

Streaming Instability and Pebble Accretion: Evidence from the Kuiper Belt



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A Solution for the Density Dichotomy Problem of Kuiper Belt Objects with Multispecies Streaming Instability and Pebble Accretion

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Abstract

Kuiper Belt objects (KBOs) show an unexpected trend, whereby large bodies have increasingly higher densities, up to five times greater than their smaller counterparts. Current explanations for this trend assume formation at constant composition, with the increasing density resulting from gravitational compaction. However, this scenario poses a timing problem to avoid early melting by decay of ²⁶Al. We aim to explain the density trend in the context of streaming instability and pebble accretion. Small pebbles experience lofting into the atmosphere of the disk, being exposed to UV and partially losing their ice via desorption. Conversely, larger pebbles are shielded and remain icier. We use a shearing box model including gas and solids, the latter split into ices and silicate pebbles. Self-gravity is included, allowing dense clumps to collapse into planetesimals. We find that the streaming instability leads to the formation of mostly icy planetesimals, albeit with an unexpected trend that the lighter ones are more silicate-rich than the heavier ones. We feed the resulting planetesimals into a pebble accretion integrator with a continuous size distribution, finding that they undergo drastic changes in composition as they preferentially accrete silicate pebbles. The density and masses of large KBOs are best reproduced if they form between 15 and 22 au. Our solution avoids the timing problem because the first planetesimals are primarily icy and ²⁶Al is mostly incorporated in the slow phase of silicate pebble accretion. Our results lend further credibility to the streaming instability and pebble accretion as formation and growth mechanisms.

Unified Astronomy Thesaurus concepts: Dwarf planets (419); Kuiper Belt (893); Pluto (1267); Hydrodynamics (1963); Planet formation (1241)



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An Analytical Theory for the Growth from Planetesimals to Planets by Polydisperse Pebble Accretion

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Abstract

Pebble accretion is recognized as a significant accelerator of planet formation. Yet only formulae for single-sized (monodisperse) distribution have been derived in the literature. These can lead to significant underestimates for Bondi accretion, for which the best accreted pebble size may not be the one that dominates the mass distribution. We derive in this paper the polydisperse theory of pebble accretion. We consider a power-law distribution in pebble radius, and we find the resulting surface and volume number density distribution functions. We derive also the exact monodisperse analytical pebble accretion rate for which 3D accretion and 2D accretion are limits. In addition, we find analytical solutions to the polydisperse 2D Hill and 3D Bondi limits. We integrate the polydisperse pebble accretion numerically for the MRN distribution, finding a slight decrease (by an exact factor 3/7) in the Hill regime compared to the monodisperse case. In contrast, in the Bondi regime, we find accretion rates 1–2 orders of magnitude higher compared to monodisperse, also extending the onset of pebble accretion to 1–2 orders of magnitude lower in mass. We find megayear timescales, within the disk lifetime, for Bondi accretion on top of planetary seeds of masses 10^{-6} to $10^{-4} M_{\oplus}$, over a significant range of the parameter space. This mass range overlaps with the high-mass end of the planetesimal initial mass function, and thus pebble accretion is possible directly following formation by streaming instability. This alleviates the need for mutual planetesimal collisions as a major contribution to planetary growth.

Unified Astronomy Thesaurus concepts: Planet formation (1241); Planetary system formation (1257)



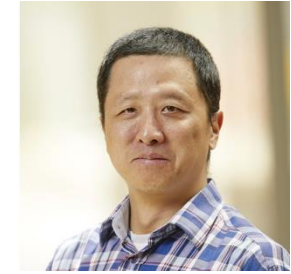
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The PFITS+ Collaboration (Planet Formation in the Southwest)



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The size-density relationship of Kuiper Belt objects

THE DENSITY OF MID-SIZED KUIPER BELT OBJECT 2002 UX25 AND THE FORMATION OF THE DWARF PLANETS

M. E. BROWN

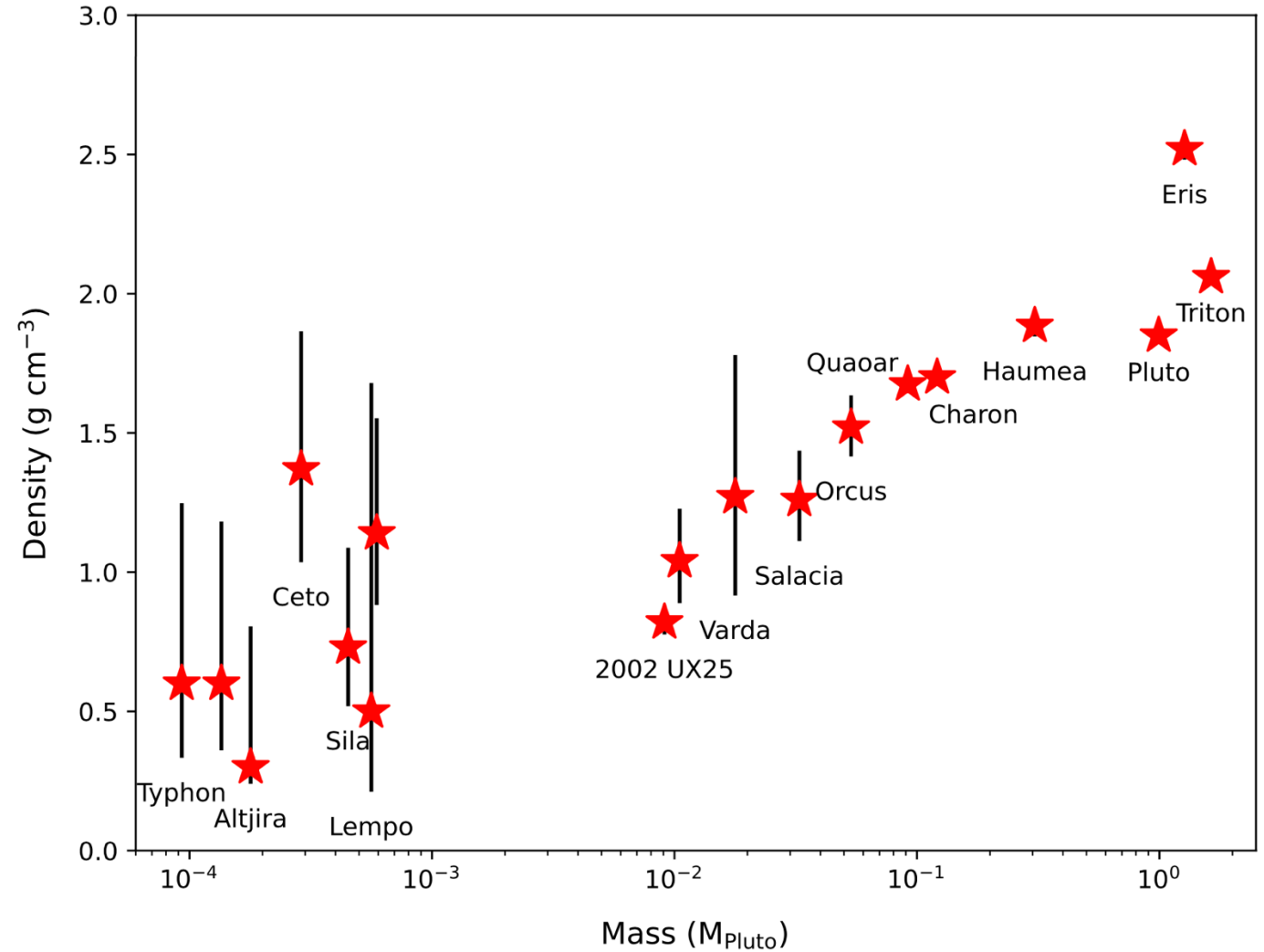
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ABSTRACT

The inferred low rock fraction of the 2002 UX25 system makes the formation of rock-rich larger objects difficult to explain in any standard coagulation scenario. For example, to create an object with the volume of Eris would require assembling ~ 40 objects of the size of 2002 UX25. Yet the assembled object, even with the additional compression, would still have a density close to 1 g cm^{-3} rather than the 2.5 g cm^{-3} density of Eris (Sicardy et al. 2011).

- Extremely low porosity;
- Biased sample;
- Compaction through giant impacts

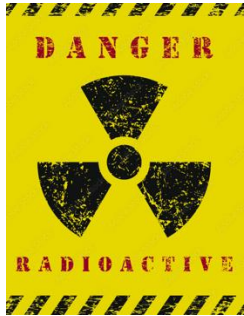
None of these alternatives appears likely. We are left in the uncomfortable state of having no satisfying mechanism to explain the formation of the icy dwarf planets. While objects up to the size of 2002 UX25 can easily be formed through standard coagulation scenarios, the rock-rich larger bodies may require a formation mechanism separate from the rest of the Kuiper belt.



Previous best bet: Porosity removal by gravitational compaction

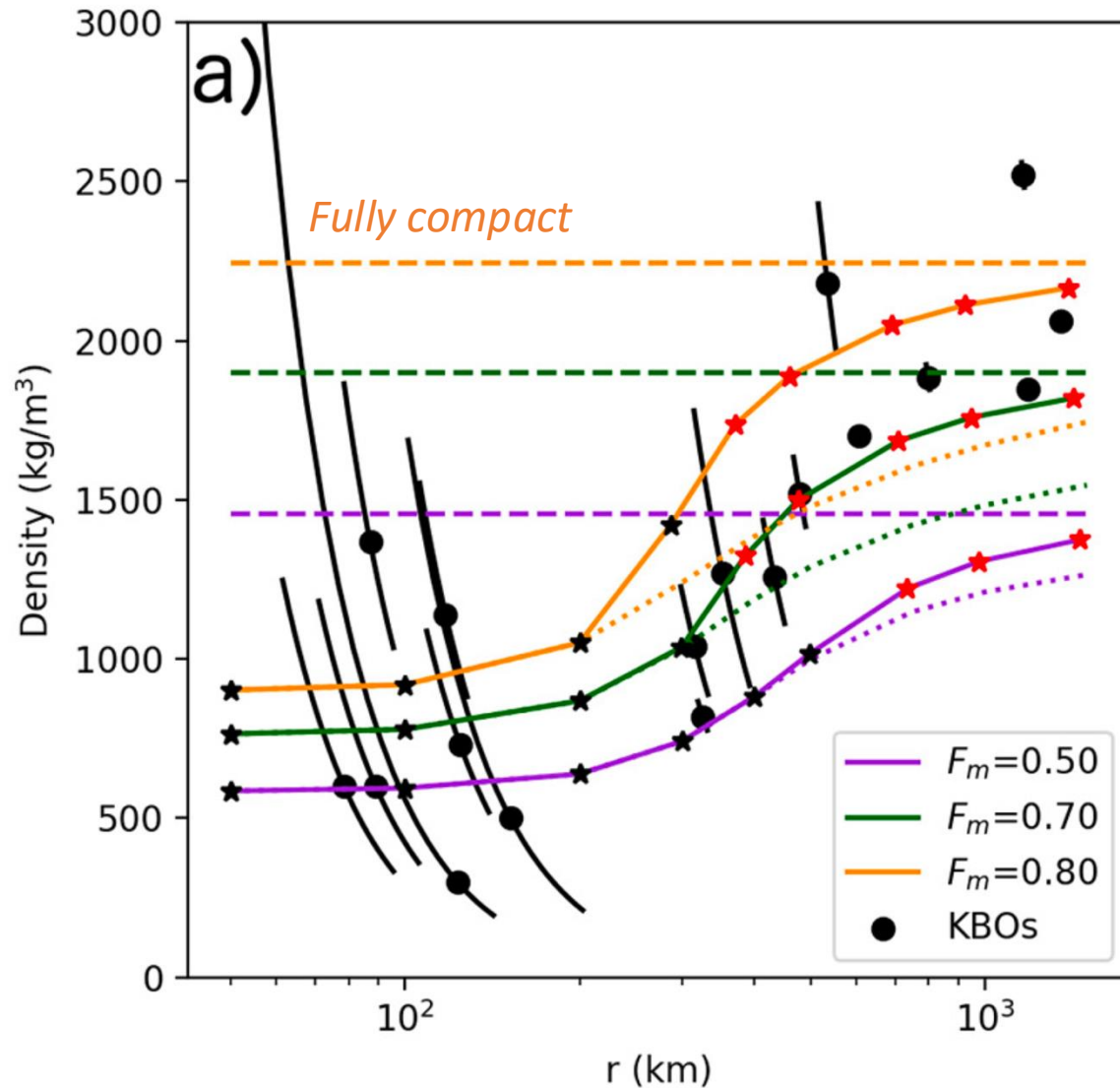
Problem

- Timing! ^{26}Al would melt if formed within 4 Myr



Assumptions

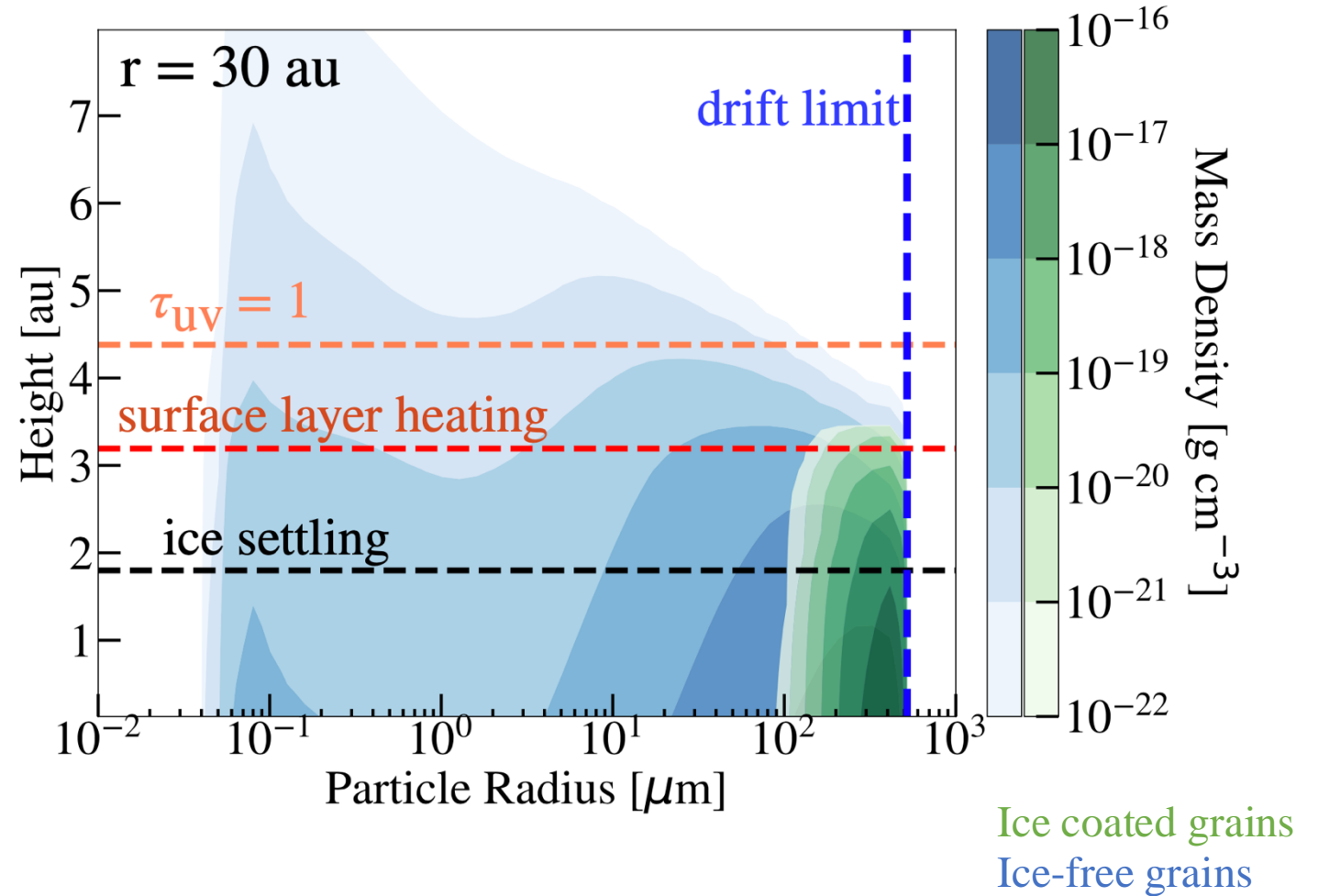
- ~~Constant composition at birth and growth~~
- ~~Growth by planetesimal accretion~~



