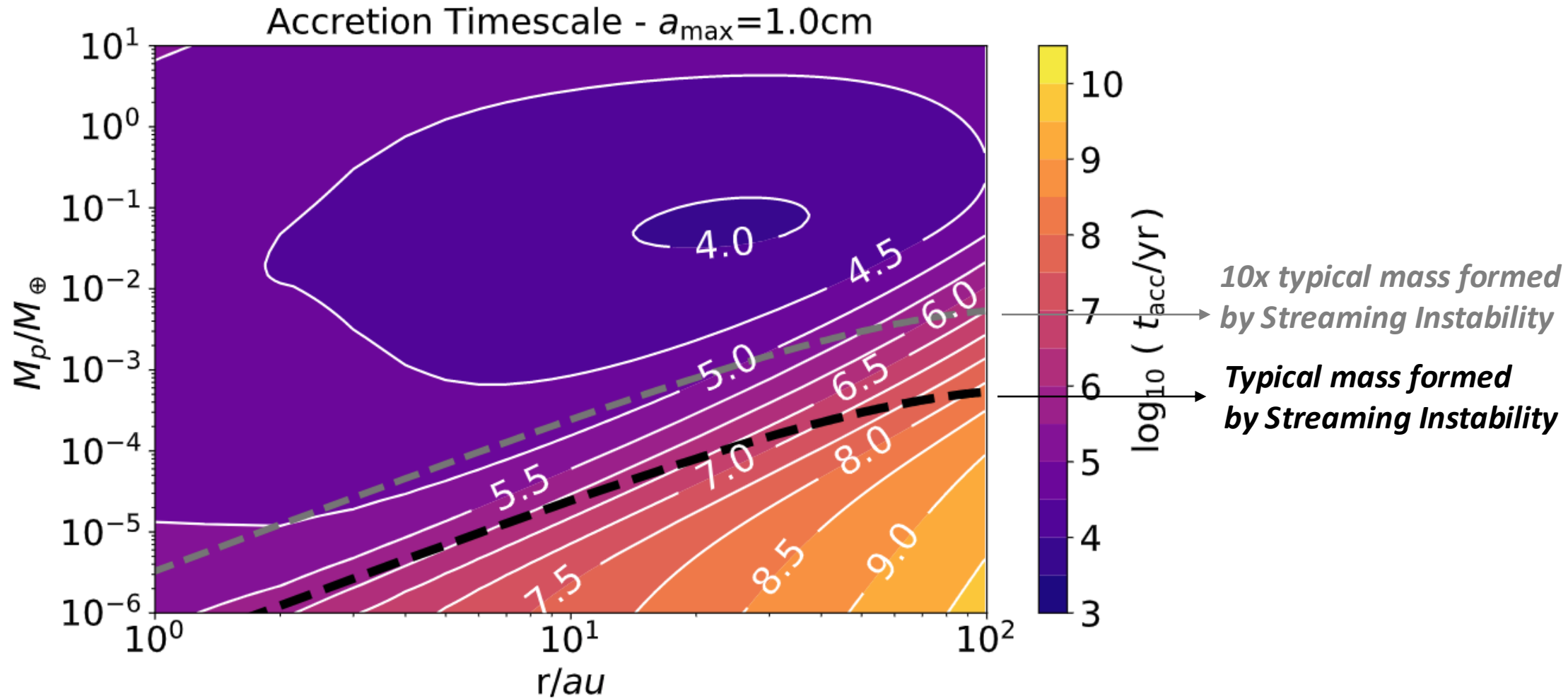
Accretion Rates



Differential Accretion Rates



Accretion Timescales



Polydisperse (Multi-Species) Pebble Accretion

$$\rho_d(a, z) = \int_0^a m(a') F(a', z) da'.$$

$$F(a, z) \equiv f(a) e^{-z^2/2H_d^2},$$

$$f(a) = \frac{3(1-p)Z\Sigma_g}{2^{5/2}\pi^{3/2}H_g\rho_\bullet^{(0)}a_{\max}^{4-k}} \sqrt{1 + a \frac{\pi}{2} \frac{\rho_\bullet(a)}{\Sigma_g\alpha}} a^{-k}.$$

$$S \equiv \frac{1}{\pi R_{\text{acc}}^2} \int_{-R_{\text{acc}}}^{R_{\text{acc}}} 2\sqrt{R_{\text{acc}}^2 - z^2} \exp\left(-\frac{z^2}{2H_d^2}\right) dz,$$

$$W(a) = \frac{3(1-p)Z\Sigma_g}{4\pi\rho_\bullet^{(0)}a_{\max}^{4-k}} a^{-k},$$

$$\hat{R}_{\text{acc}}^{(\text{Bondi})} = \left(\frac{4\tau_f}{t_B}\right)^{1/2} R_B,$$

$$\hat{R}_{\text{acc}}^{(\text{Hill})} = \left(\frac{\text{St}}{0.1}\right)^{1/3} R_H,$$

$$\delta v \equiv \Delta v + \Omega R_{\text{acc}},$$

$$R_{\text{acc}} \equiv \hat{R}_{\text{acc}} \exp[-\chi(\tau_f/t_p)^\gamma],$$

$$\frac{\partial \Sigma_d(a)}{\partial a} \propto a^{-p};$$

$$\rho_\bullet \propto a^{-q}; \quad t_p \equiv \frac{GM_p}{(\Delta v + \Omega R_H)^3}$$

$$\dot{M}(a) = \int_0^a \frac{\partial \dot{M}(a')}{\partial a'} da',$$

$$\frac{\partial \dot{M}(a)}{\partial a} = \pi R_{\text{acc}}^2(a) \delta v(a) S(a) m(a) f(a).$$

$$\dot{M}_{\text{2D, Hill}} = 2 \times 10^{2/3} \Omega R_H^2 \int_0^{a_{\max}} \text{St}(a)^{2/3} m(a) W(a) da.$$

$$\begin{aligned} \dot{M}_{\text{3D, Bondi}} &= \frac{4\pi R_B \Delta v^2}{\Omega} \\ &\times \int_0^{a_{\max}} \text{St} e^{-2\psi} m(a) f(a) \\ &\times \left[1 + 2 \left(\text{St} \frac{\Omega R_B}{\Delta v} \right)^{1/2} e^{-\psi} \right] da, \quad \psi \equiv \chi[\text{St}/(\Omega t_p)]^\gamma. \end{aligned}$$

Analytical theory of polydisperse (multi-species) pebble accretion

Monodisperse (single species)

$$\xi \equiv \left(\frac{R_{\text{acc}}}{2H_d} \right)^2$$

$$\dot{M}_{3D} = \lim_{\xi \rightarrow 0} \dot{M} = \pi R_{\text{acc}}^2 \rho_{d0} \delta v,$$

$$\dot{M}_{2D} = \lim_{\xi \rightarrow \infty} \dot{M} = 2R_{\text{acc}} \Sigma_d \delta v,$$

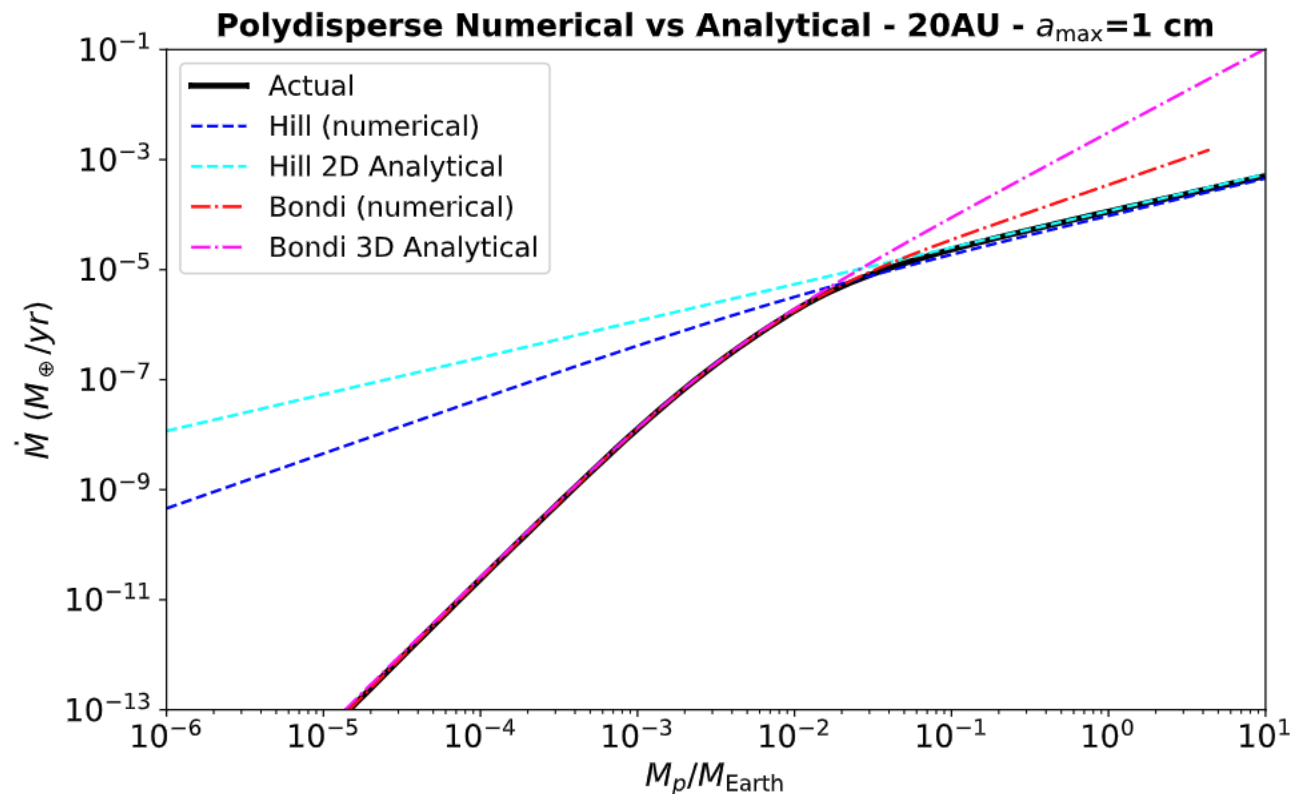
Lambrechts & Johansen (2012)

Polydisperse (multiple species)

$$\dot{M}_{2D, \text{Hill}} = \frac{6(1-p)}{14-5q-3k} \left(\frac{\text{St}_{\text{max}}}{0.1} \right)^{2/3} \Omega R_H^2 Z \Sigma_g.$$

$$\dot{M}_{3D, \text{Bondi}} \approx C_1 \frac{\gamma_l \left(\frac{b_1+1}{s}, j_1 a_{\text{max}}^s \right)}{s j_1^{(b_1+1)/s}} + C_2 \frac{\gamma_l \left(\frac{b_2+1}{s}, j_2 a_{\text{max}}^s \right)}{s j_2^{(b_2+1)/s}} + C_3 \frac{\gamma_l \left(\frac{b_3+1}{s}, j_3 a_{\text{max}}^s \right)}{s j_3^{(b_3+1)/s}} + C_4 \frac{\gamma_l \left(\frac{b_4+1}{s}, j_4 a_{\text{max}}^s \right)}{s j_4^{(b_4+1)/s}},$$

Lyra et al. (2023)



