Polydisperse Pebble Accretion Doing away with planetesimal accretion



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The size-density relationship of Kuiper Belt objects

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THE DENSITY OF MID-SIZED KUIPER BELT OBJECT 2002 UX25 AND THE FORMATION OF THE DWARF PLANETS

M. E. Brown

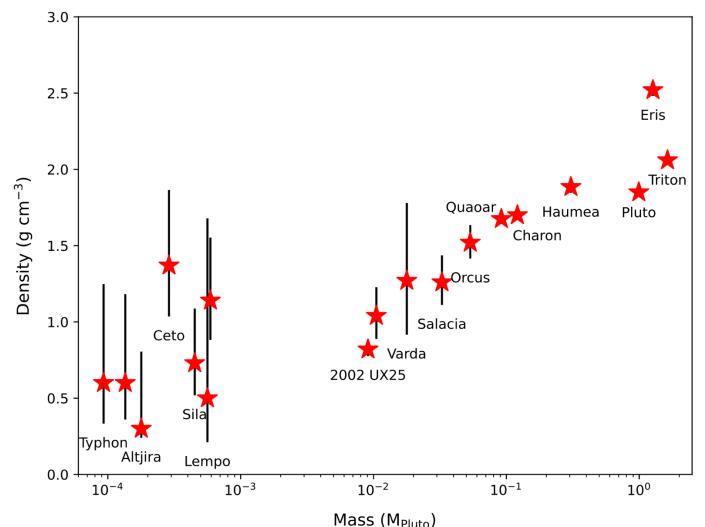
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ABSTRACT

The inferred low rock fraction of the 2002 UX25 system higher, from out invoking makes the formation of rock-rich larger objects difficult to in to increase in occurs for explain in any standard coagulation scenario. For example, and an object with the volume of Eris would require racterization, a diameter of assembling ~40 objects of the size of 2002 UX25. Yet the assembled object, even with the additional compression, would epresentation still have a density close to 1 g cm⁻³ rather than the 2.5 g cm⁻³ thigher from the size of Eris (Sicardy et al. 2011).

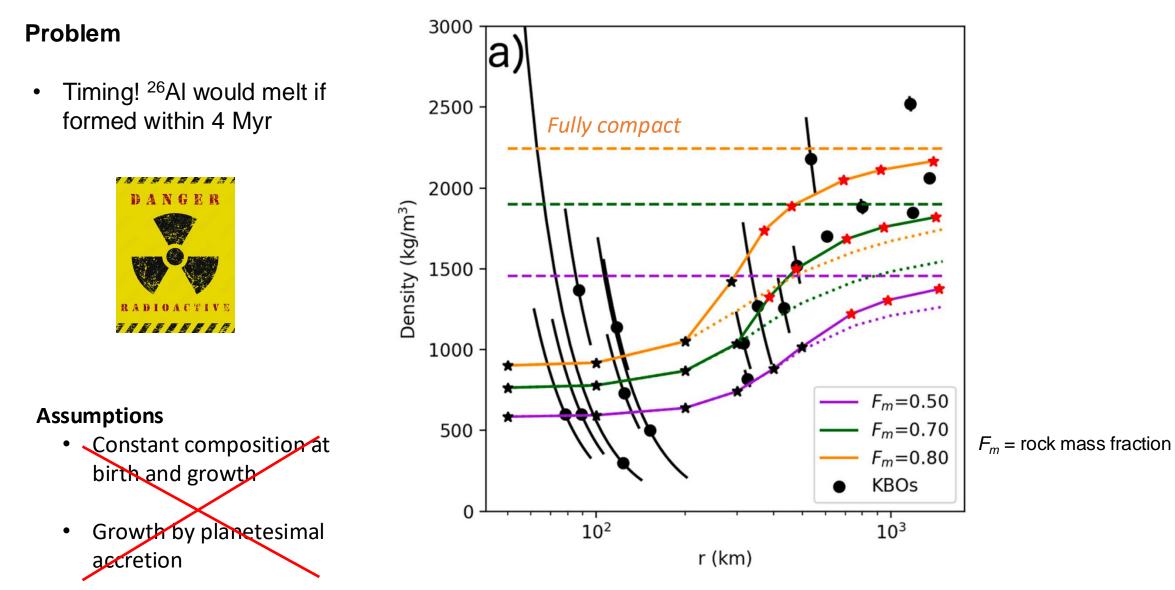
- Extremely low porosity;
- Biased sample;
- Compaction through giant impacts

None of these alternatives appears likely. We are left in the uncomfortable state of having no satisfying mechanism to explain the formation of the icy dwarf planets. While objects up to the size of 2002 UX25 can easily be formed through standard coagulation scenarios, the rock-rich larger bodies may require a formation mechanism separate from the rest of the Kuiper belt.



Data; Thomas (2000), Stansberry et al. (2006), Grundy et al. (2007), Brown et al. (2011), Stansberry et al. (2012), Brown (2013), Fornasier et al. (2013), Vilenius, et al. (2014), Nimmo et al. (2016), Ortiz et al. (2017), Brown and Butler (2017), Grundy et al. (2019), Morgado et al. (2023), Pereira et al. (2023).