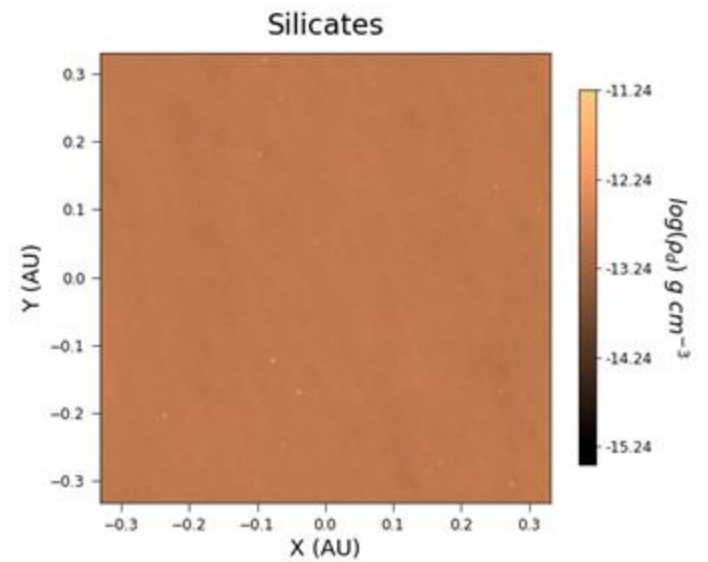
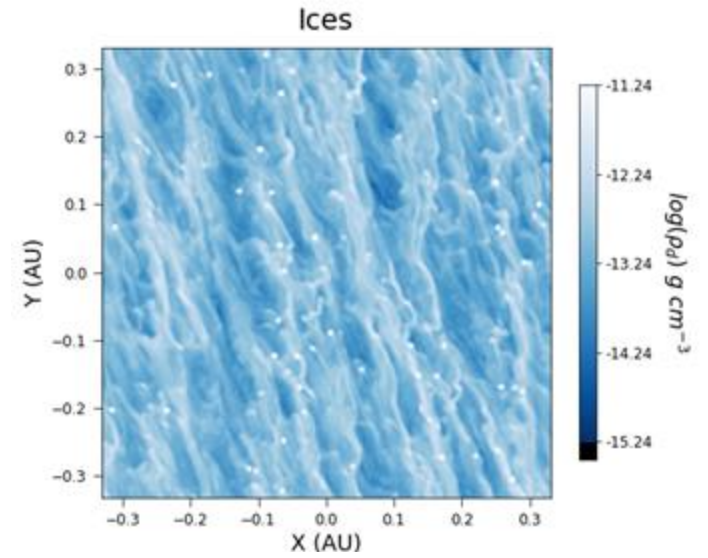
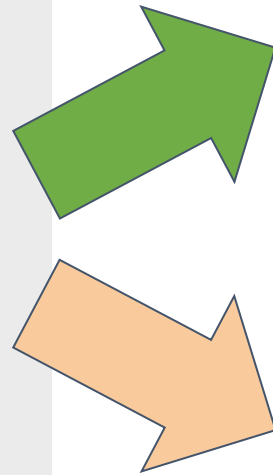
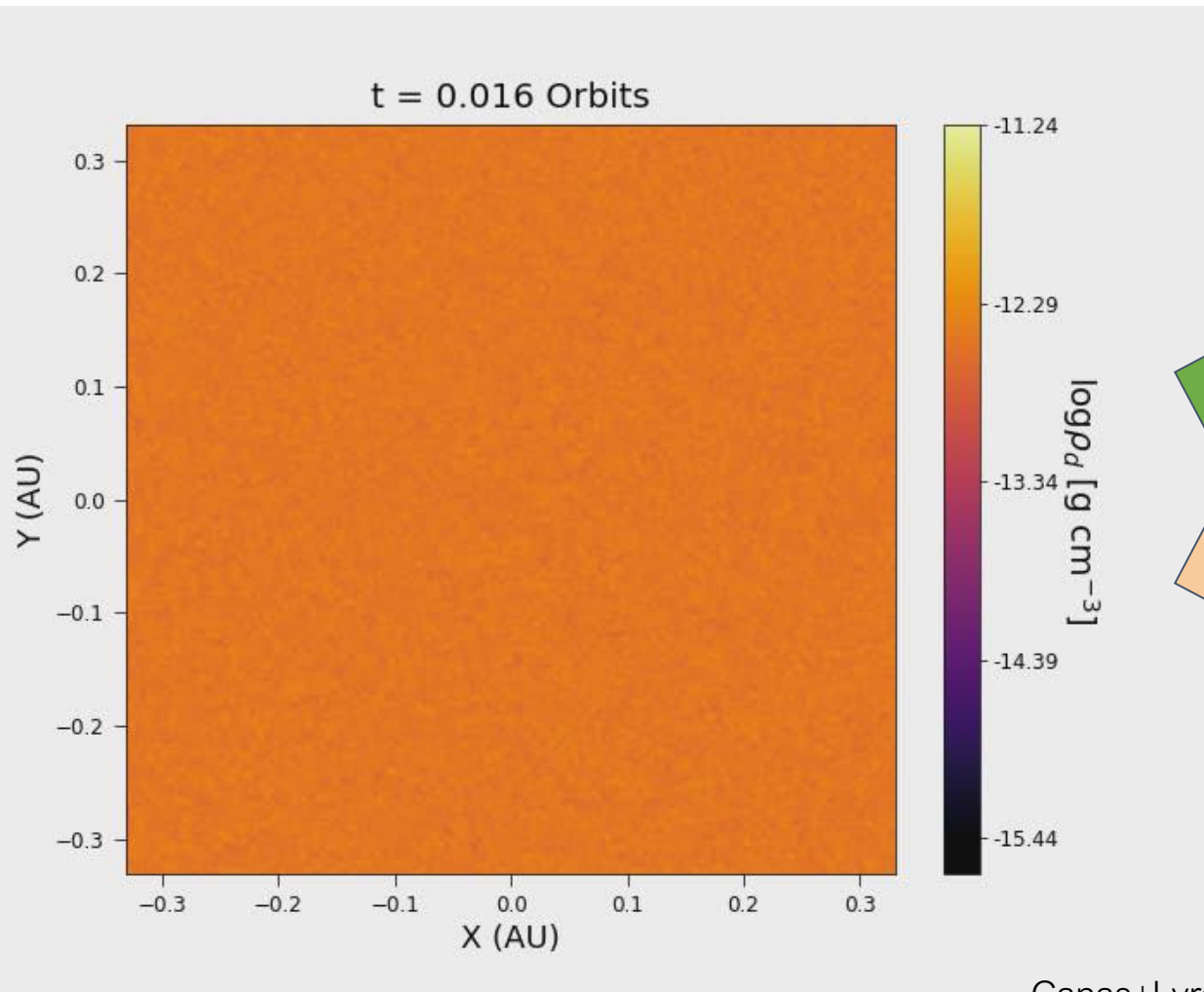
Abandoning Constant Composition

Heating and UV irradiation remove ice on Myr timescales (Harrison & Schoen 1967)

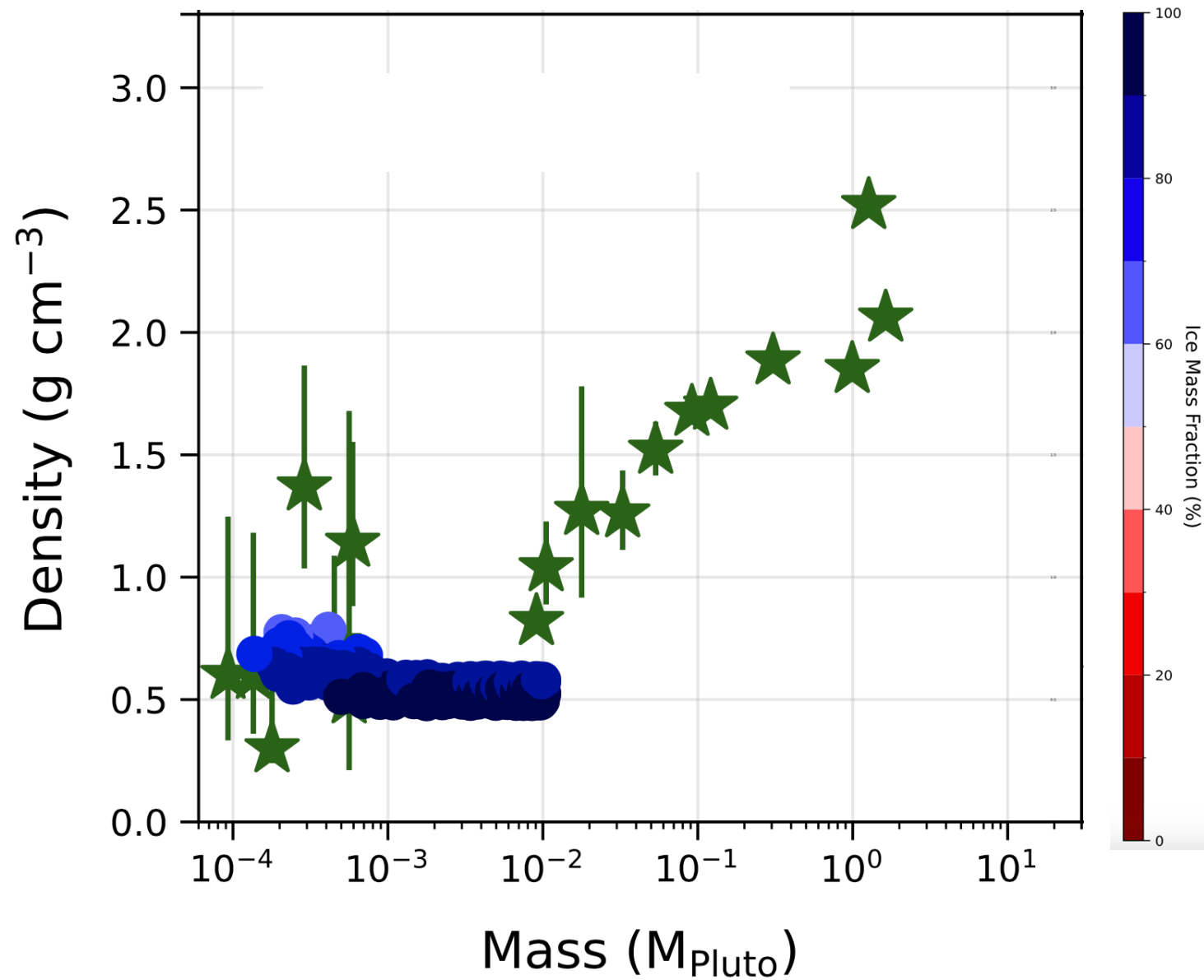
- Small grains lofted in the atmosphere lose ice
- Big grains are shielded and remain icy.