Analytical Solutions

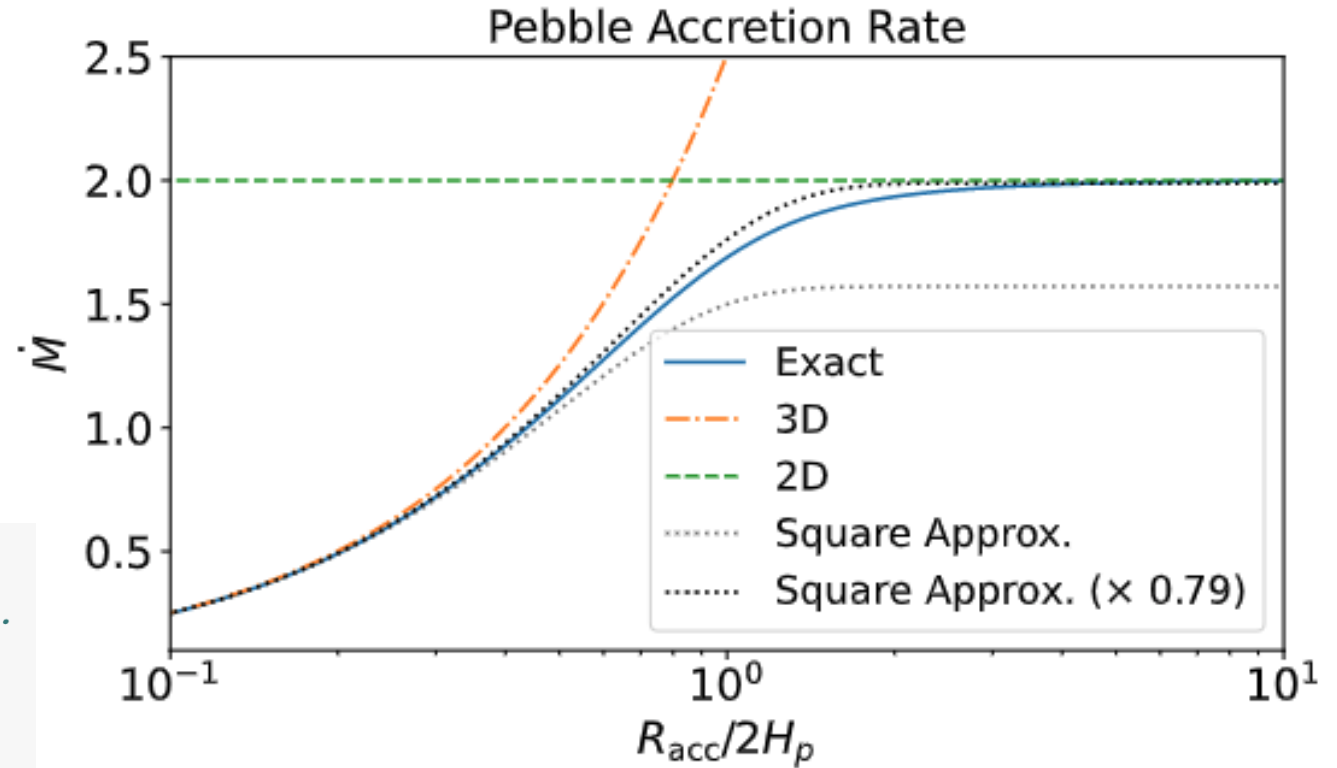
$$\dot{M} = \pi R_{\text{acc}}^2 \rho_{d0} S \delta v.$$

$$S \equiv \frac{1}{\pi R_{\text{acc}}^2} \int_{-R_{\text{acc}}}^{R_{\text{acc}}} 2 \sqrt{R_{\text{acc}}^2 - z^2} \exp\left(-\frac{z^2}{2H_d^2}\right) dz,$$

$$S = e^{-\xi} [I_0(\xi) + I_1(\xi)], \quad \xi \equiv \left(\frac{R_{\text{acc}}}{2H_d}\right)^2$$

```
y = (x/2)**2
# Modified Bessel function of the first kind of real order.
I0 = sp.special.iv(0, y)
I1 = sp.special.iv(1, y)

Sint = np.exp(-y) * (I0 + I1)
rho_int = rhop * Sint
Mdot = pi*r**2 * rho_int * deltav
```



Analytical Solutions

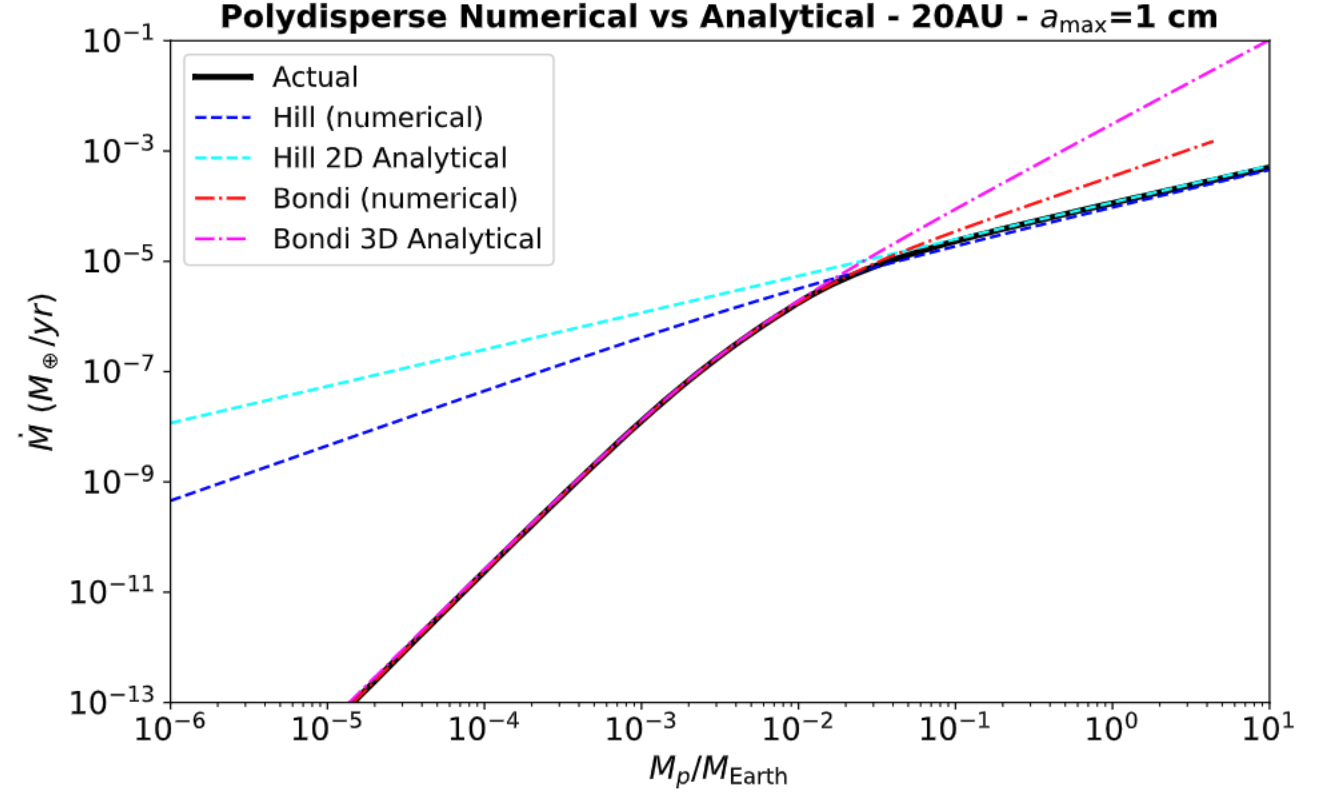
$$\dot{M}_{2D,Hill} = \frac{6(1-p)}{14-5q-3k} \left(\frac{St_{\max}}{0.1} \right)^{2/3} \Omega R_H^2 Z \Sigma_g.$$

$$\dot{M}_{3D,Bondi} \approx C_1 \frac{\gamma_l \left(\frac{b_1+1}{s}, j_1 a_{\max}^s \right)}{s j_1^{(b_1+1)/s}} + C_2 \frac{\gamma_l \left(\frac{b_2+1}{s}, j_2 a_{\max}^s \right)}{s j_2^{(b_2+1)/s}} + C_3 \frac{\gamma_l \left(\frac{b_3+1}{s}, j_3 a_{\max}^s \right)}{s j_3^{(b_3+1)/s}} + C_4 \frac{\gamma_l \left(\frac{b_4+1}{s}, j_4 a_{\max}^s \right)}{s j_4^{(b_4+1)/s}}.$$

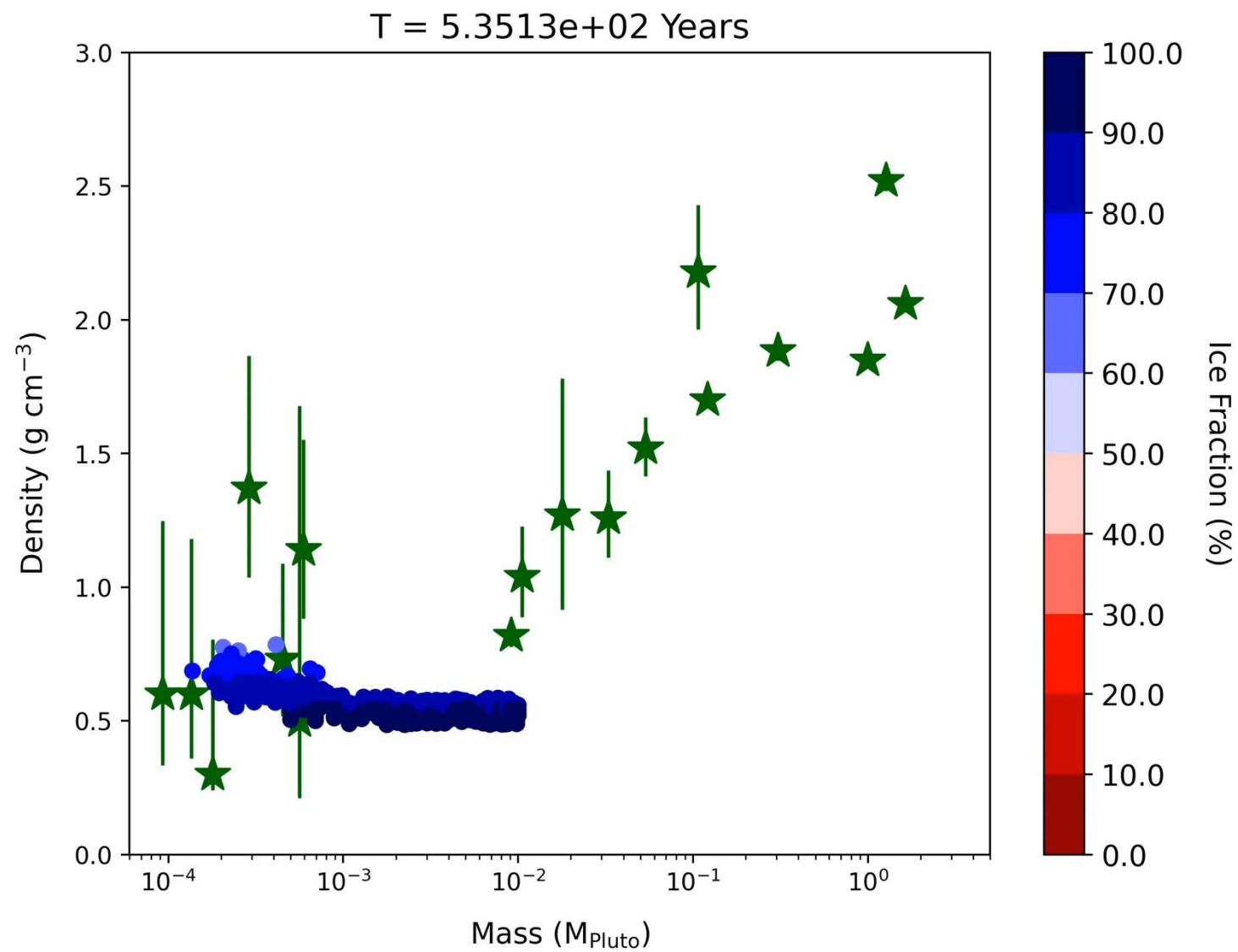
```
gamma11 = sp.special.gammainc((b1+1)/s, j1*a**s)*sp.special.gamma((b1+1)/s)
gamma12 = sp.special.gammainc((b2+1)/s, j2*a**s)*sp.special.gamma((b2+1)/s)
gamma13 = sp.special.gammainc((b3+1)/s, j3*a**s)*sp.special.gamma((b3+1)/s)
gamma14 = sp.special.gammainc((b4+1)/s, j4*a**s)*sp.special.gamma((b4+1)/s)
```

```
G1 = C1*gamma11/s/j1**((b1+1)/s)
G2 = C2*gamma12/s/j2**((b2+1)/s)
G3 = C3*gamma13/s/j3**((b3+1)/s)
G4 = C4*gamma14/s/j4**((b4+1)/s)
```

```
Mbondi3d = G1 + G2 + G3 + G4
```

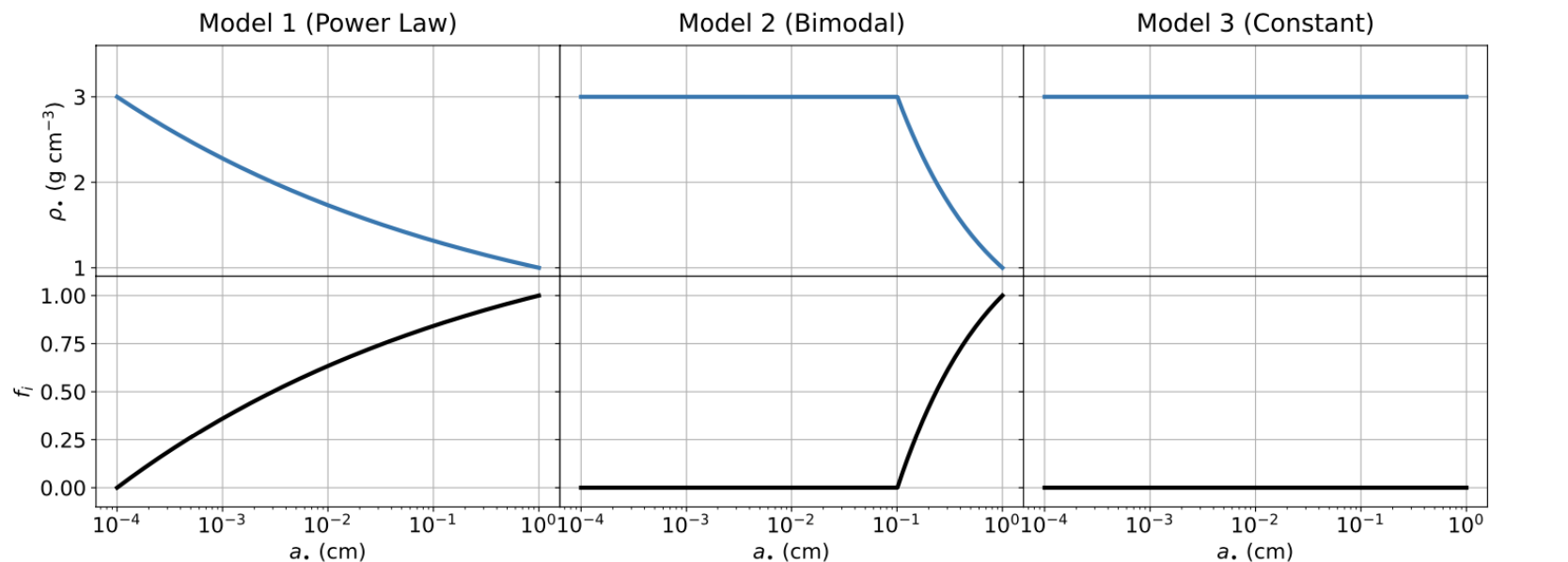


Growing Pluto by silicate pebble accretion

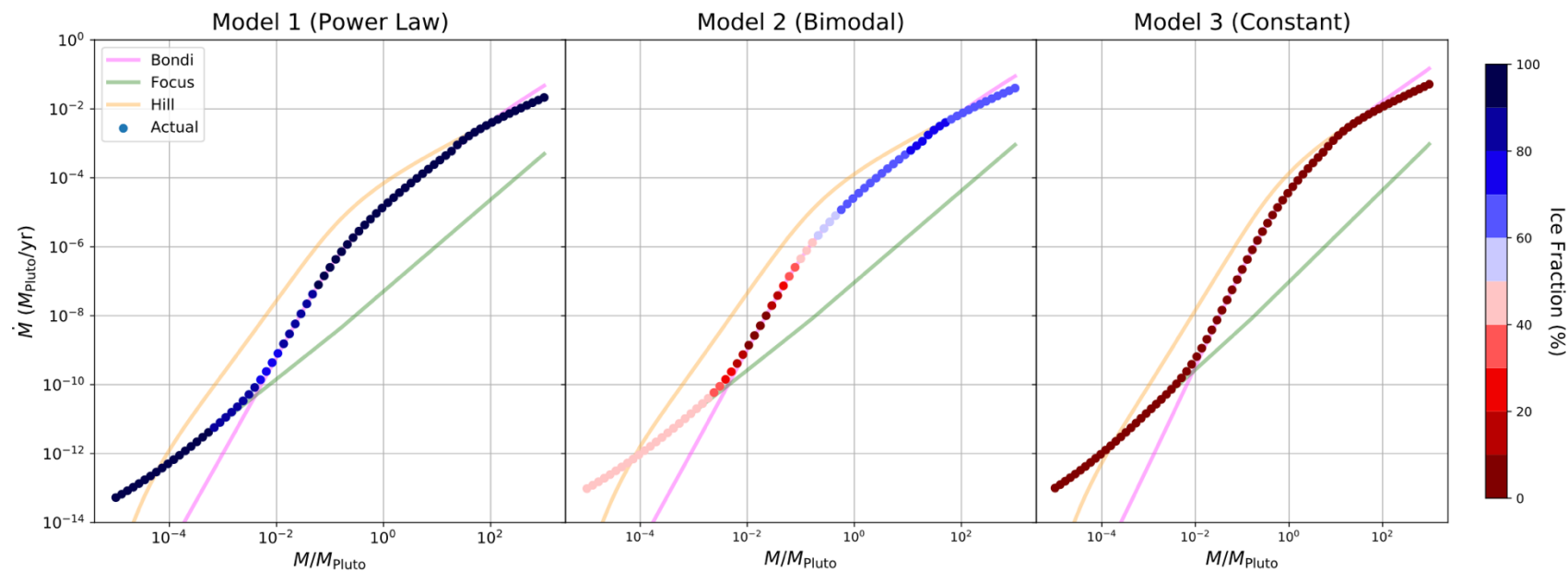


Pebble Internal Density

Ice Volume Fraction

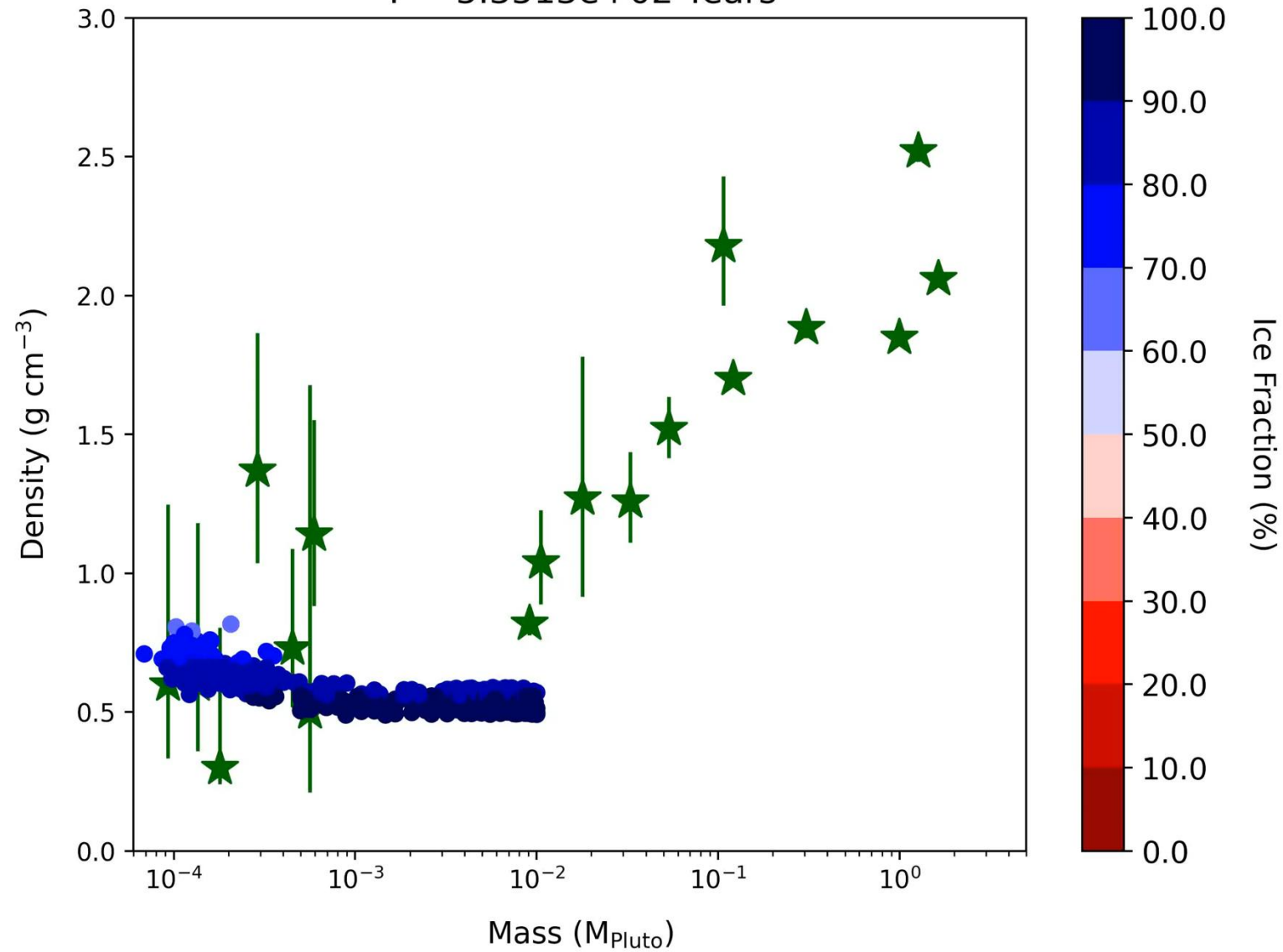


Mass Accretion rate

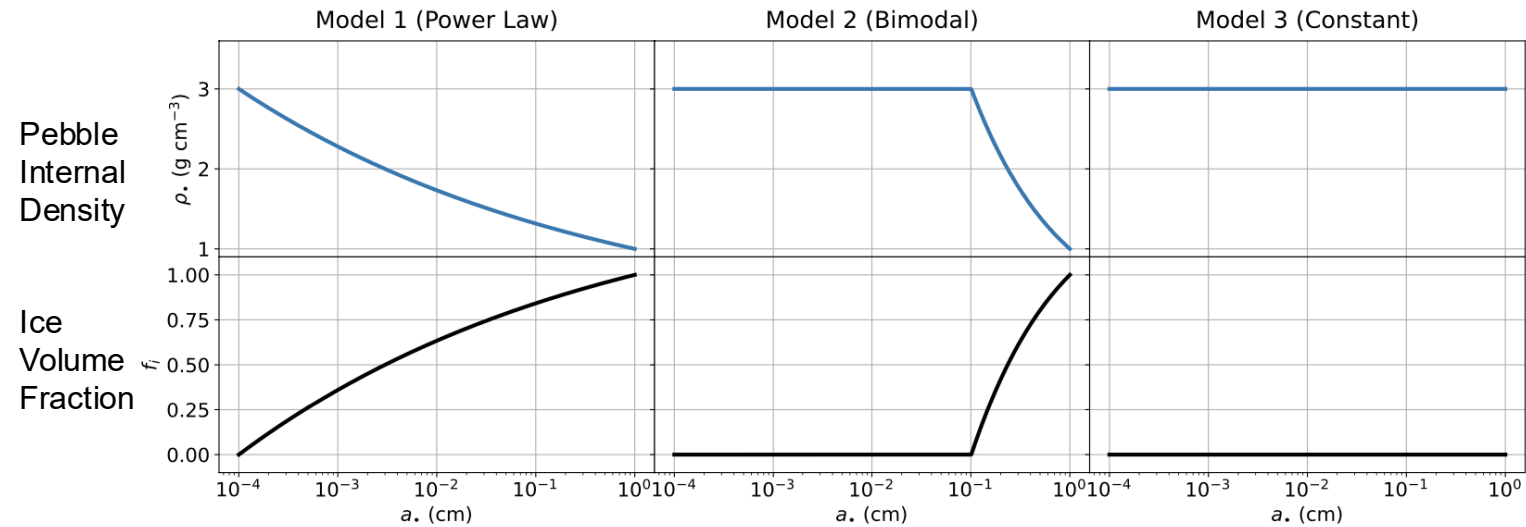


Growing Pluto by silicate pebble accretion

$T = 5.3513 \times 10^2$ Years



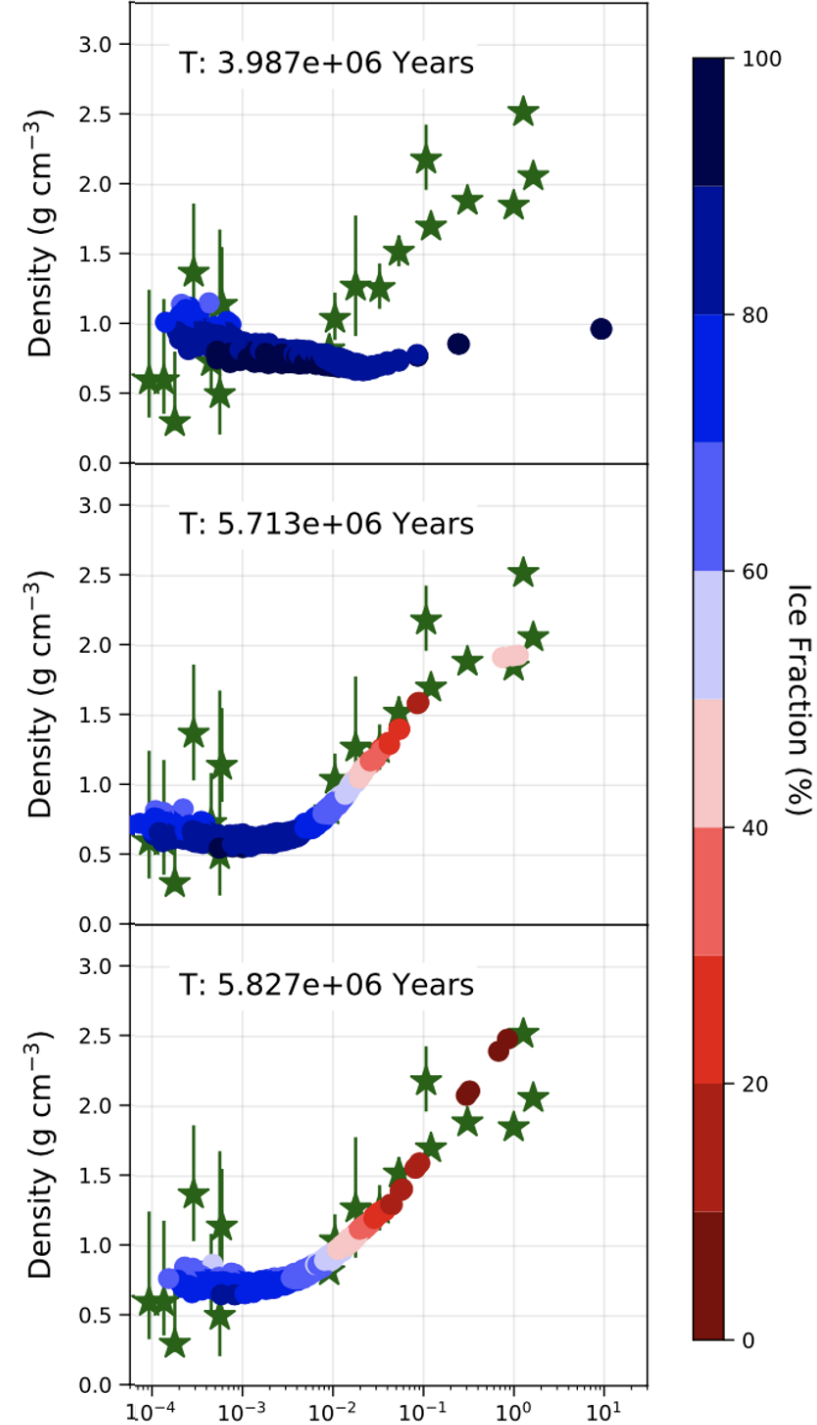