Split into icy and silicate pebbles



The first planetesimals are icy



The first planetesimals won't melt

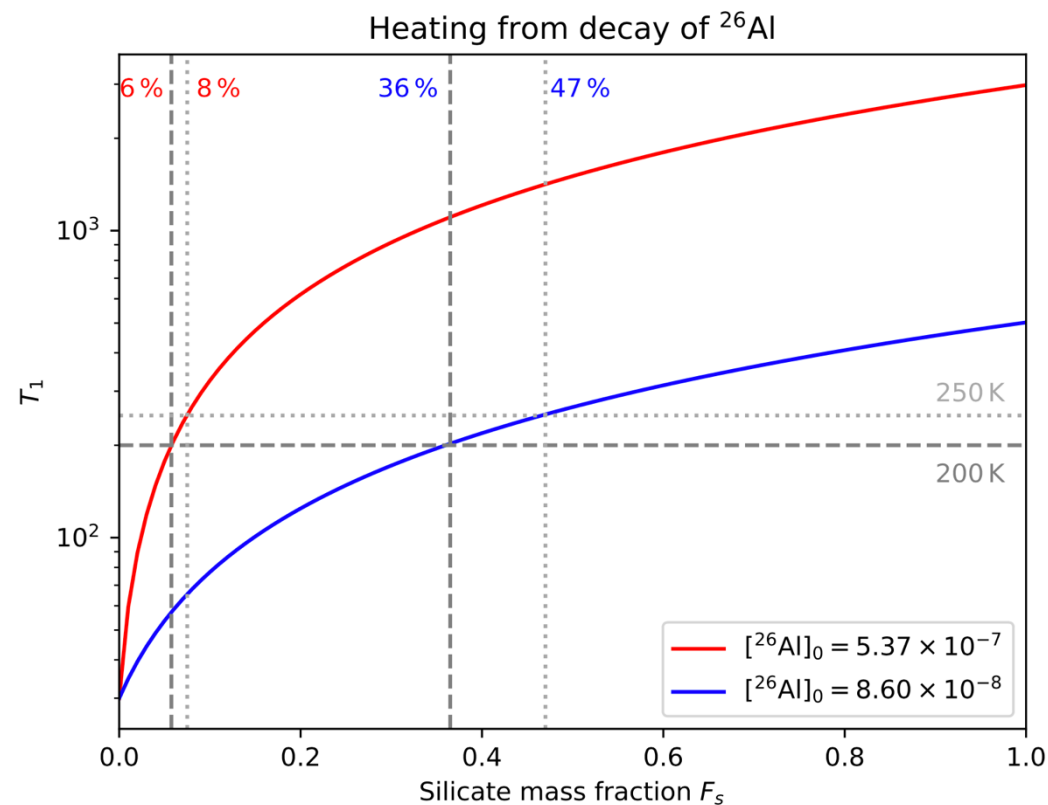
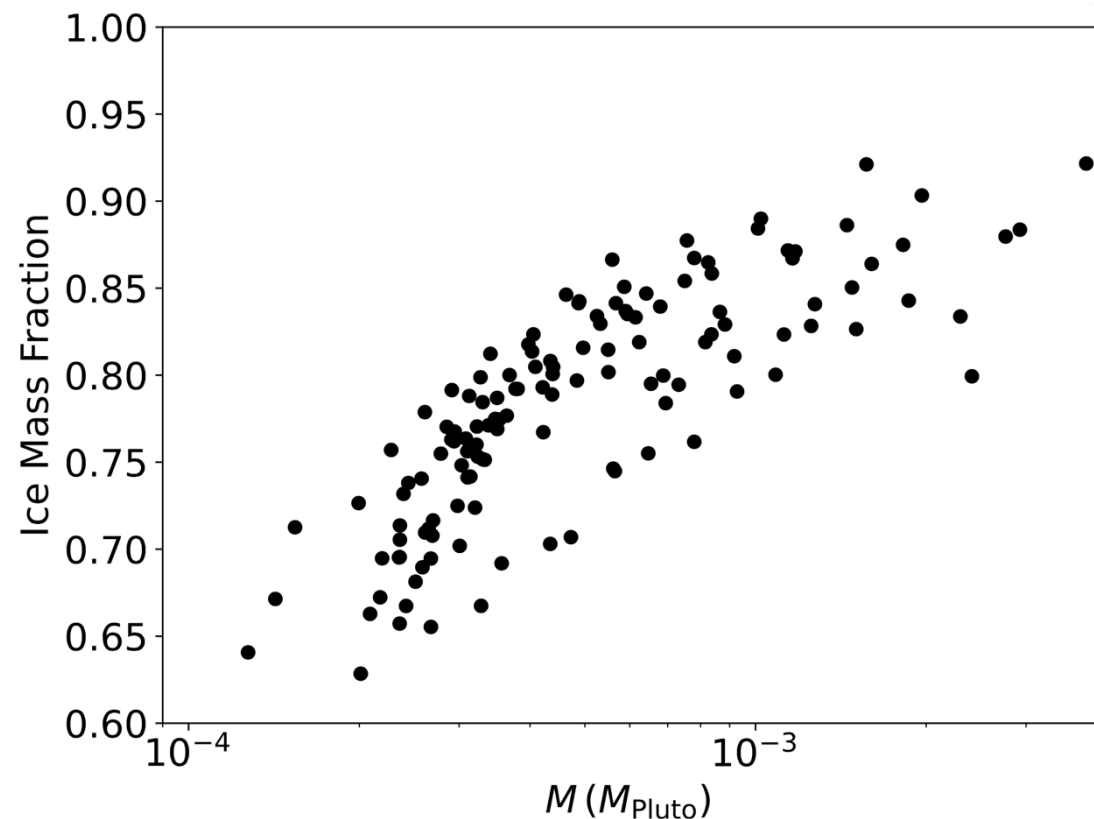
$$\mathcal{H} = \rho F_s [^{26}\text{Al}]_0 \mathcal{H}_0 e^{-\lambda t}$$

$$Q(t) = V \int_0^t \mathcal{H}(t') dt'$$

$$= M_p F_s [^{26}\text{Al}]_0 \mathcal{H}_0 \lambda^{-1} (1 - e^{-\lambda t})$$

$$Q = M_p c_p \Delta T$$

$$\Delta T = F_s [^{26}\text{Al}]_0 \mathcal{H}_0 \lambda^{-1} c_p^{-1}$$



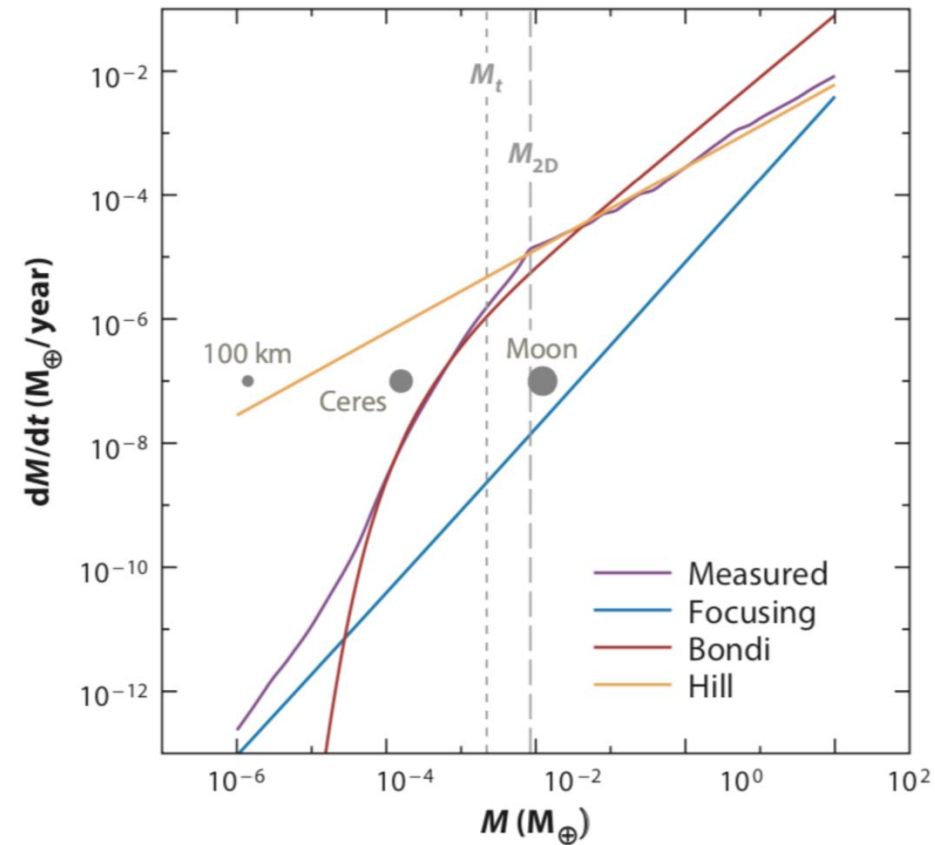
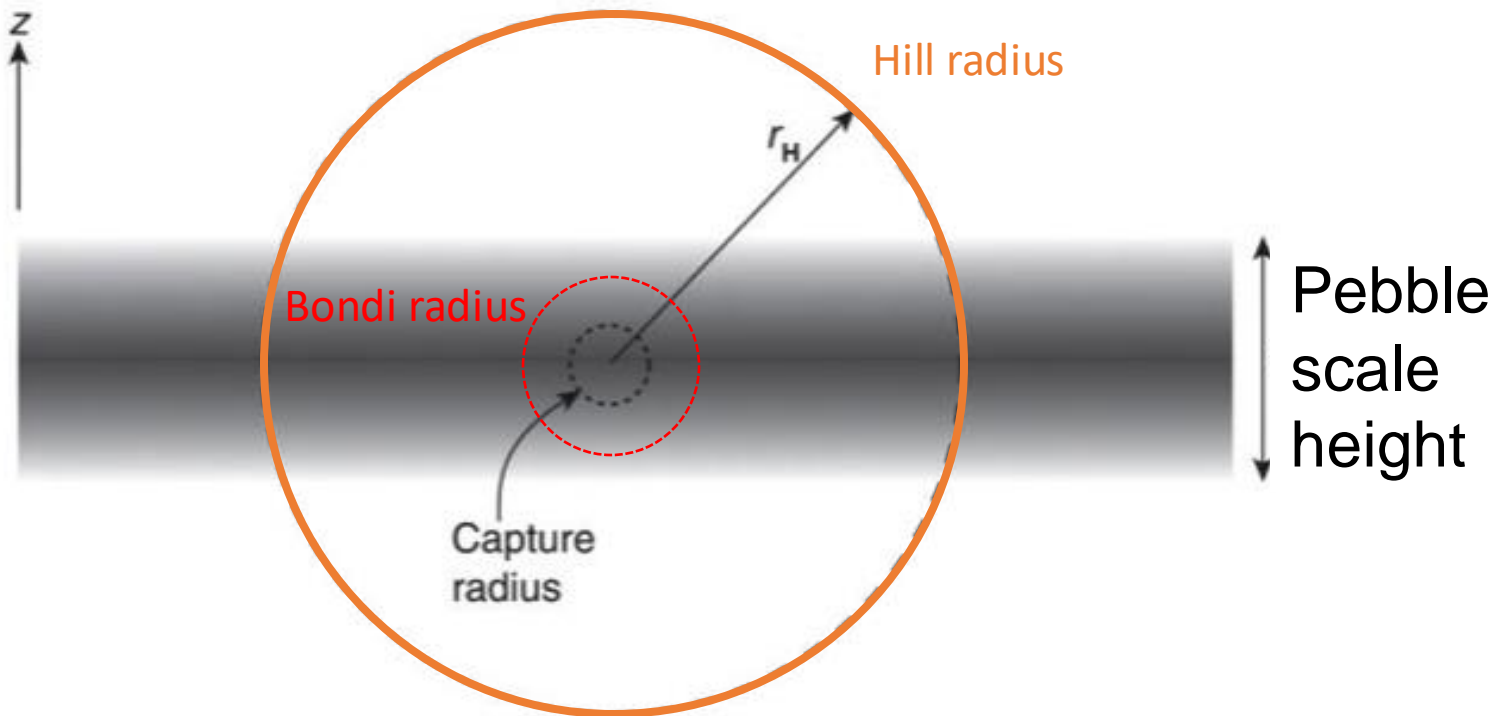
Pebble Accretion: Geometric, Bondi, and Hill regime

Bondi accretion - Bound against **headwind**

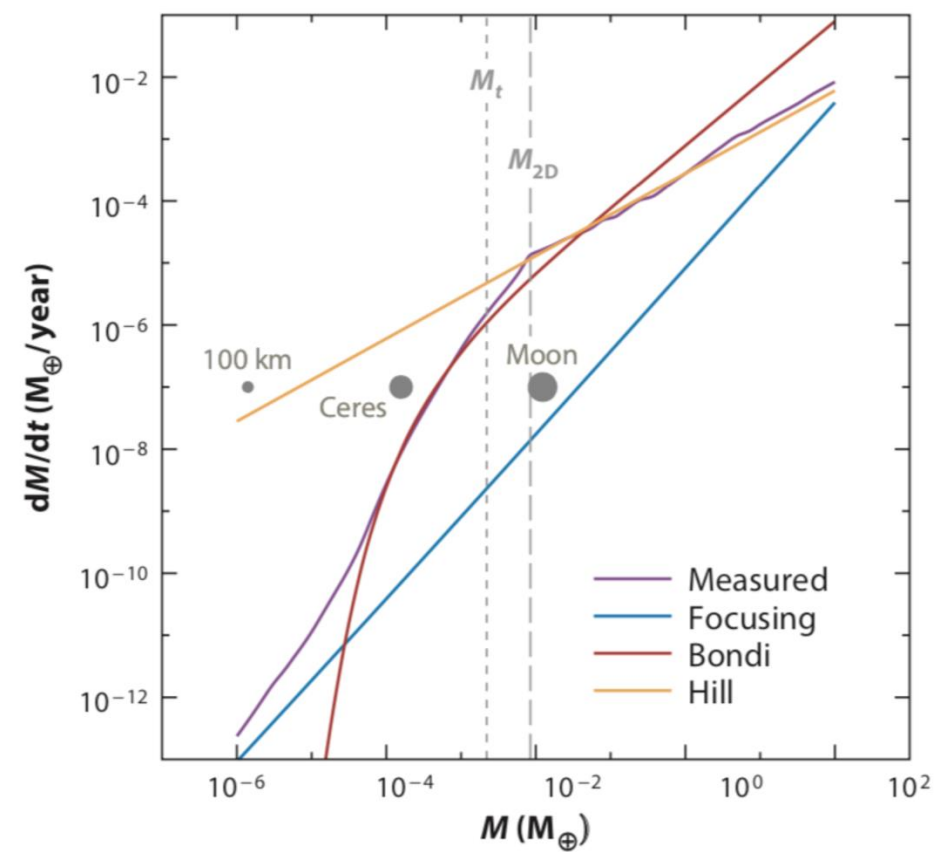
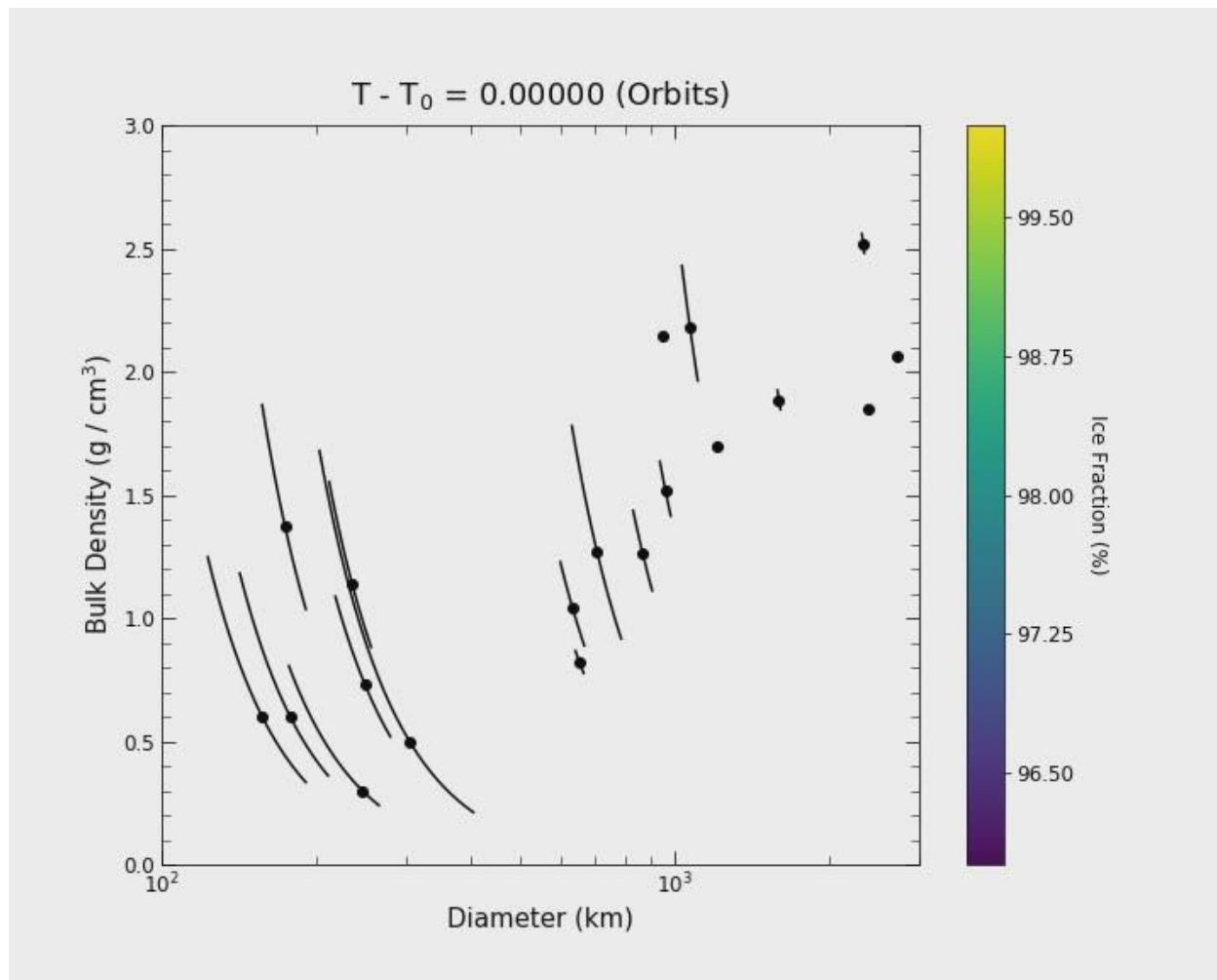
Hill accretion - Bound against **stellar tide**

$$\xi \equiv \left(\frac{R_{\text{acc}}}{2H_d} \right)^2 \quad \begin{aligned} \dot{M}_{3D} &= \lim_{\xi \rightarrow 0} \dot{M} = \pi R_{\text{acc}}^2 \rho_{d0} \delta v, \\ \dot{M}_{2D} &= \lim_{\xi \rightarrow \infty} \dot{M} = 2 R_{\text{acc}} \Sigma_d \delta v, \end{aligned}$$