Resulting Densities vs Mass relations



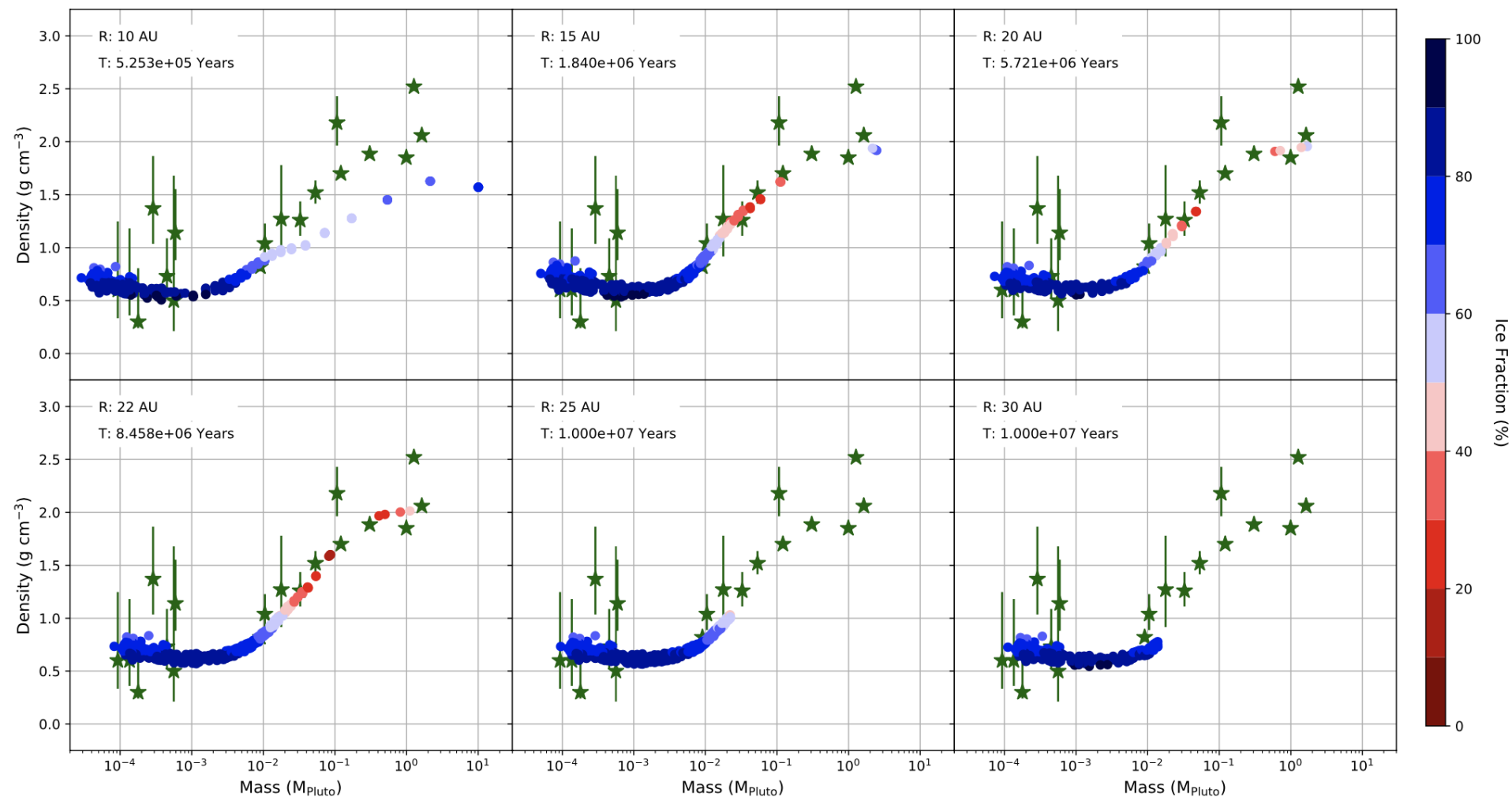
Model 1 (Power Law)

Model 2 (Bimodal)

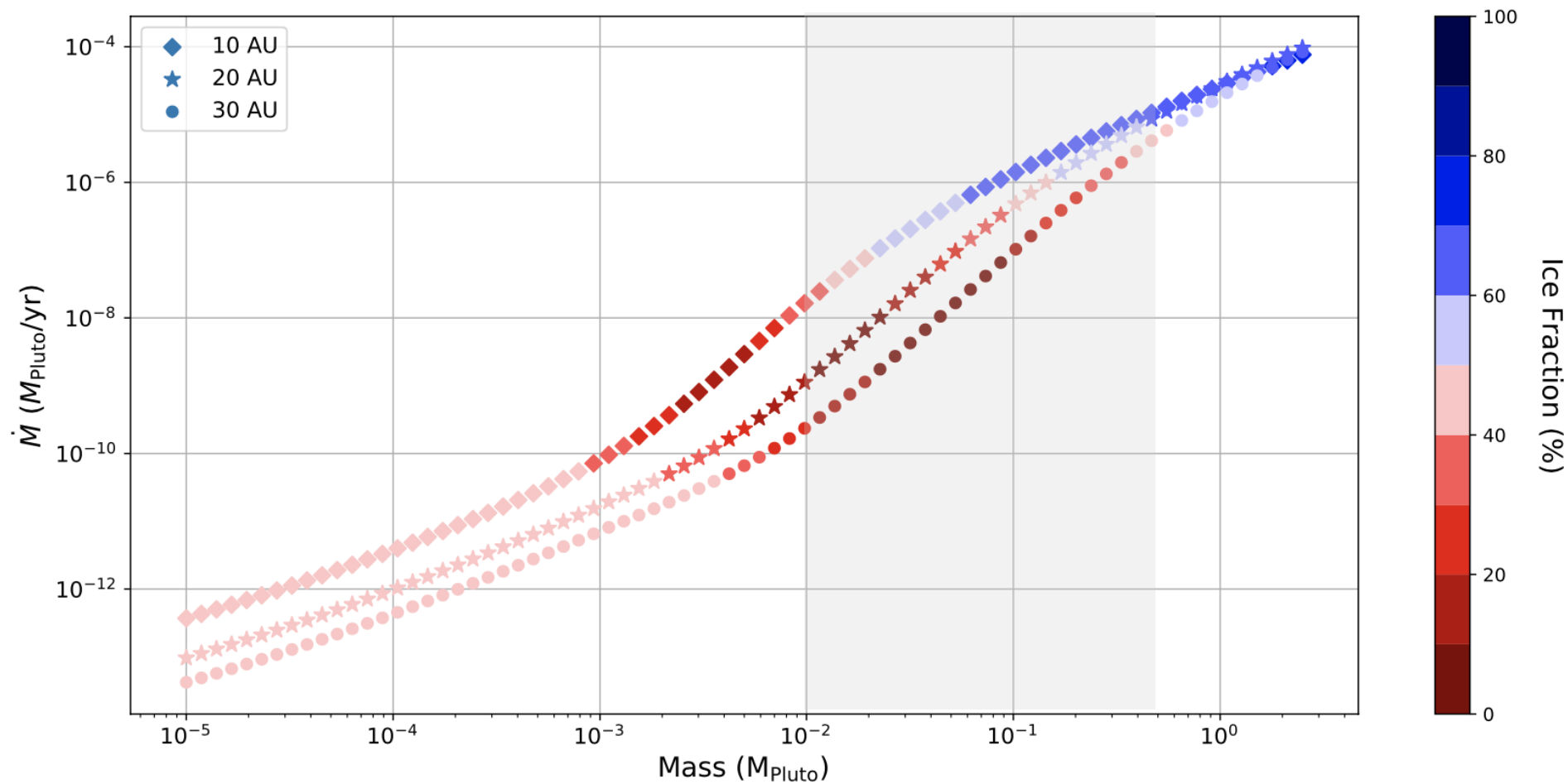
Model 3 (Constant)



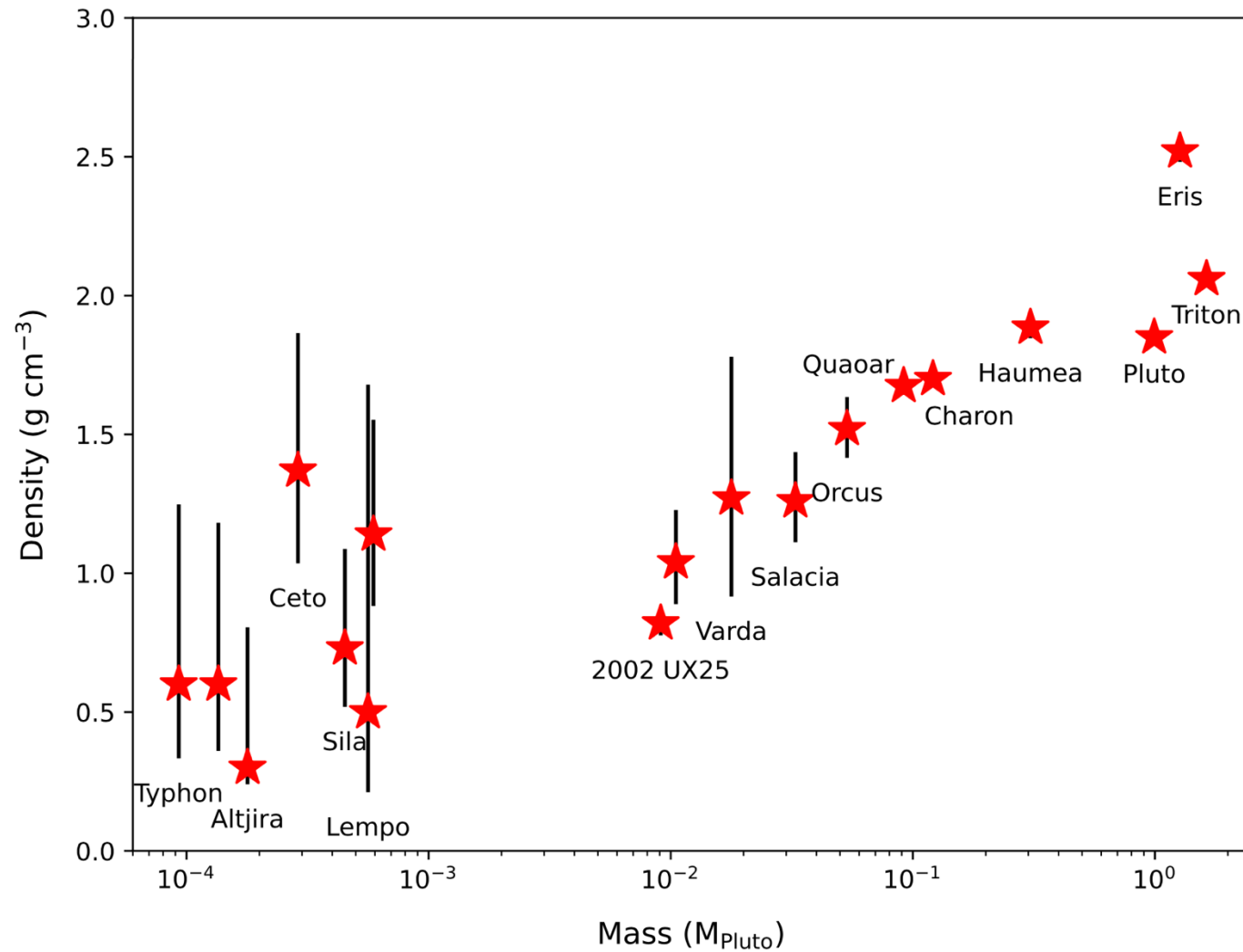
Distance Range 15 - 25AU



The window of silicate accretion

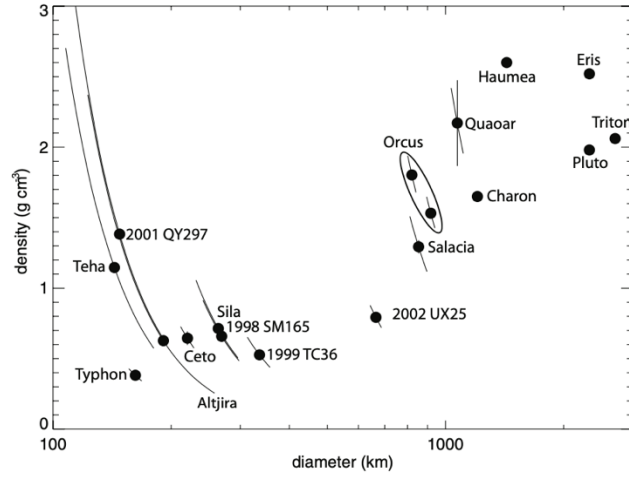


Where are the missing Kuiper Belt binaries?