Mass Accretion rates

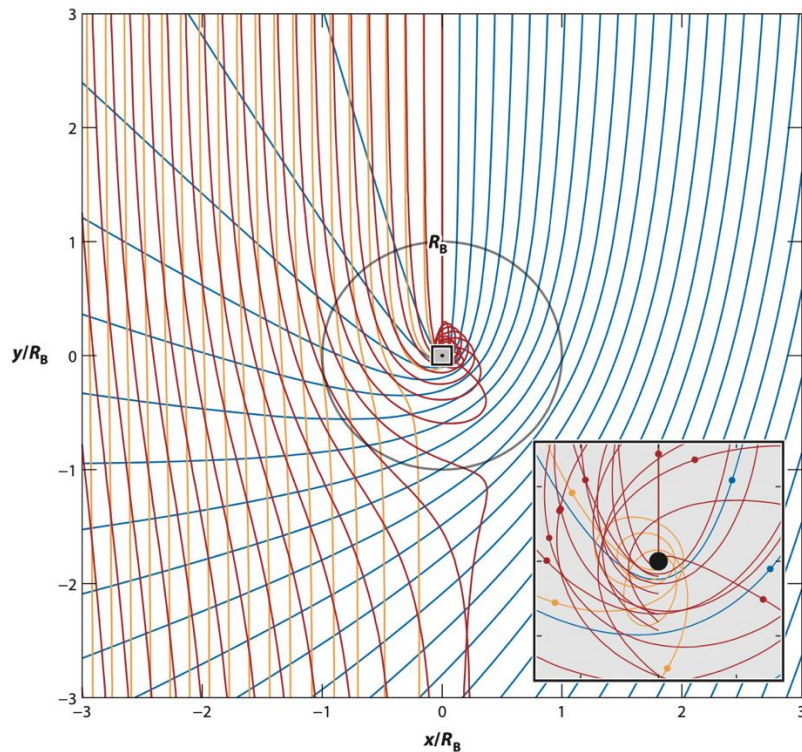


Integrate pebble accretion



Pebble Accretion: Pebbles of different size accrete differently

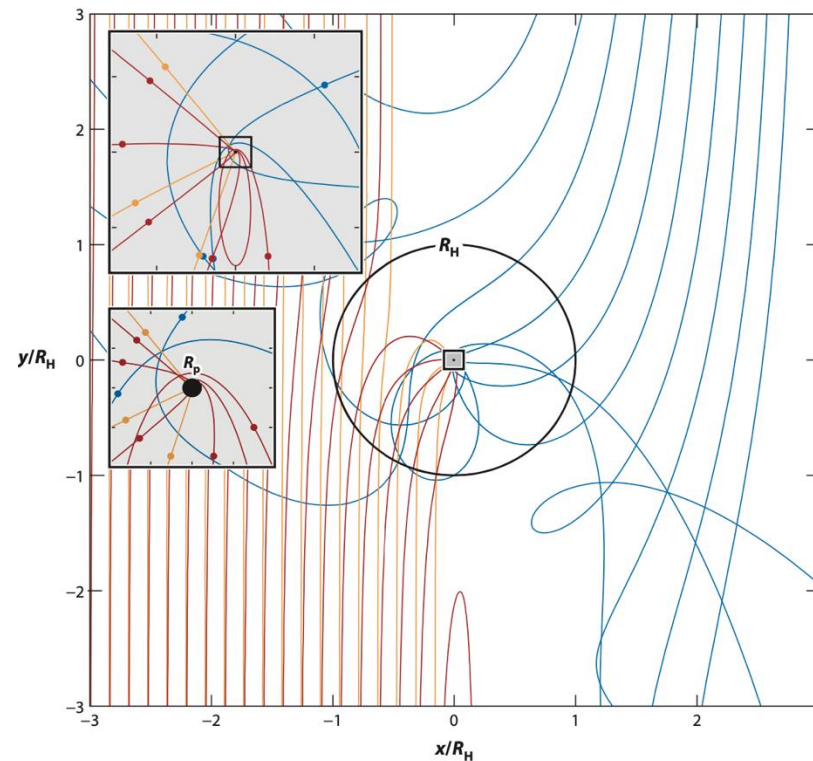
Bondi Regime



Best accreted pebble

Drag time \sim Bondi Time

Hill Regime



Best accreted pebble

Drag time \sim Orbital Time

Polydisperse (Multi-Species) Pebble Accretion

$$\rho_d(a, z) = \int_0^a m(a') F(a', z) da'.$$

$$F(a, z) \equiv f(a) e^{-z^2/2H_d^2},$$

$$f(a) = \frac{3(1-p)Z\Sigma_g}{2^{5/2}\pi^{3/2}H_g\rho_\bullet^{(0)}a_{\max}^{4-k}} \sqrt{1 + a\frac{\pi}{2}\frac{\rho_\bullet(a)}{\Sigma_g\alpha}} a^{-k}.$$

$$S \equiv \frac{1}{\pi R_{\text{acc}}^2} \int_{-R_{\text{acc}}}^{R_{\text{acc}}} 2\sqrt{R_{\text{acc}}^2 - z^2} \exp\left(-\frac{z^2}{2H_d^2}\right) dz,$$

$$W(a) = \frac{3(1-p)Z\Sigma_g}{4\pi\rho_\bullet^{(0)}a_{\max}^{4-k}} a^{-k},$$

$$\hat{R}_{\text{acc}}^{(\text{Bondi})} = \left(\frac{4\tau_f}{t_B}\right)^{1/2} R_B,$$

$$\hat{R}_{\text{acc}}^{(\text{Hill})} = \left(\frac{\text{St}}{0.1}\right)^{1/3} R_H,$$

$$\delta v \equiv \Delta v + \Omega R_{\text{acc}},$$

$$R_{\text{acc}} \equiv \hat{R}_{\text{acc}} \exp[-\chi(\tau_f/t_p)^\gamma],$$

$$\frac{\partial \Sigma_d(a)}{\partial a} \propto a^{-p};$$

$$\rho_\bullet \propto a^{-q}; \quad t_p \equiv \frac{GM_p}{(\Delta v + \Omega R_H)^3}$$

$$\dot{M}(a) = \int_0^a \frac{\partial \dot{M}(a')}{\partial a'} da',$$

$$\frac{\partial \dot{M}(a)}{\partial a} = \pi R_{\text{acc}}^2(a) \delta v(a) S(a) m(a) f(a).$$

$$\dot{M}_{\text{2D, Hill}} = 2 \times 10^{2/3} \Omega R_H^2 \int_0^{a_{\max}} \text{St}(a)^{2/3} m(a) W(a) da.$$

$$\begin{aligned} \dot{M}_{\text{3D, Bondi}} &= \frac{4\pi R_B \Delta v^2}{\Omega} \\ &\times \int_0^{a_{\max}} \text{St} e^{-2\psi} m(a) f(a) \\ &\times \left[1 + 2 \left(\text{St} \frac{\Omega R_B}{\Delta v} \right)^{1/2} e^{-\psi} \right] da, \quad \psi \equiv \chi[\text{St}/(\Omega t_p)]^\gamma. \end{aligned}$$

Analytical theory of polydisperse (multi-species) pebble accretion

Monodisperse (single species)

$$\xi \equiv \left(\frac{R_{\text{acc}}}{2H_d} \right)^2$$

$$\dot{M}_{3D} = \lim_{\xi \rightarrow 0} \dot{M} = \pi R_{\text{acc}}^2 \rho_{d0} \delta v,$$

$$\dot{M}_{2D} = \lim_{\xi \rightarrow \infty} \dot{M} = 2R_{\text{acc}} \Sigma_d \delta v,$$

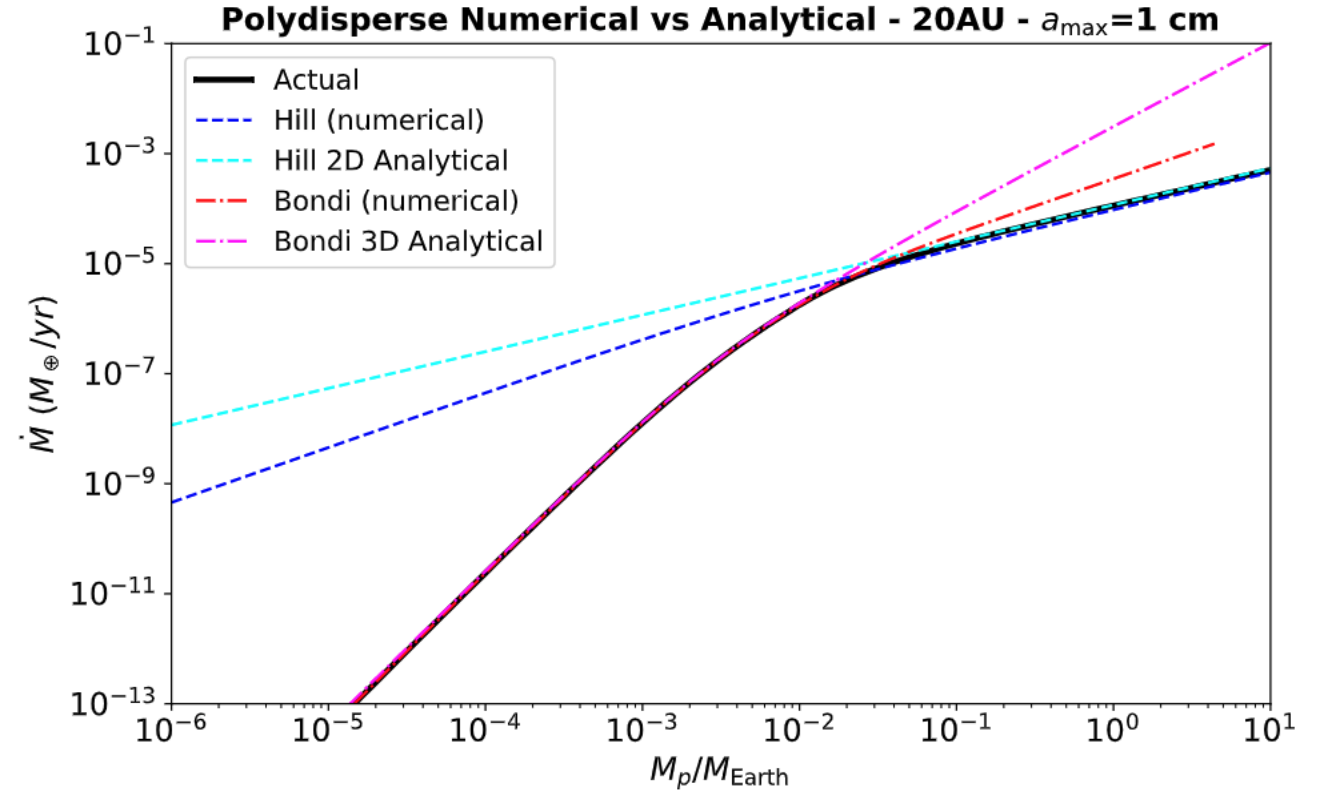
Lambrechts & Johansen (2012)

Polydisperse (multiple species)

$$\dot{M}_{2D,\text{Hill}} = \frac{6(1-p)}{14-5q-3k} \left(\frac{\text{St}_{\text{max}}}{0.1} \right)^{2/3} \Omega R_H^2 Z \Sigma_g.$$

$$\dot{M}_{3D,\text{Bondi}} \approx C_1 \frac{\gamma_l \left(\frac{b_1+1}{s}, j_1 a_{\text{max}}^s \right)}{s j_1^{(b_1+1)/s}} + C_2 \frac{\gamma_l \left(\frac{b_2+1}{s}, j_2 a_{\text{max}}^s \right)}{s j_2^{(b_2+1)/s}} + C_3 \frac{\gamma_l \left(\frac{b_3+1}{s}, j_3 a_{\text{max}}^s \right)}{s j_3^{(b_3+1)/s}} + C_4 \frac{\gamma_l \left(\frac{b_4+1}{s}, j_4 a_{\text{max}}^s \right)}{s j_4^{(b_4+1)/s}},$$

Lyra et al. (2023)



Lyra et al. 2023

Analytical Solution for General Monodisperse (single species) Pebble Accretion

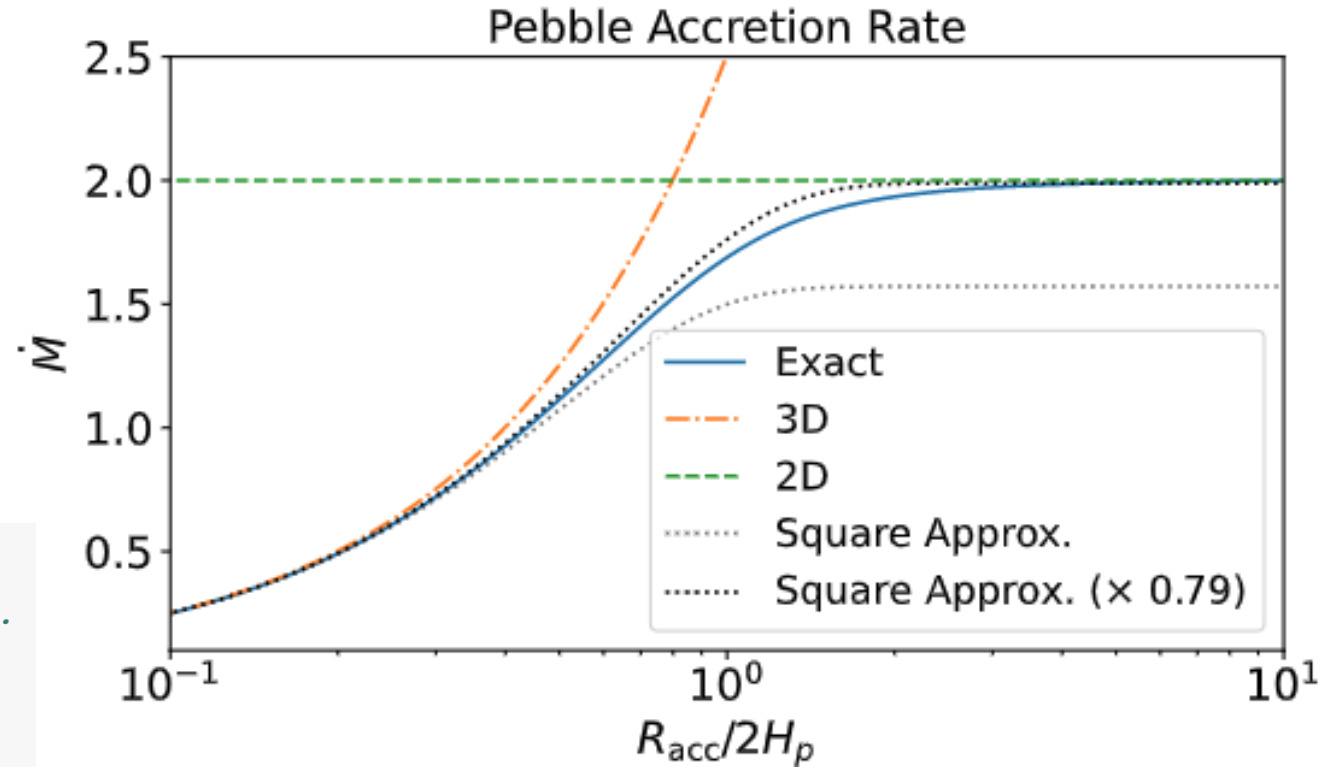
$$\dot{M} = \pi R_{\text{acc}}^2 \rho_{d0} S \delta v.$$

$$S \equiv \frac{1}{\pi R_{\text{acc}}^2} \int_{-R_{\text{acc}}}^{R_{\text{acc}}} 2 \sqrt{R_{\text{acc}}^2 - z^2} \exp\left(-\frac{z^2}{2H_d^2}\right) dz,$$

$$S = e^{-\xi} [I_0(\xi) + I_1(\xi)], \quad \xi \equiv \left(\frac{R_{\text{acc}}}{2H_d}\right)^2$$

```
y = (x/2)**2
# Modified Bessel function of the first kind of real order.
I0 = sp.special.iv(0, y)
I1 = sp.special.iv(1, y)

Sint = np.exp(-y) * (I0 + I1)
rho_int = rhop * Sint
Mdot = pi*r**2 * rho_int * deltav
```



Analytical Solutions for 2D and 3D Polydisperse (multi-species) Pebble Accretion

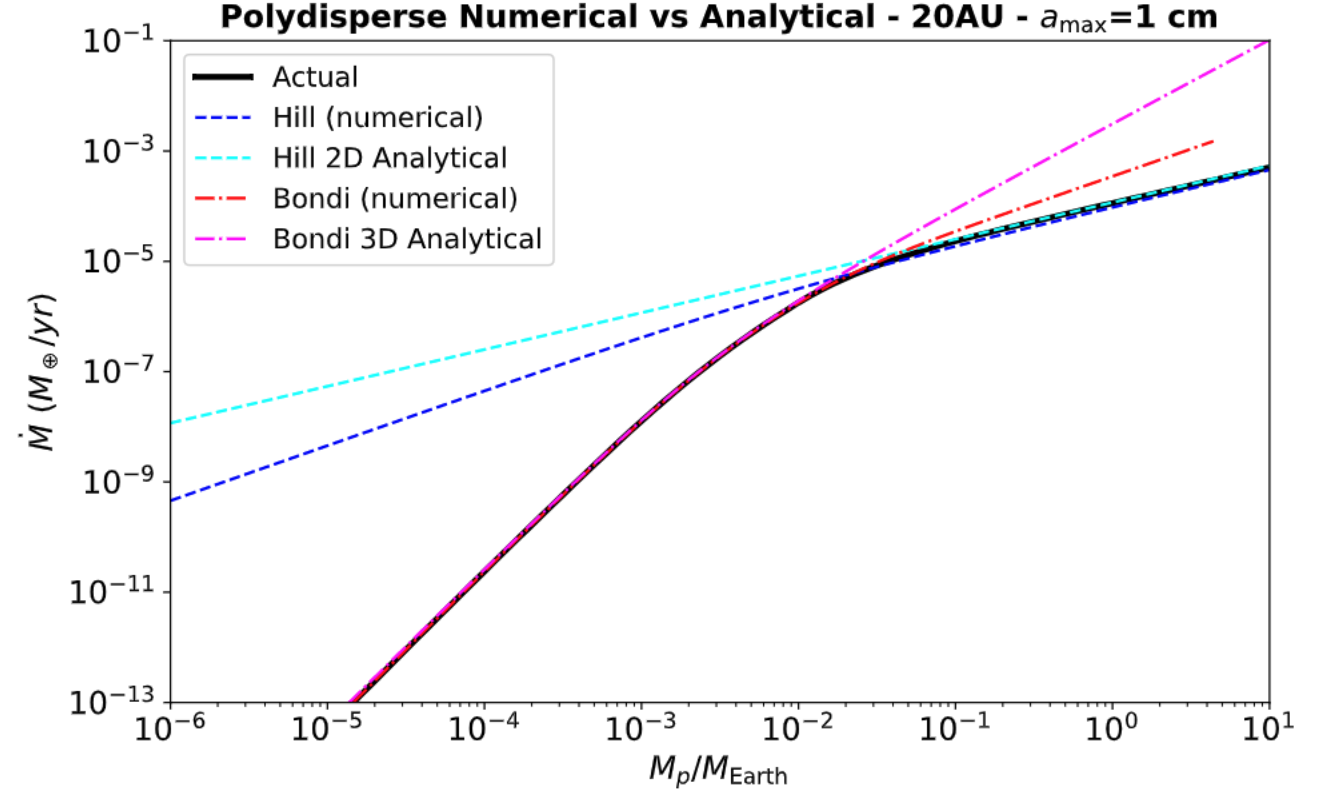
$$\dot{M}_{2D,Hill} = \frac{6(1-p)}{14-5q-3k} \left(\frac{St_{\max}}{0.1} \right)^{2/3} \Omega R_H^2 Z \Sigma_g.$$

$$\dot{M}_{3D,Bondi} \approx C_1 \frac{\gamma_l \left(\frac{b_1+1}{s}, j_1 a_{\max}^s \right)}{s j_1^{(b_1+1)/s}} + C_2 \frac{\gamma_l \left(\frac{b_2+1}{s}, j_2 a_{\max}^s \right)}{s j_2^{(b_2+1)/s}} + C_3 \frac{\gamma_l \left(\frac{b_3+1}{s}, j_3 a_{\max}^s \right)}{s j_3^{(b_3+1)/s}} + C_4 \frac{\gamma_l \left(\frac{b_4+1}{s}, j_4 a_{\max}^s \right)}{s j_4^{(b_4+1)/s}}.$$

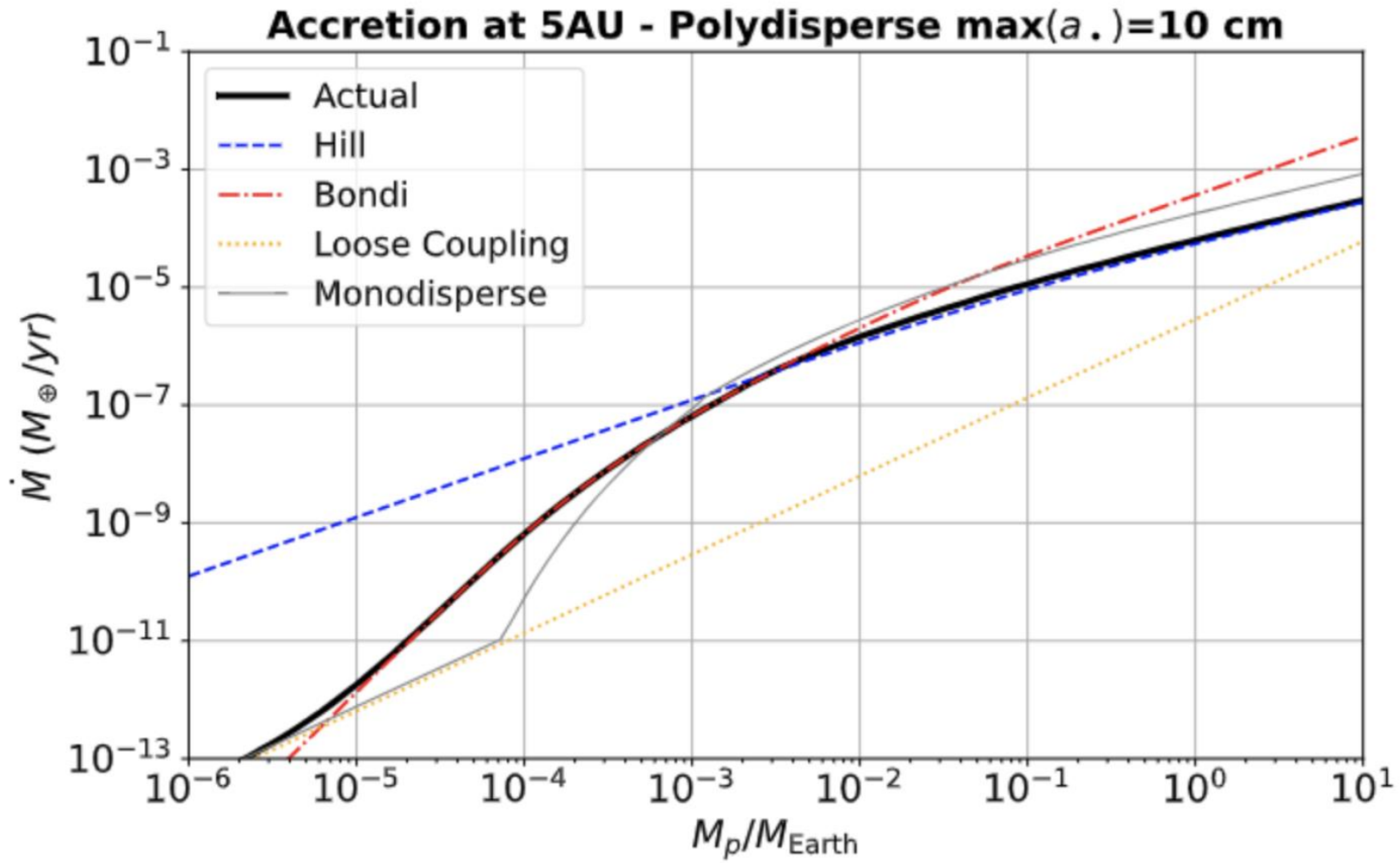
```
gamma11 = sp.special.gammainc((b1+1)/s, j1*a**s)*sp.special.gamma((b1+1)/s)
gamma12 = sp.special.gammainc((b2+1)/s, j2*a**s)*sp.special.gamma((b2+1)/s)
gamma13 = sp.special.gammainc((b3+1)/s, j3*a**s)*sp.special.gamma((b3+1)/s)
gamma14 = sp.special.gammainc((b4+1)/s, j4*a**s)*sp.special.gamma((b4+1)/s)
```

```
G1 = C1*gamma11/s/j1**((b1+1)/s)
G2 = C2*gamma12/s/j2**((b2+1)/s)
G3 = C3*gamma13/s/j3**((b3+1)/s)
G4 = C4*gamma14/s/j4**((b4+1)/s)
```

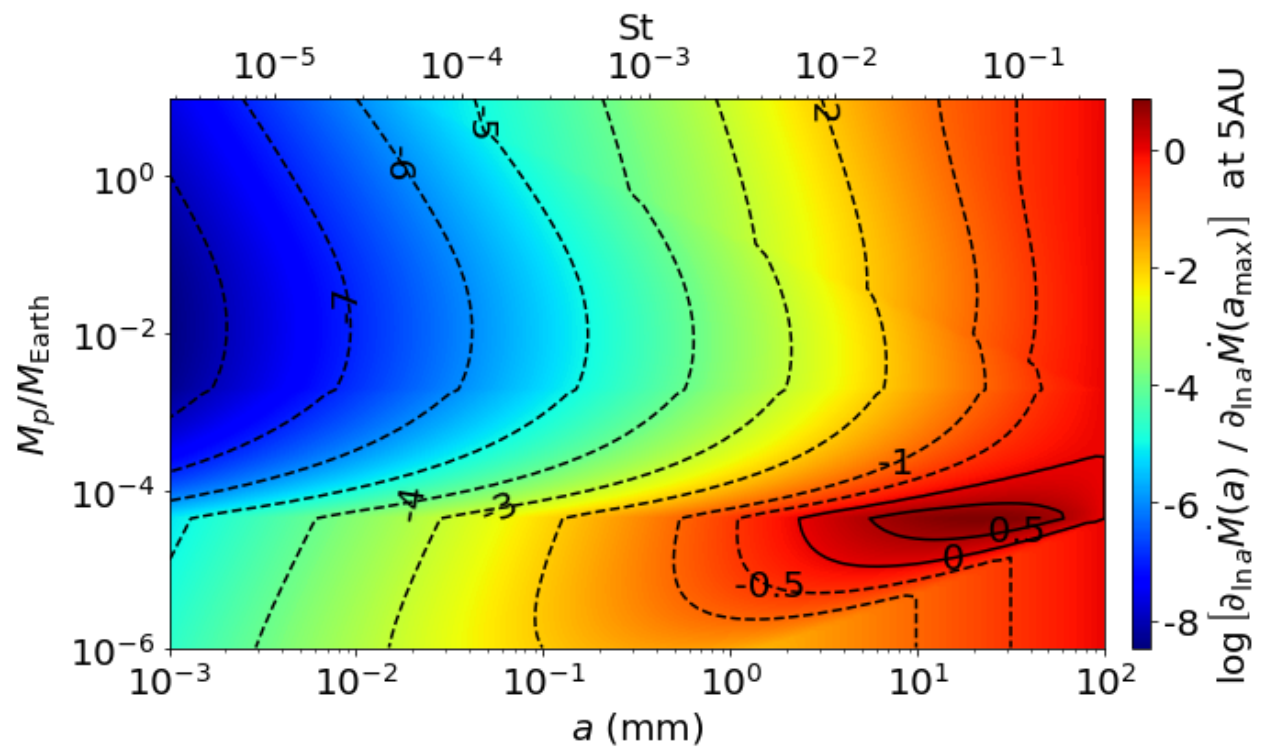
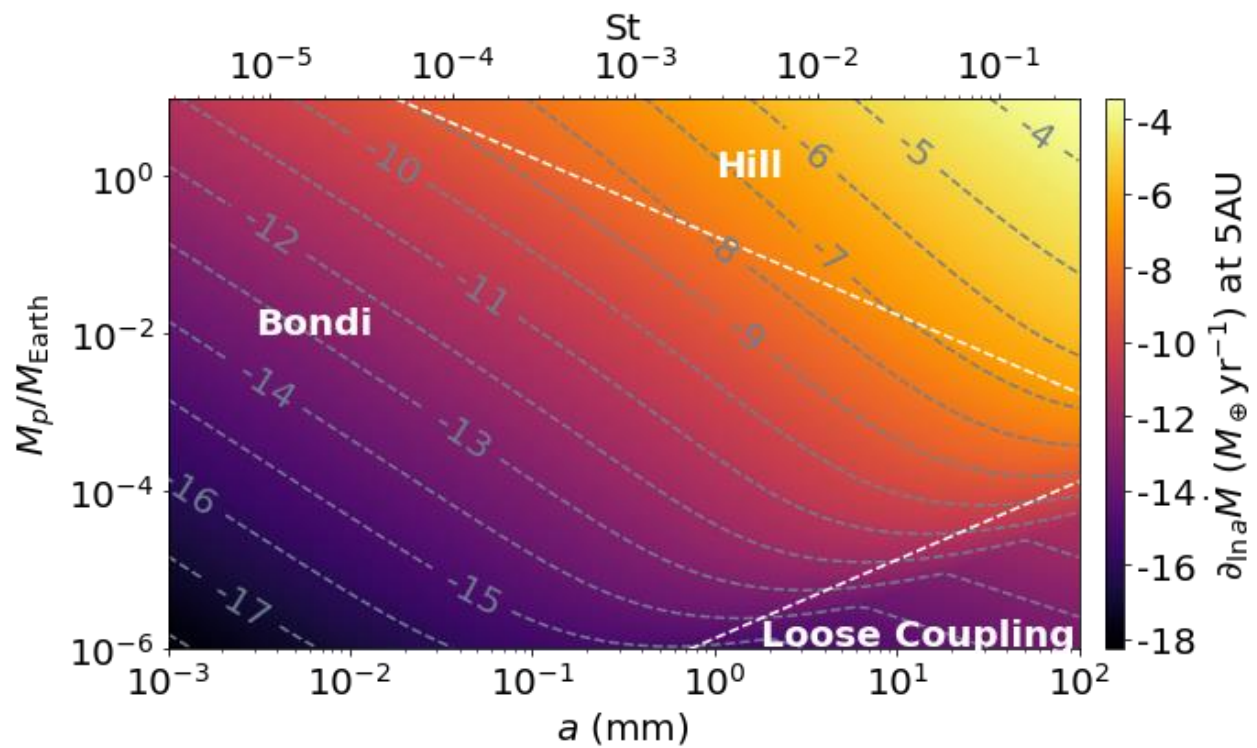
```
Mbondi3d = G1 + G2 + G3 + G4
```