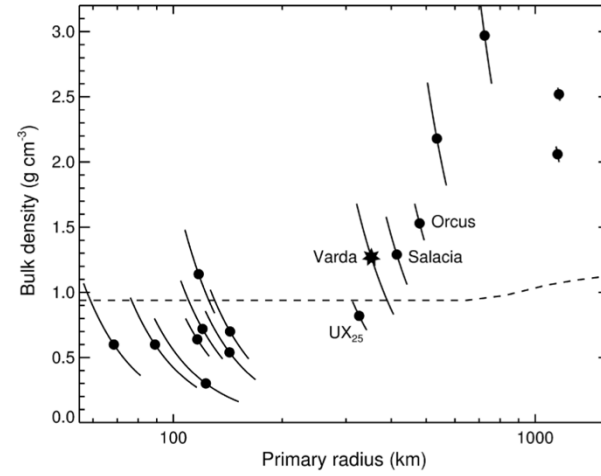


Mass gap

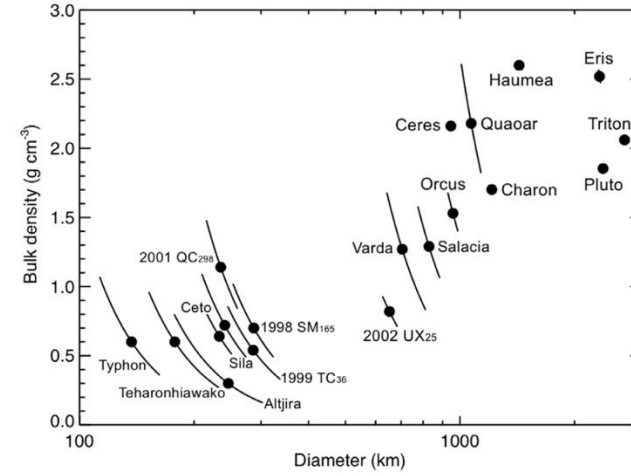
Brown 2013



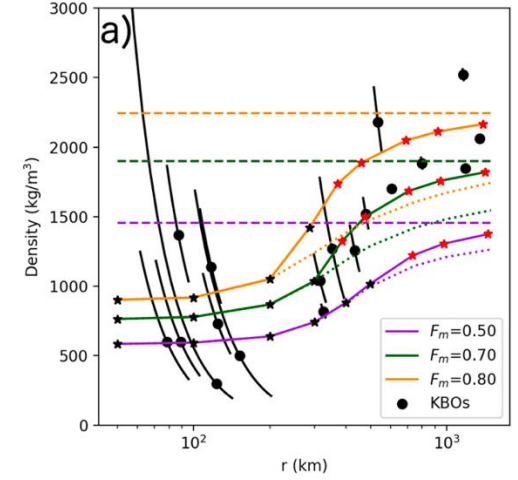
Grundy et al. 2015



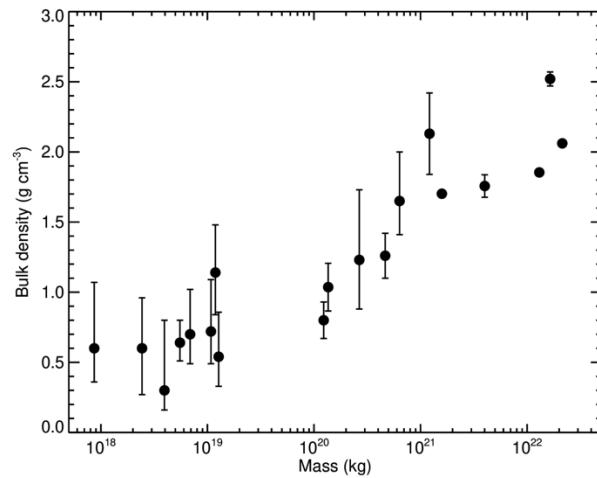
McKinnon et al. 2017



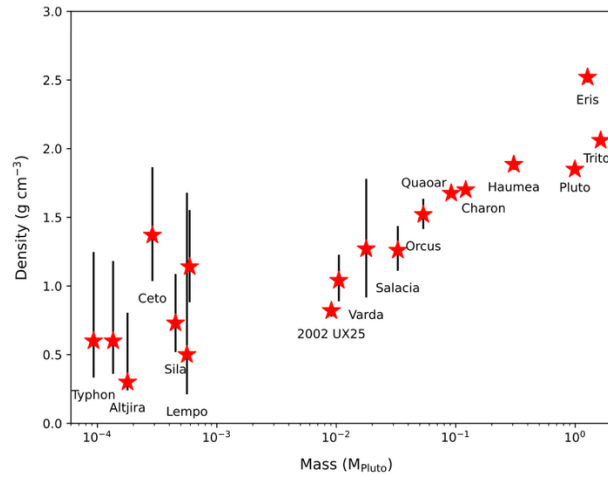
Bierson & Nimmo 2019



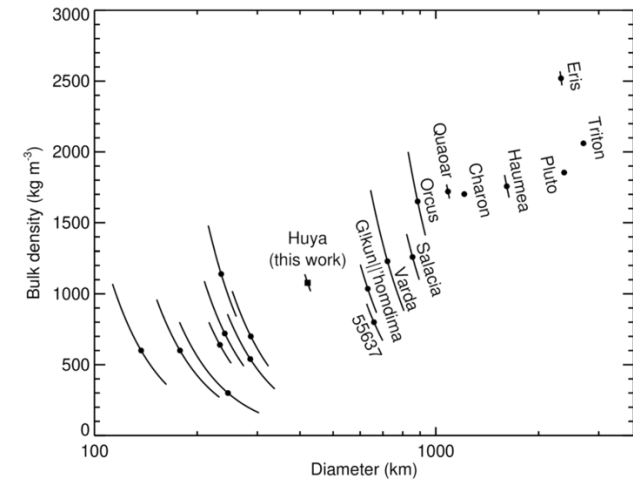
Noll et al. 2020

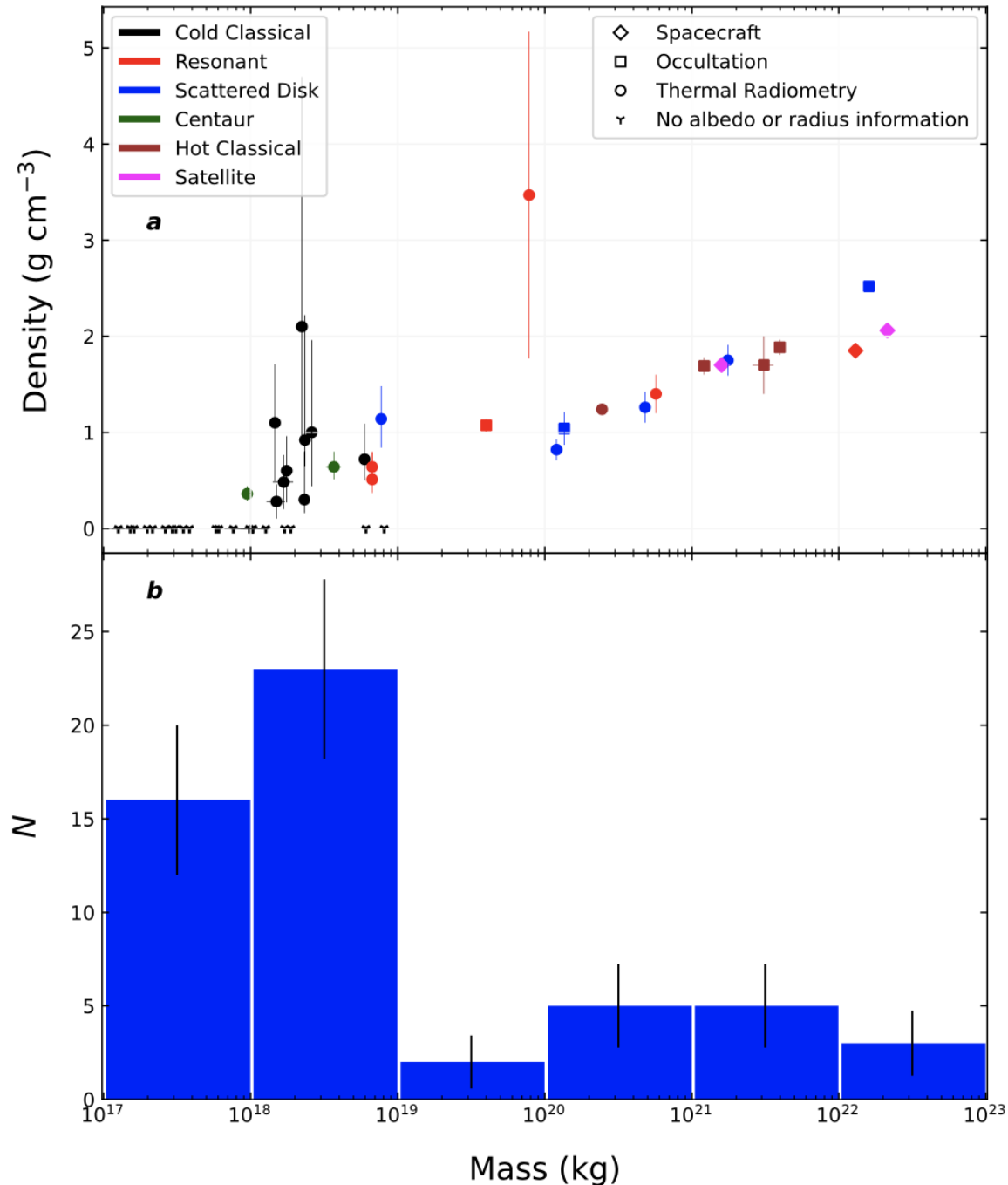


Canas et al. 2024



Rommel et al. 2025





Where are the missing Kuiper Belt binaries?

- Gap between 10^{19} and 10^{20} kg (10^{-3} and 10^{-2} Pluto masses)
- Population difference
 - Cold Classicals are on low-mass side of the gap
- Likely not small number statistics
- *Low mass side is the high-mass end of the planetesimal initial mass function.*

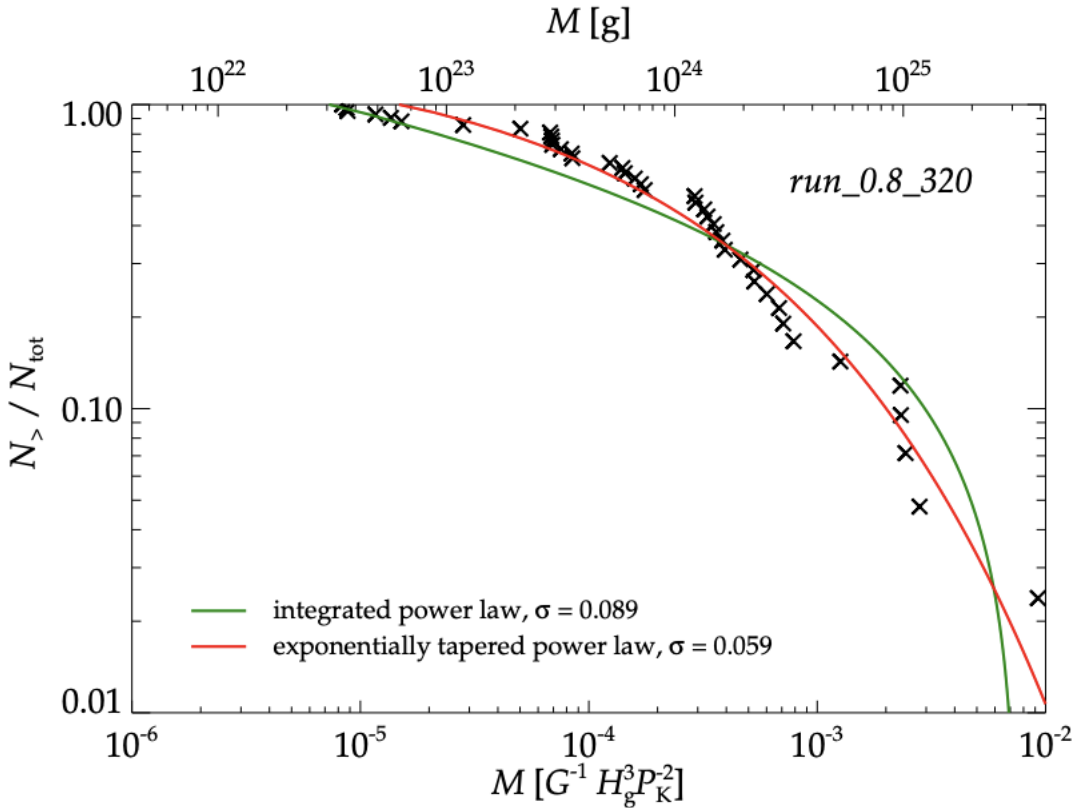
Low-mass end: consistent with high-mass end of Streaming Instability

A&A 597, A69 (2017)
DOI: 10.1051/0004-6361/201629561
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Astronomy
&
Astrophysics

Initial mass function of planetesimals formed by the streaming instability

Urs Schäfer^{1,2}, Chao-Chin Yang², and Anders Johansen²



Schafer et al. (2017)

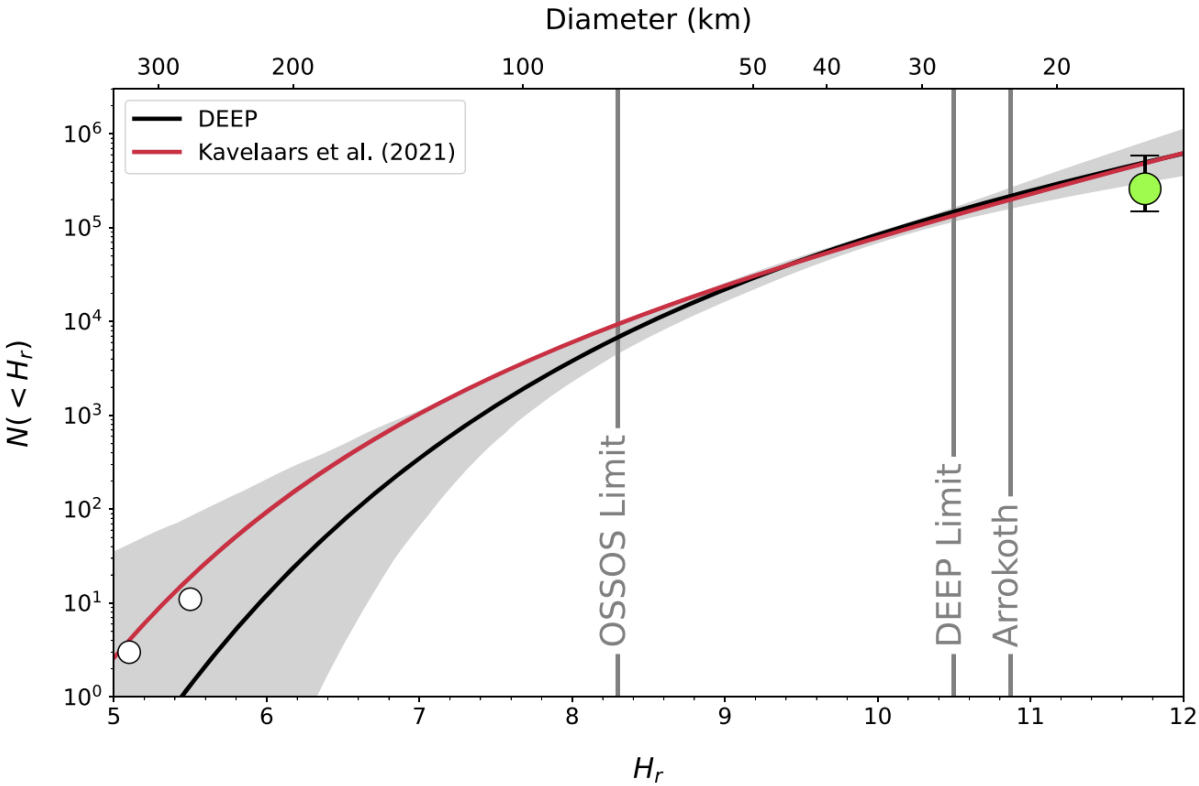
THE PLANETARY SCIENCE JOURNAL, 5:50 (18pp), 2024 February
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<https://doi.org/10.3847/PSJ/ad1528>



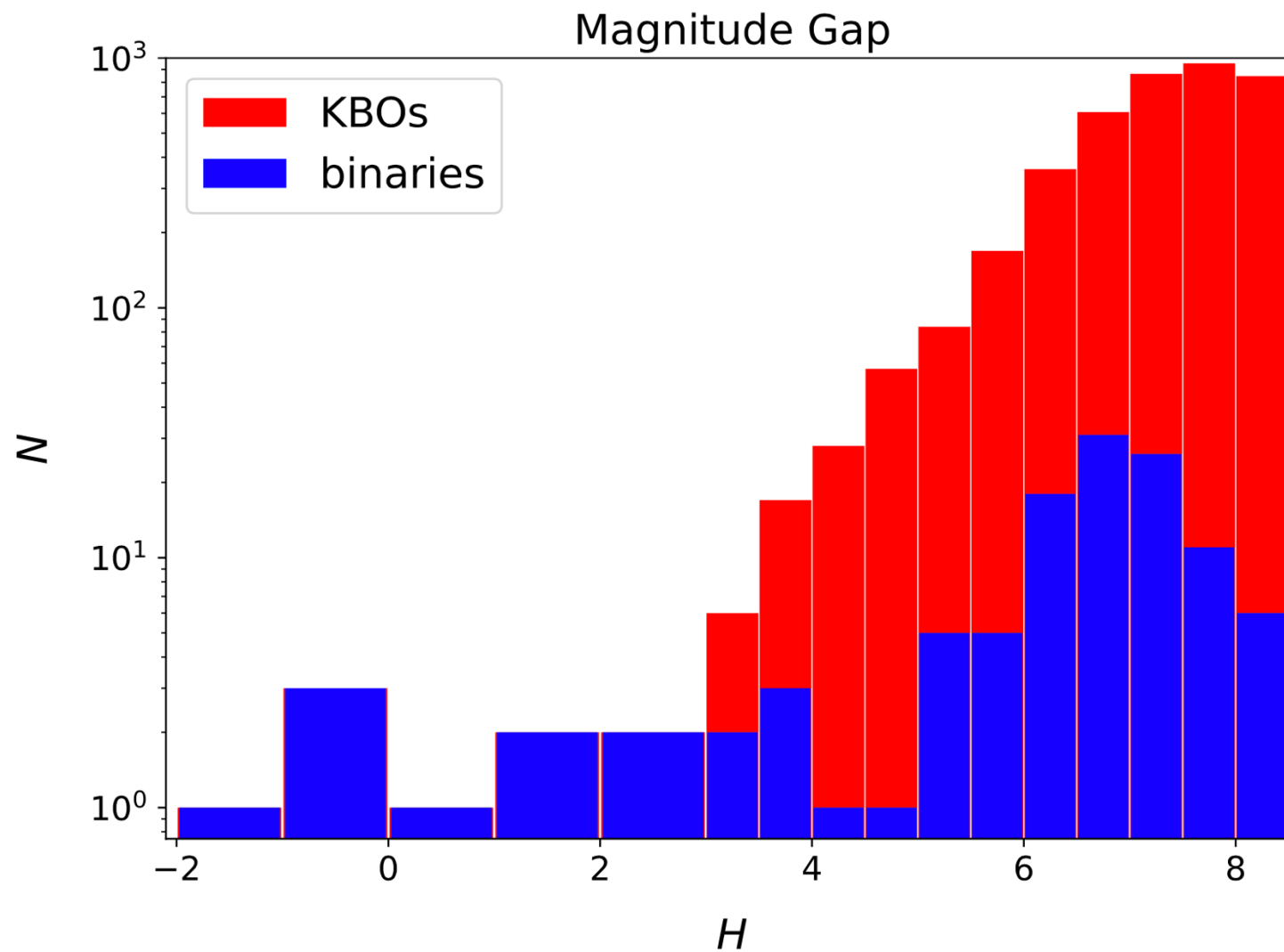
The DECam Ecliptic Exploration Project (DEEP). V. The Absolute Magnitude Distribution of the Cold Classical Kuiper Belt

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David E. Trilling⁵, Ryder Strauss⁵, William J. Oldroyd⁵, Chadwick A. Trujillo⁵, Colin Orion Chandler^{4,5,6},
Matthew J. Holman⁷, Hilke E. Schlichting⁸, and Andrew McNeill^{9,10}

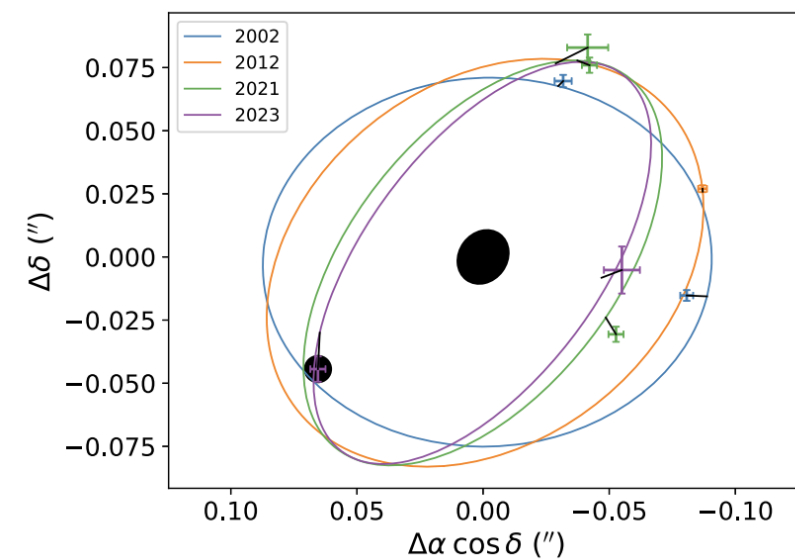
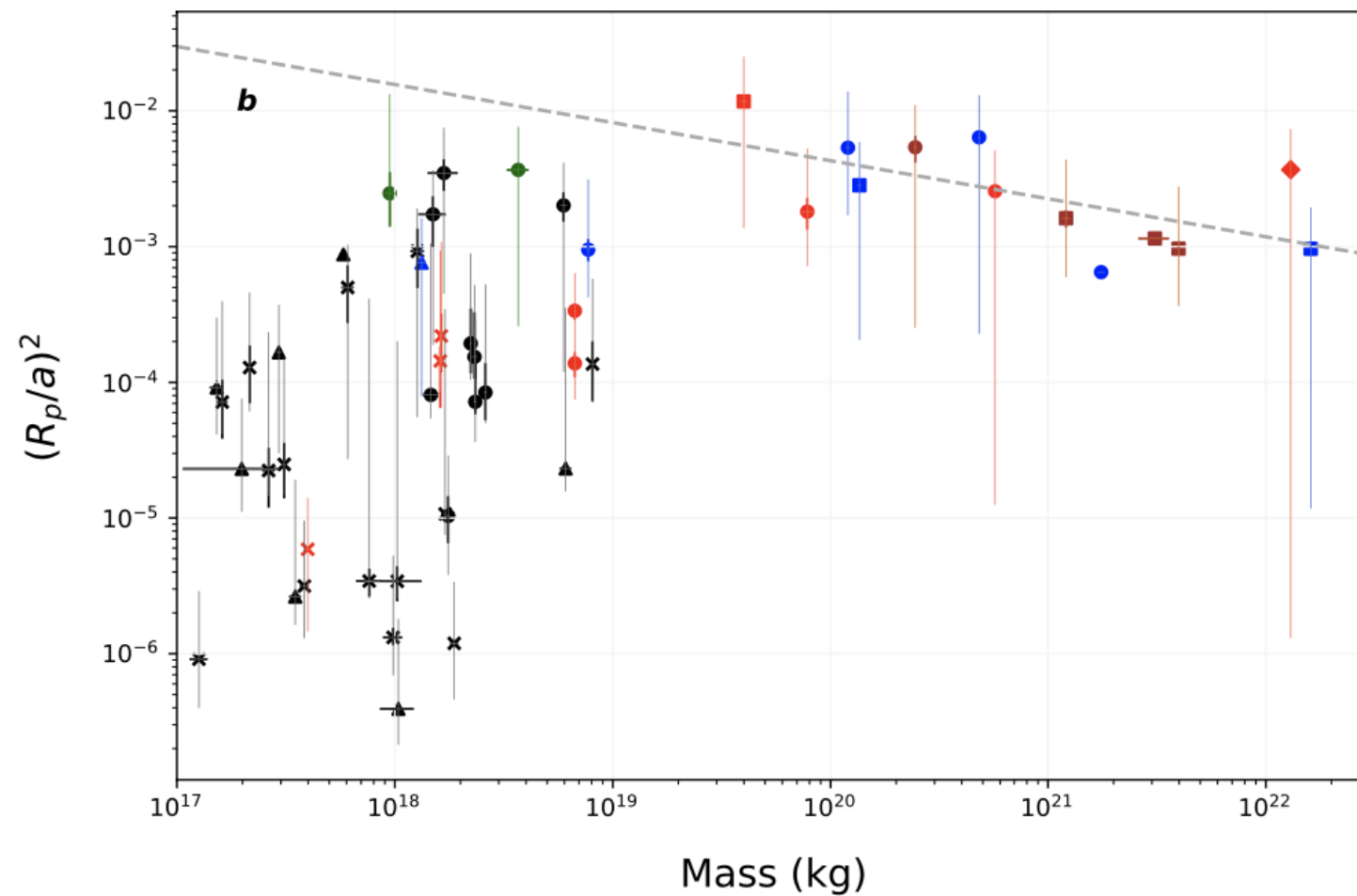