Accretion Rates

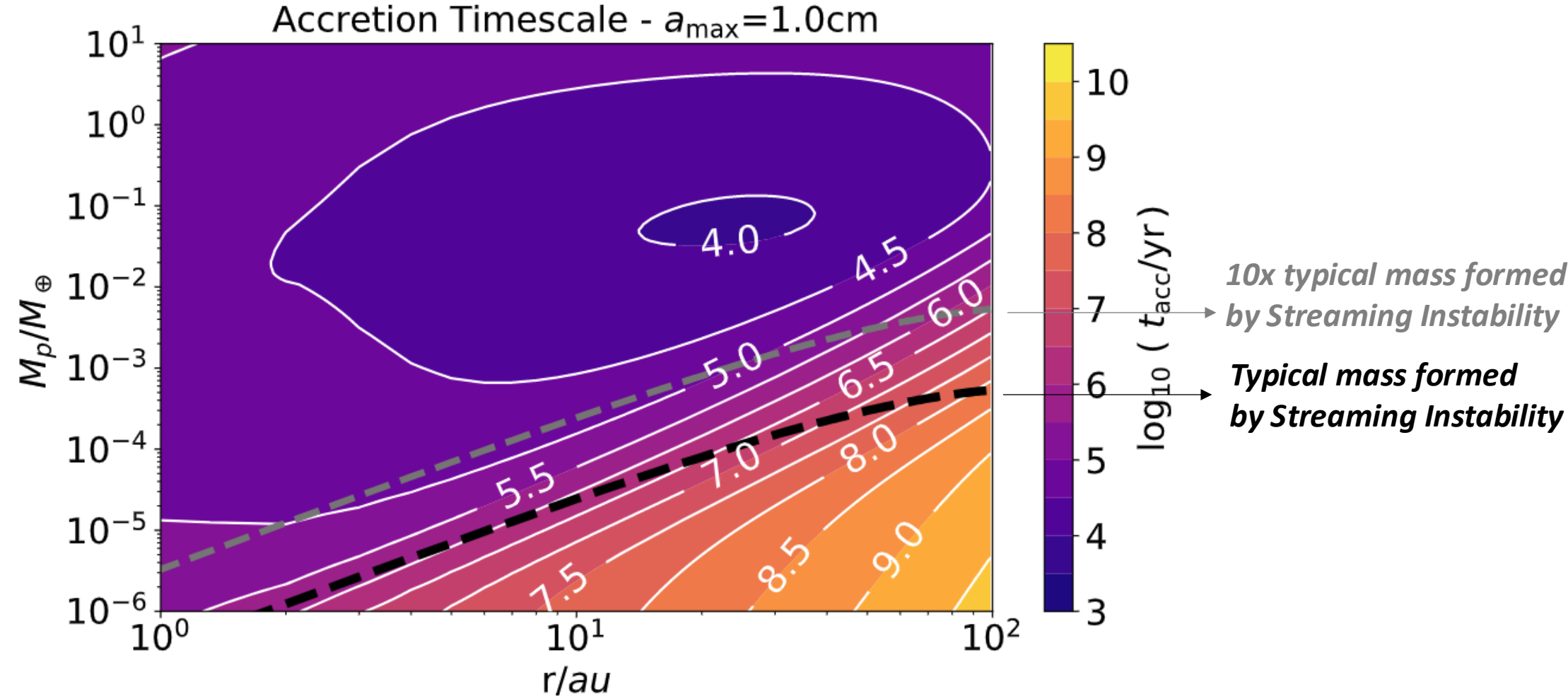


Accretion Rates

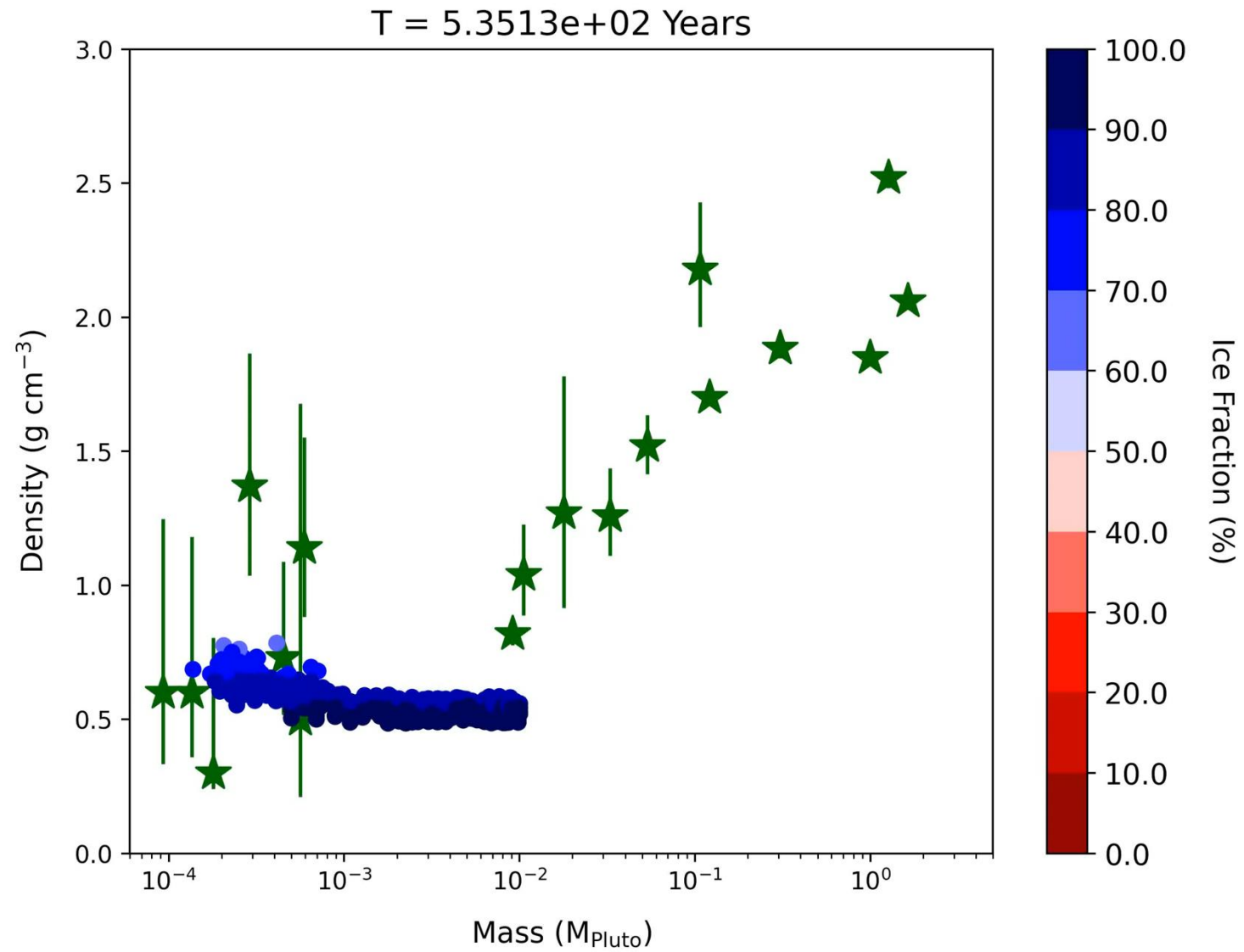


Accretion Timescales

Myr accretion timescales possible on top of planetesimals produces by Streaming Instability