Kavelaars et al (2021); Napier et al. (2024)

High-mass side. Observational bias?

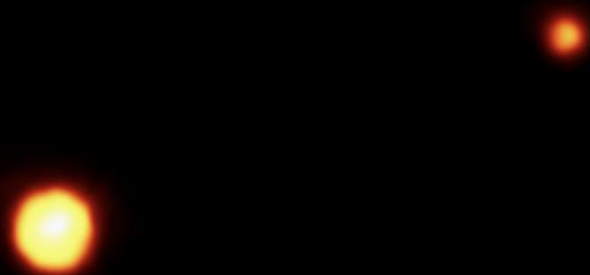


Non-Keplerian orbits



Rommel et al. (2025)

Pluto and Charon



Haumea,
Hi'aka, and
Namaka



Gonggong and Xianglu

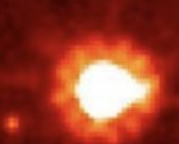


Orcus and Vanth



Many of the largest Kuiper Belt Objects have large mass fraction satellites!

Eris and Dysnomia



Makemake and MK2



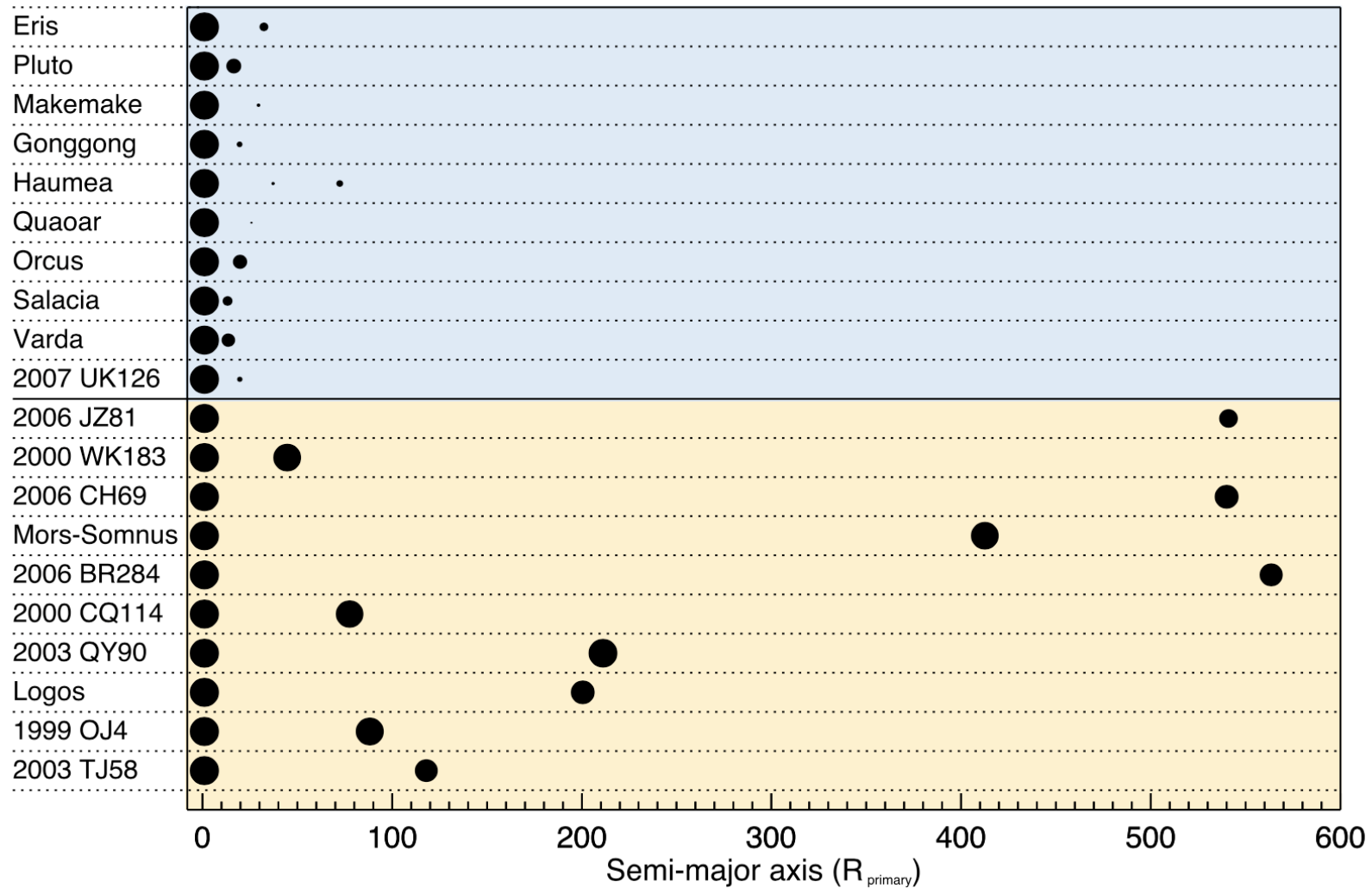
Varda and Ilmarë



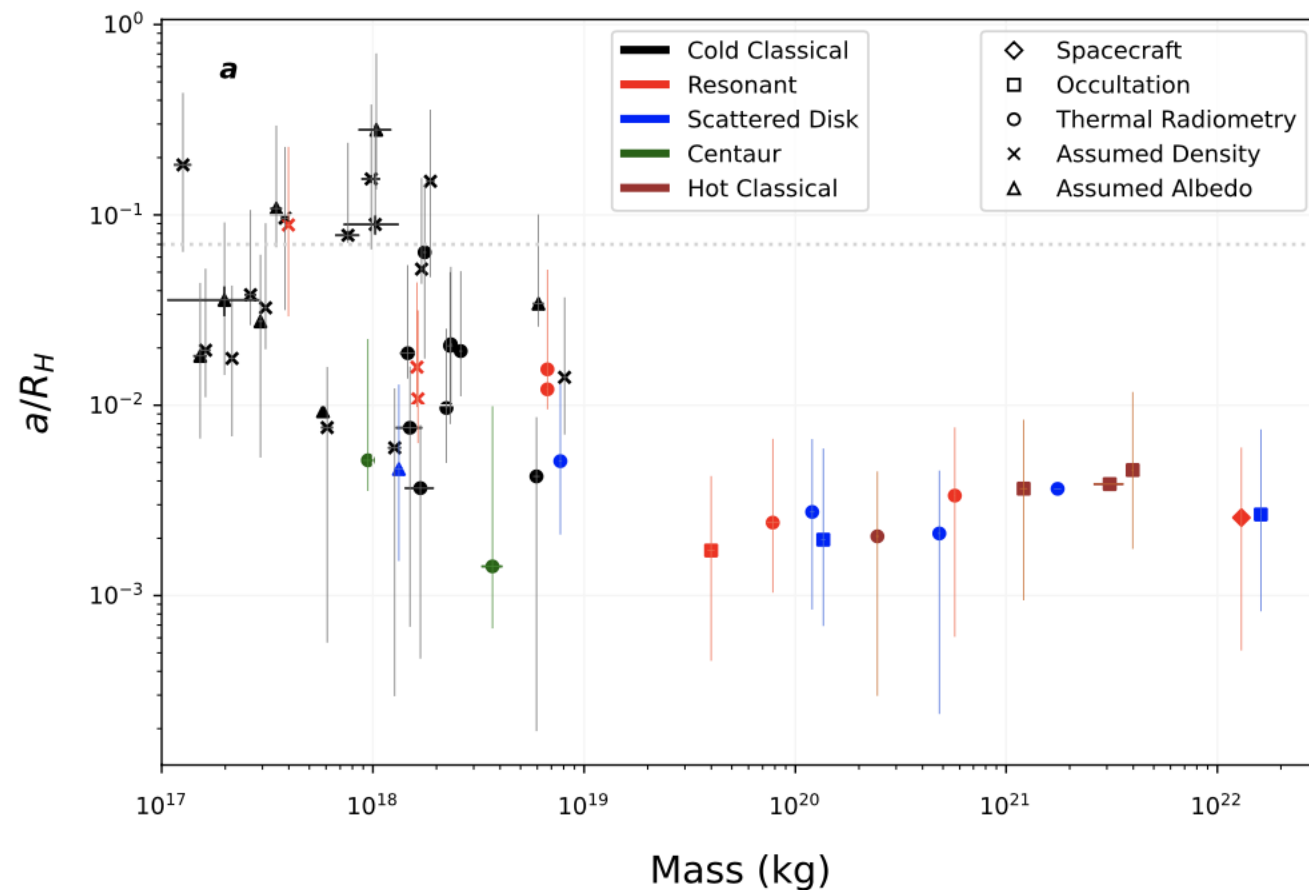
Quaoar and Weywot



Different system architecture

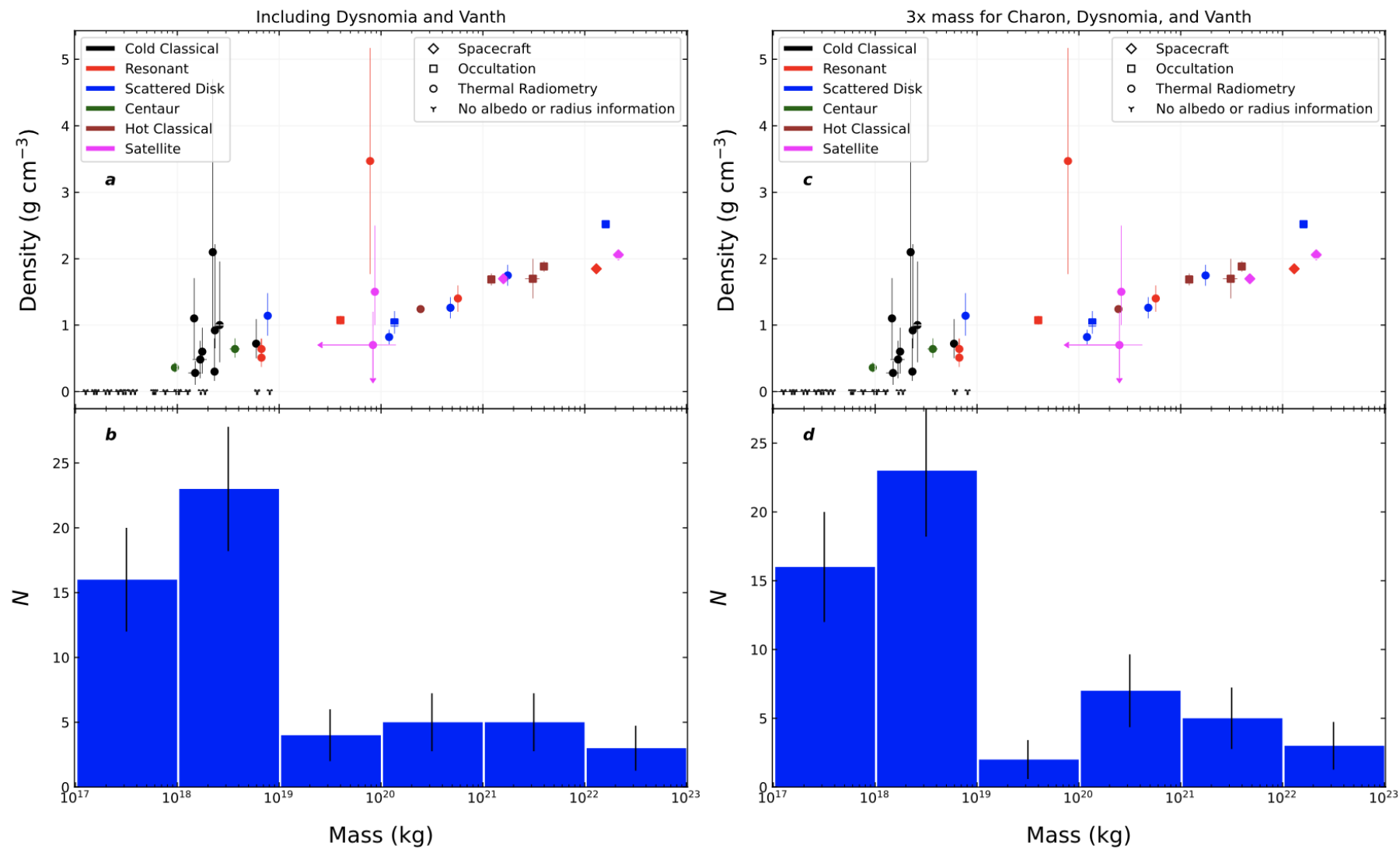


Did the high-mass objects lose their primordial satellites?