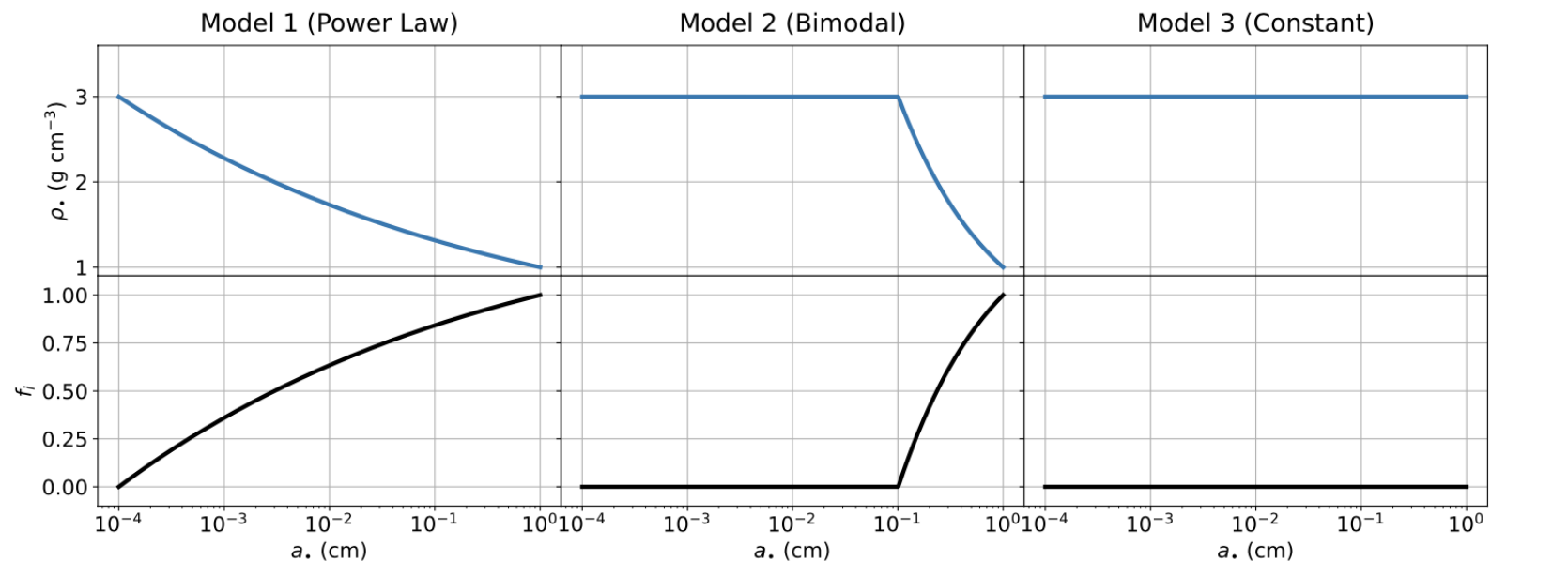


Growing Pluto by silicate pebble accretion

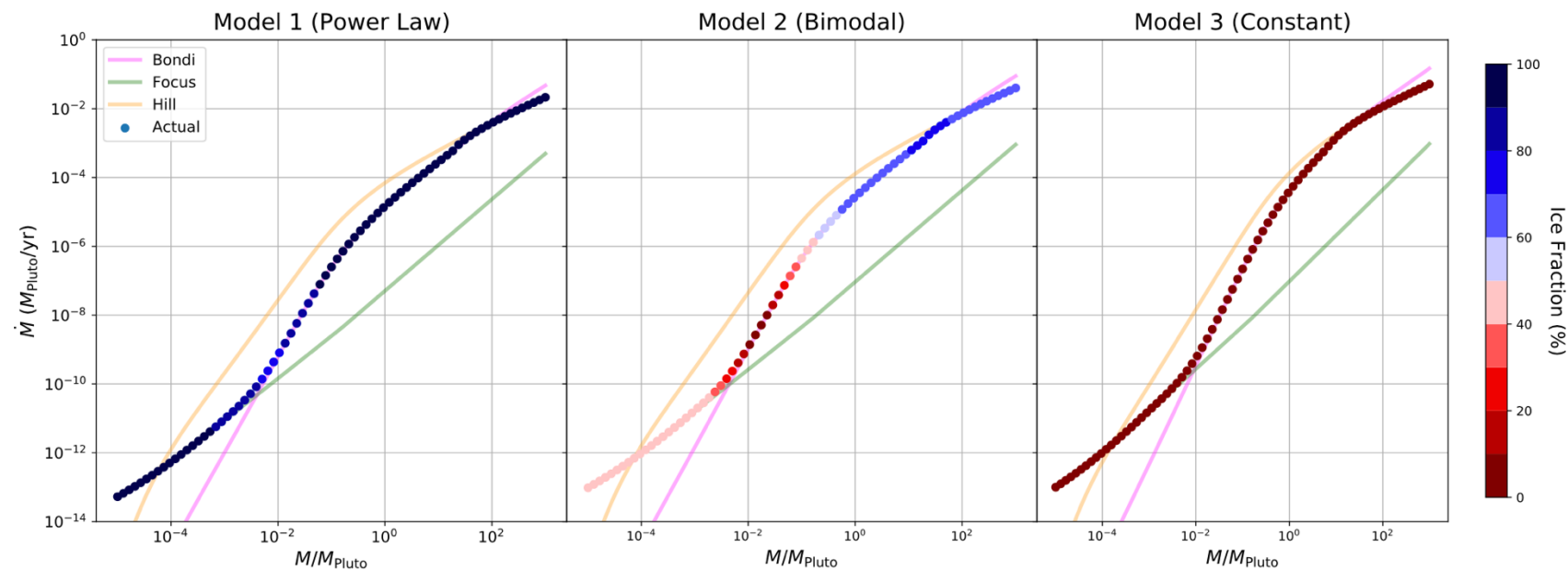


Pebble Internal Density

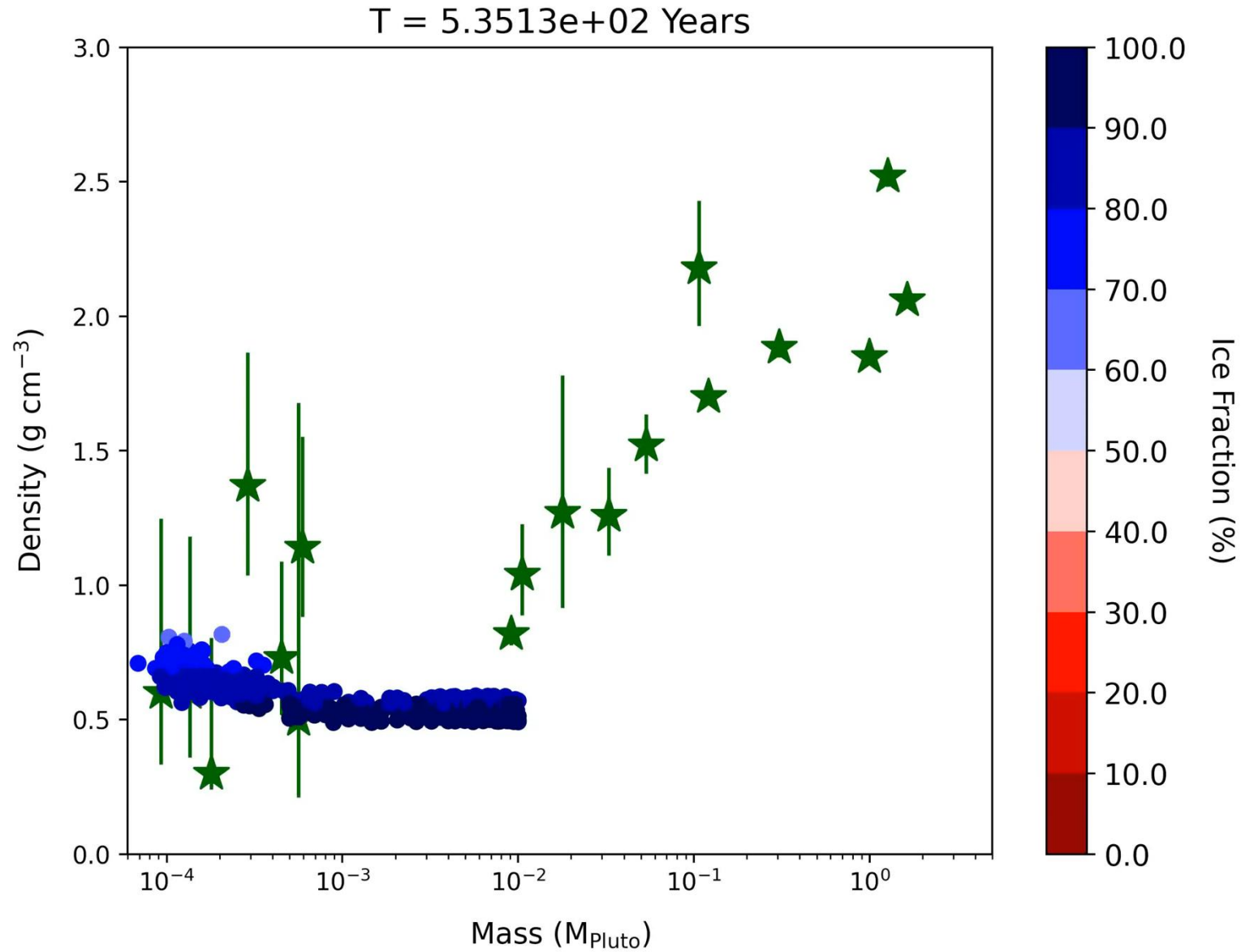
Ice Volume Fraction



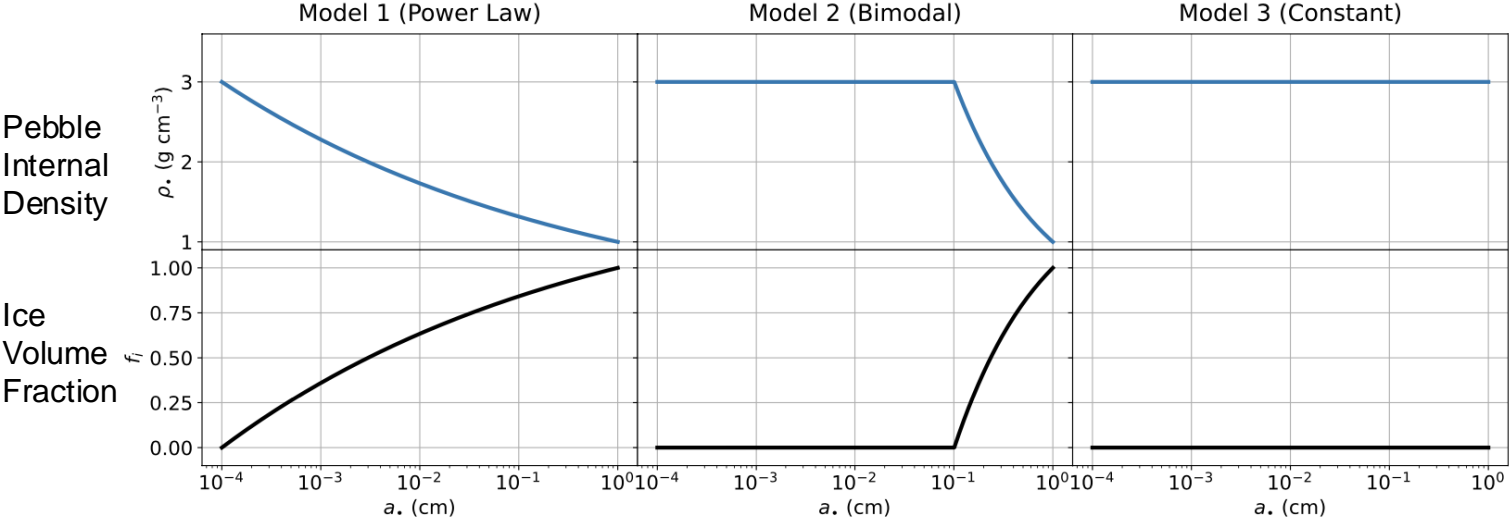
Mass Accretion rate



Growing Pluto by silicate pebble accretion



Resulting Densities vs Mass relations

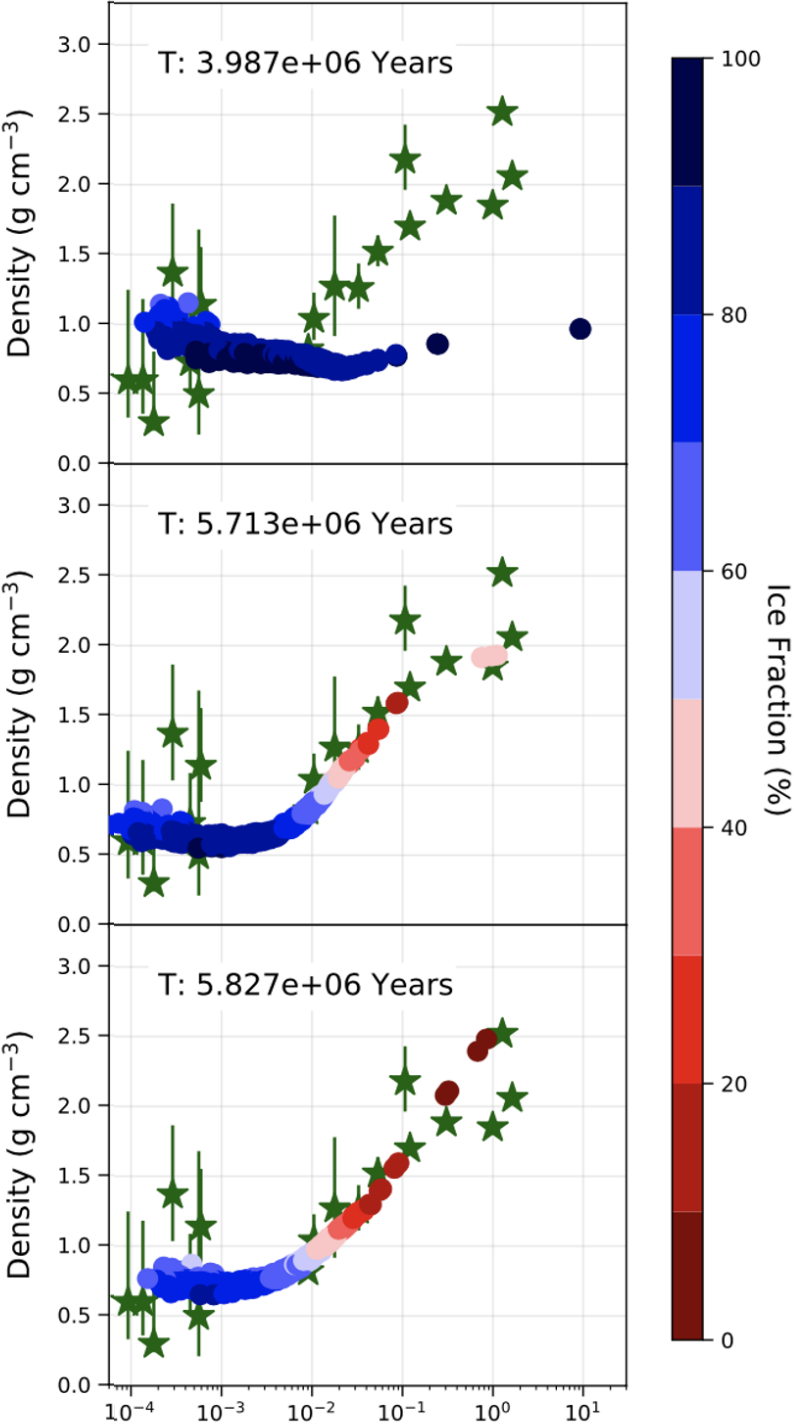


Canas+Lyra et al. 2024

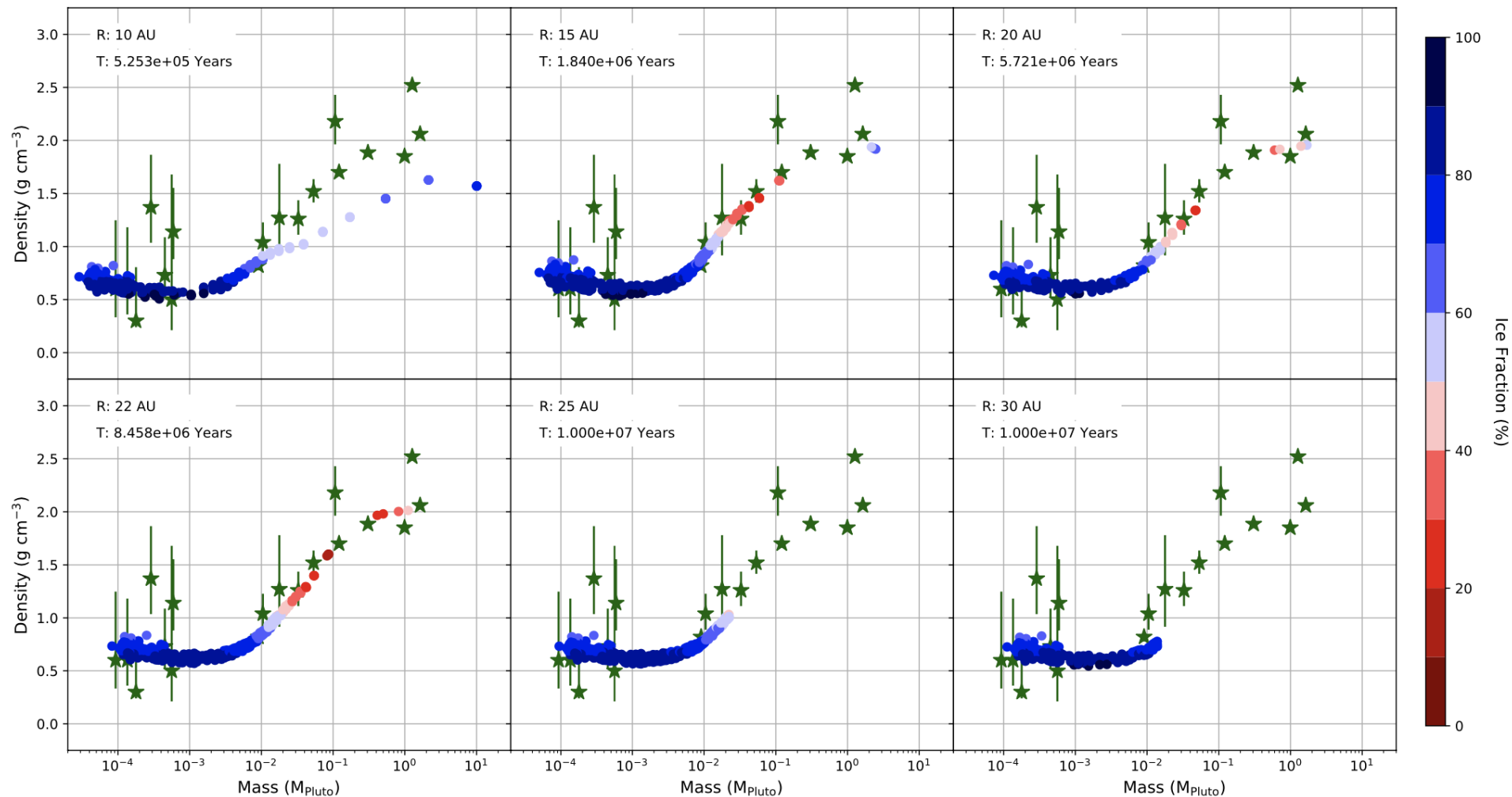
Model 1 (Power Law)

Model 2 (Bimodal)

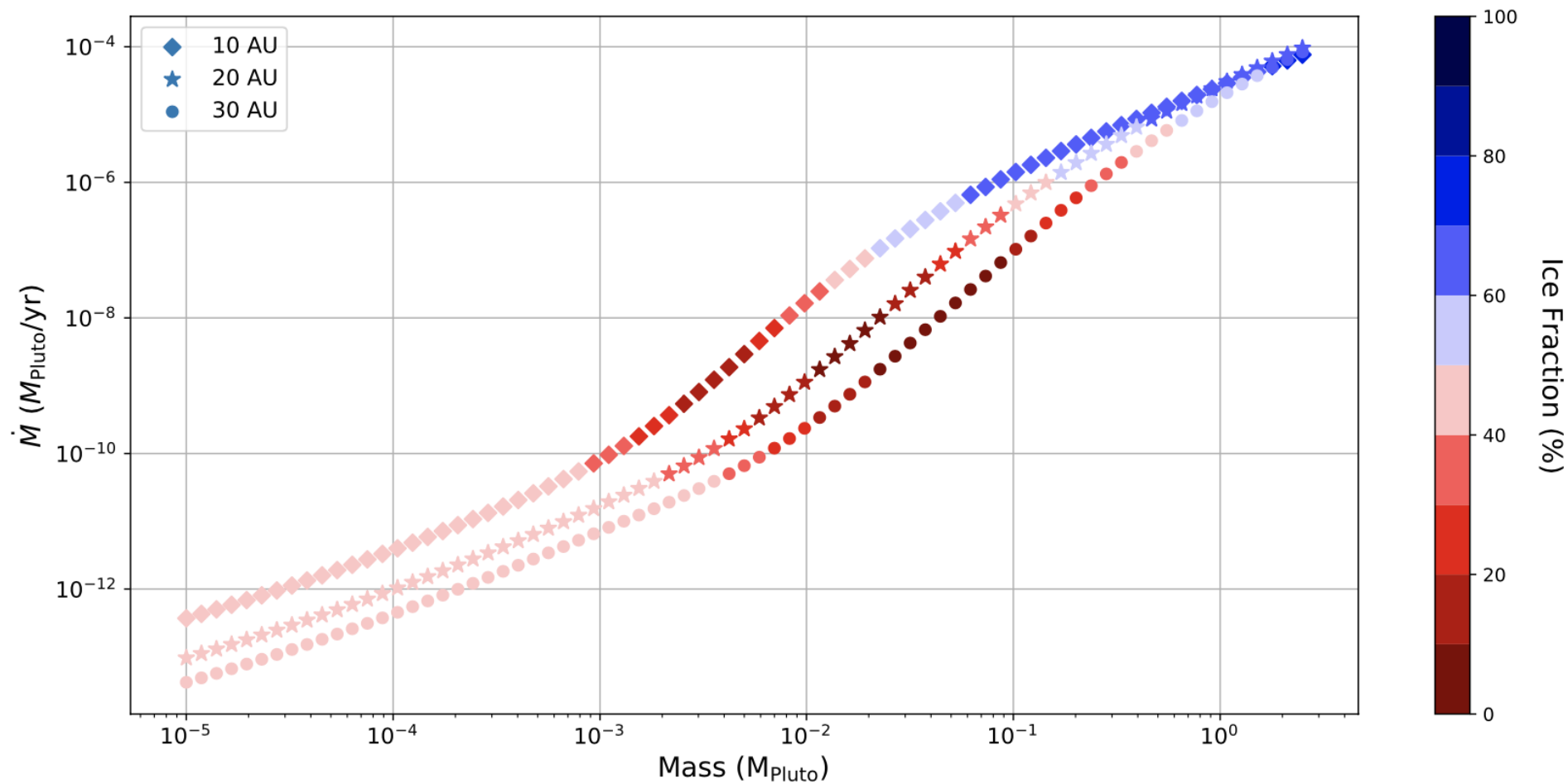
Model 3 (Constant)



Distance Range 15 - 25AU

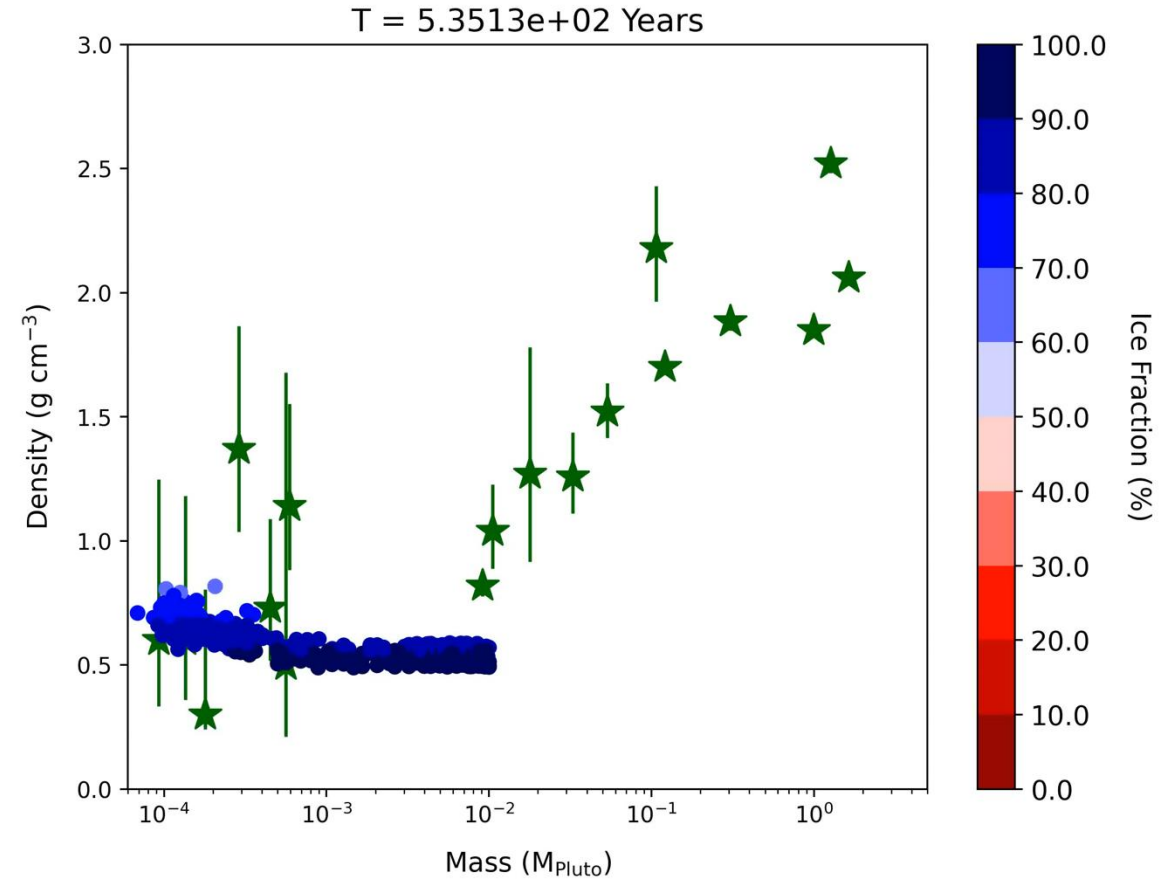


The window of silicate accretion



Conclusions

- Polydisperse Bondi accretion 1-2 orders of magnitude more efficient than monodisperse
 - Best accreted pebbles are those of drag time \sim Bondi time, not the largest ones
 - The largest ones dominate the mass budget, but accrete poorly
- Onset of Bondi accretion 1-2 orders of magnitude lower in mass compared to monodisperse
 - Bondi accretion possible on top of Streaming Instability planetary embryos within disk lifetime
 - Reaches 100-350km objects within Myr timescales
- Analytical solution to
 - Monodisperse general case
 - Polydisperse 2D Hill and 3D Bondi
- KBO density problem:
 - Two different pebble populations, maintained by ice desorption off small grains
 - Streaming instability: icy-rich small objects; nearly uniform composition
 - Polydisperse pebble accretion: silicate-rich larger objects; varied composition
 - Melting avoided by
 - ice-rich formation
 - ^{26}Al incorporated mostly in long ($>$ Myr) phase of silicate accretion
 - KBOs best reproduced between 15-25 AU



Positive Feedback: How a Synergy Between the Streaming Instability and Dust Coagulation Forms Planetesimals

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ABSTRACT

Context. One of the most important open questions in planet formation is how dust grains in a protoplanetary disk manage to overcome growth barriers and form the ~100km planet building blocks that we call planetesimals. There appears to be a gap between the largest grains that can be produce by coagulation, and the smallest grains that are needed for the streaming instability (SI) to form planetesimals.
Aims. Here we explore a novel hypothesis: That dust coagulation and the SI work in tandem. That they form a feedback loop where each one boosts the action of the other to bridge the gap between dust grains and planetesimals.
Methods. We develop a semi-analytical model of dust concentration due to the SI, and an analytic model of how the SI affects the fragmentation and radial drift barriers. We then combine those to model our proposed feedback loop.
Results. In the fragmentation-limited regime, we find a powerful synergy between the SI and dust growth that drastically increases both grain sizes and densities. We find that a midplane dust-to-gas ratio of $\epsilon \geq 0.3$ is a sufficient condition for the feedback loop to reach the planetesimal-forming region for turbulence values $10^{-4} \leq \alpha \leq 10^{-3}$ and grain sizes $0.01 \leq St \leq 0.1$. In contrast, the drift-limited regime only shows grain growth, without significant dust accumulation. Planet formation in the drift-limited portion of the disk may require other processes (particle traps) to halt radial drift.

Key words. planetesimals – planet formation



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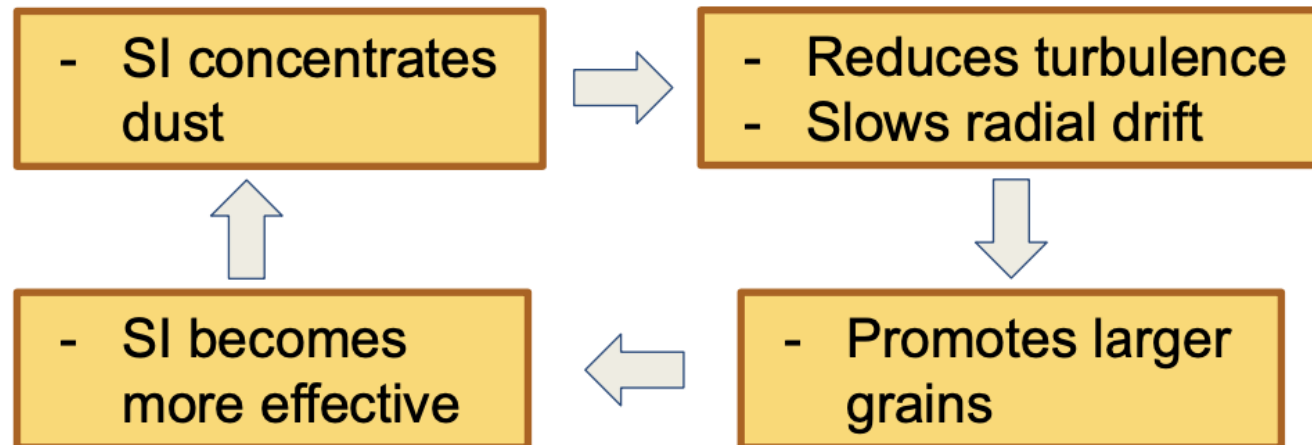
Positive Feedback – A Synergy Between the SI and Dust Coagulation

Motivation: The SI only forms planetesimals if

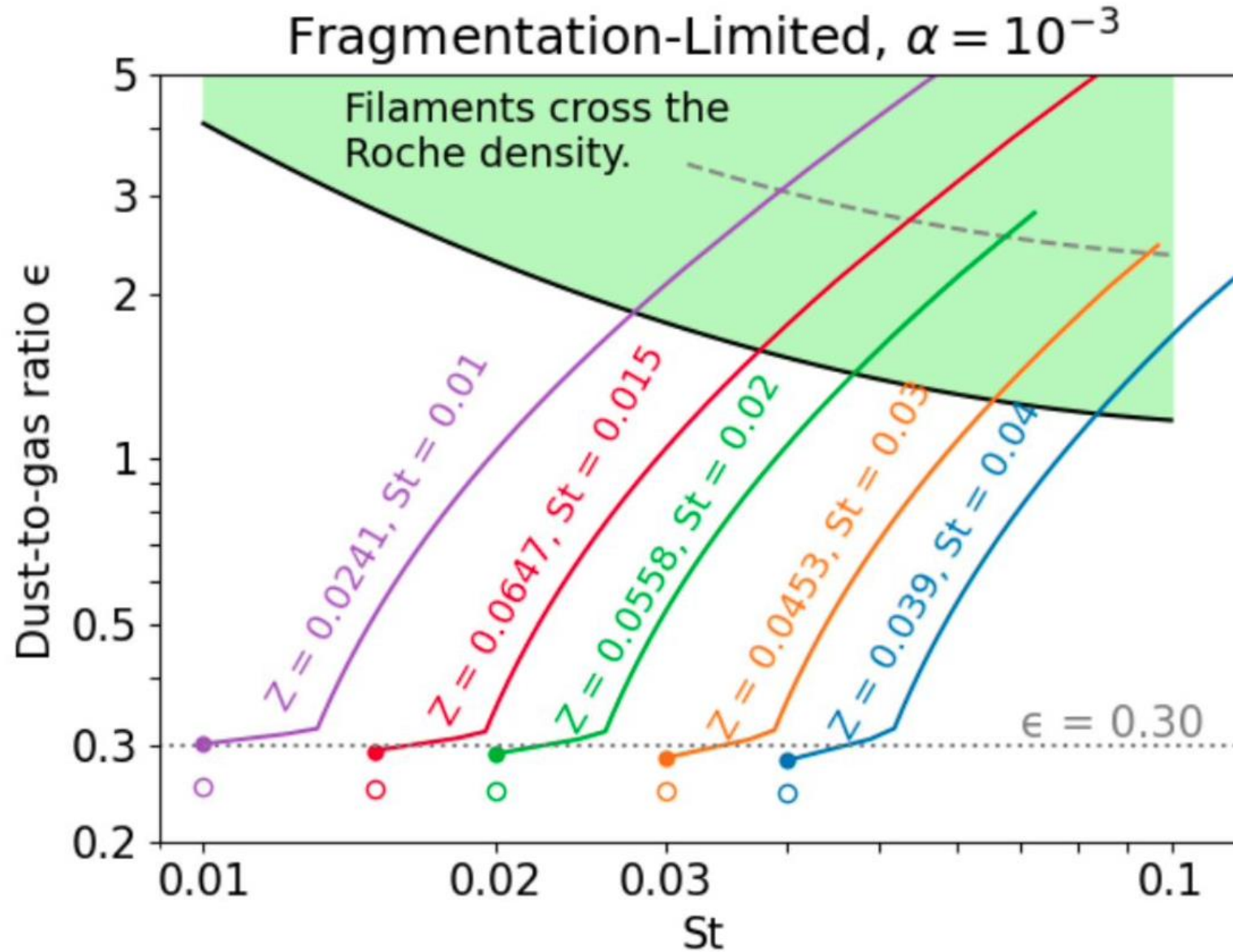
- Grains are larger than dust models predict, AND/OR
- Z is larger than dust models predict.

Our proposal: A **feedback loop** between the SI and dust coagulation

- The SI creates the conditions for better dust growth.
- Dust growth creates the conditions for better SI.



Positive Feedback – A Synergy Between the SI and Dust Coagulation



Summary

Positive Feedback: How a Synergy Between the Streaming Instability and Dust Coagulation Forms Planetesimals

- **Feedback loop** between the SI + dust coagulation largely overcomes the fragmentation barrier