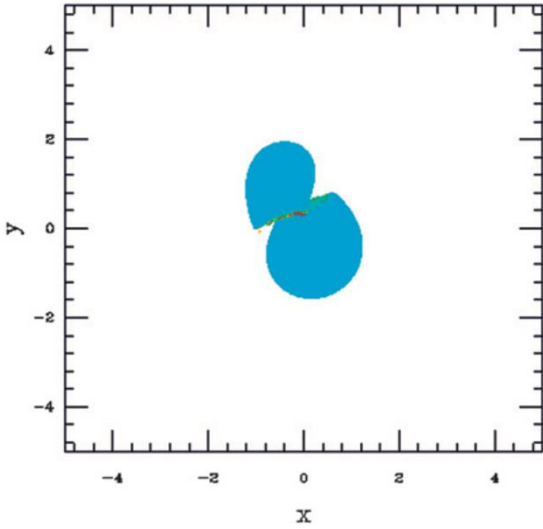
UWBs

Satellites

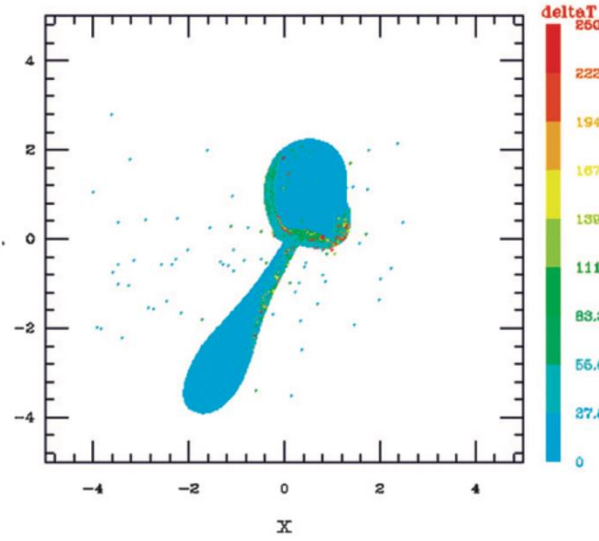


Pluto-Charon without strength high loss of mass scenario

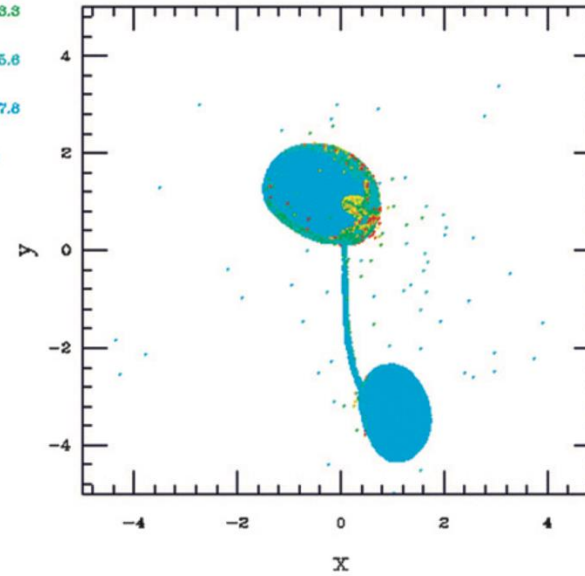
Time = 0.861108 hr



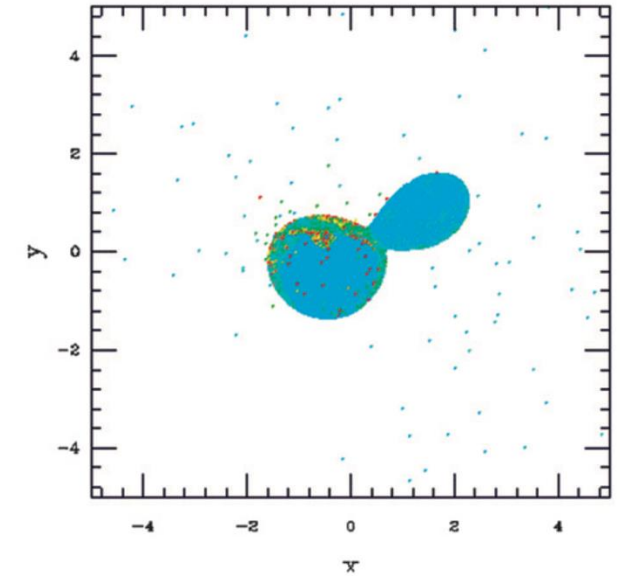
Time = 3.22824 hr



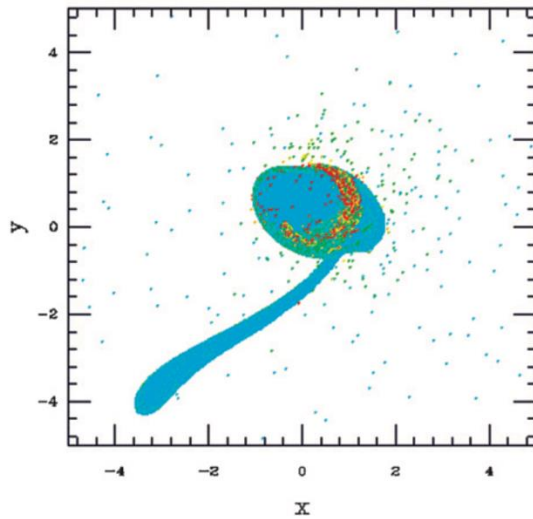
Time = 5.91801 hr



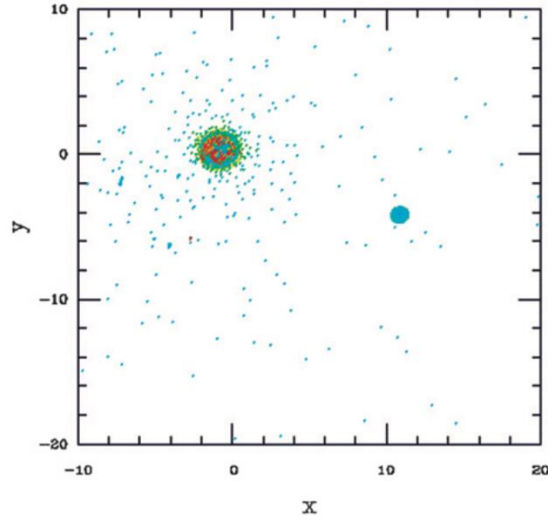
Time = 9.68417 hr



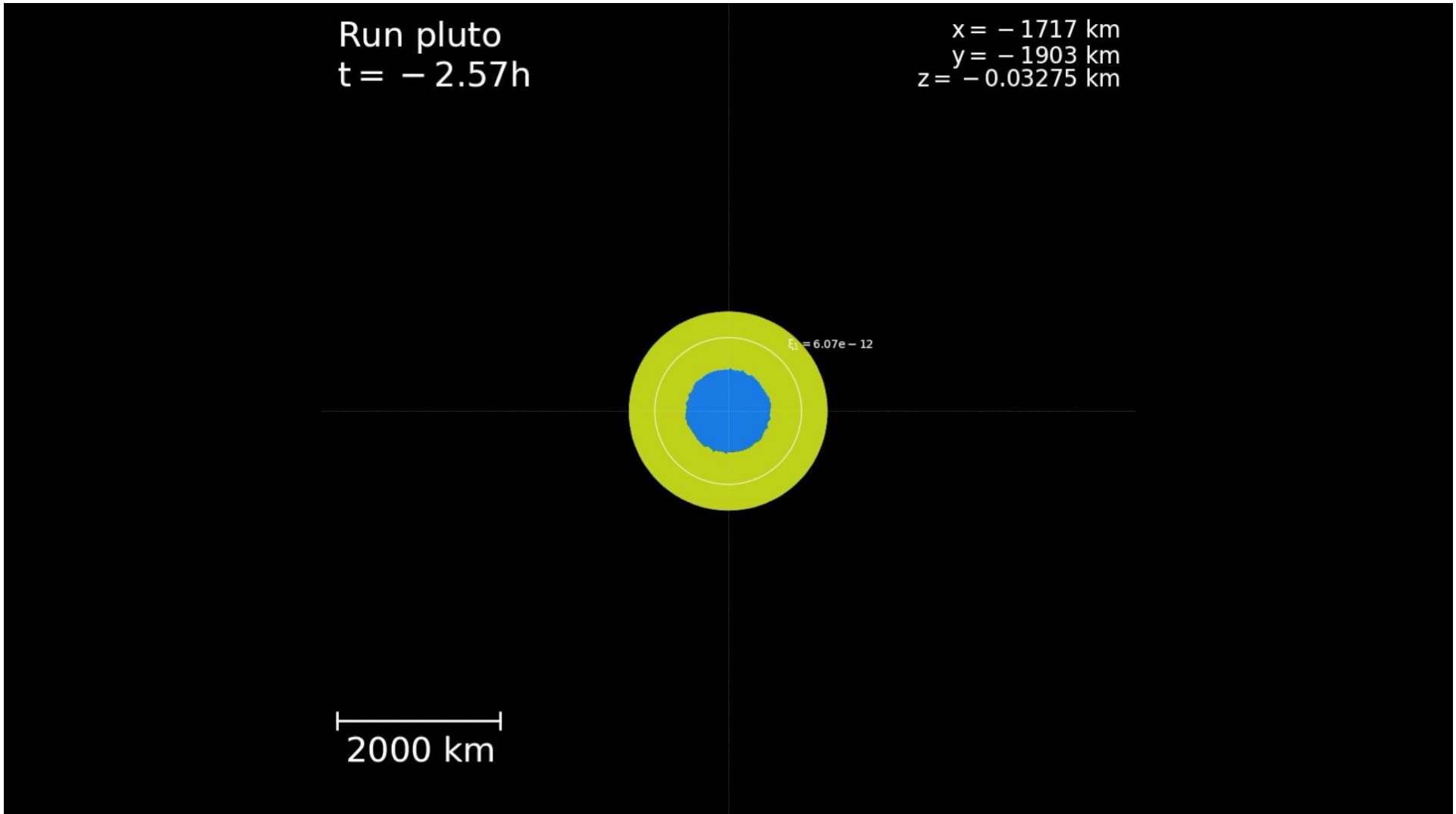
Time = 13.4503 hr



Time = 30.666 hr



Pluto-Charon with strength “kiss and capture” low loss of mass scenario



Conclusions

- KBO density problem:
 - Two different pebble populations, maintained by ice desorption off small grains
 - Streaming instability: icy-rich small objects; nearly uniform composition
 - Pebble accretion: silicate-rich larger objects; varied composition
 - Melting avoided by
 - ice-rich formation
 - ^{26}Al incorporated mostly in long ($>\text{Myr}$) phase of silicate accretion
 - KBOs best reproduced between 15-25 AU
- Missing Binaries
 - Cold classicals capped at 10^{-3} Pluto masses
 - Gap between 10^{-3} and 10^{-2} Pluto masses for non-cold classicals
 - Formation imprint?
 - Dynamical loss?
 - Observation bias?