Previous best bet: Porosity removal by gravitational compaction



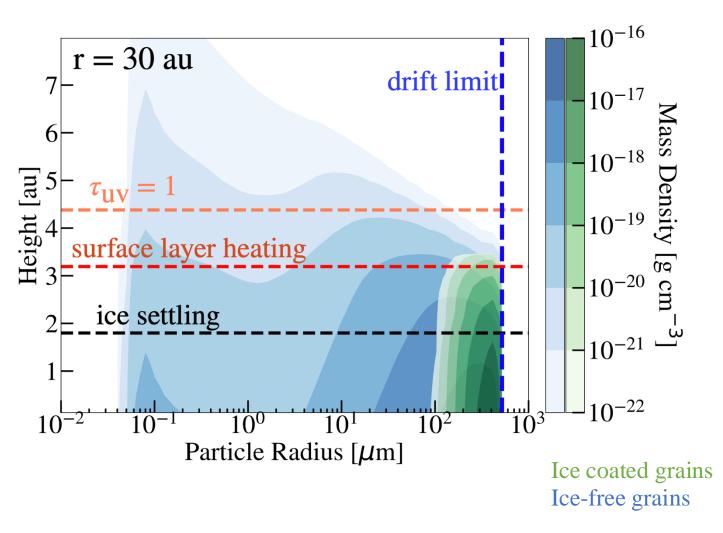
Bierson & Nimmo (2019)

Abandoning Constant Composition

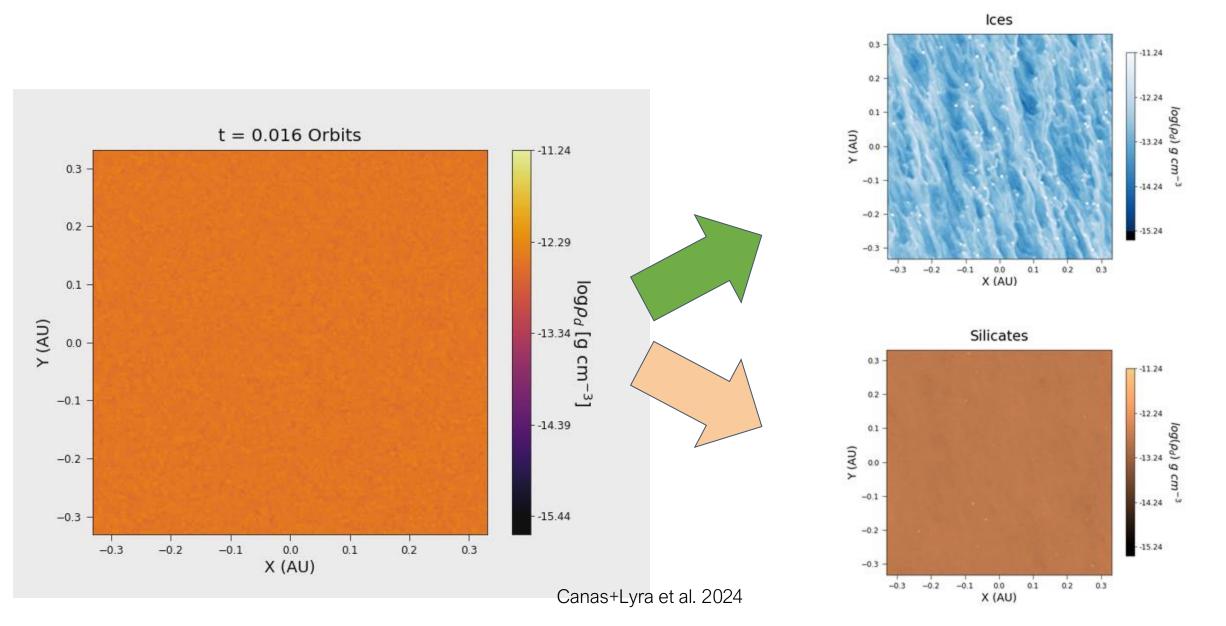
Heating and UV irradiation remove ice on Myr timescales (Harrison & Schoen 1967)

- Small grains lofted in the atmosphere lose ice
- Big grains are shielded and remain icy.

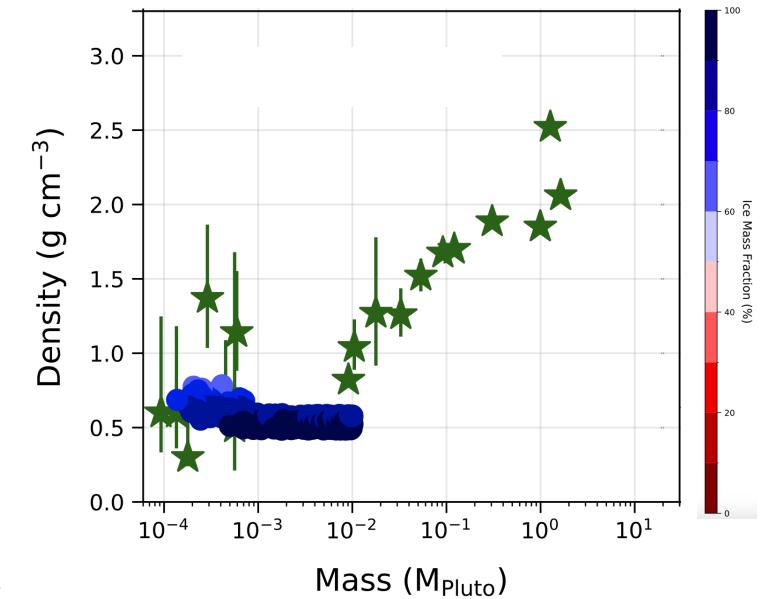




Split into icy and silicate pebbles



The first planetesimals are icy



The first planetesimals won't melt

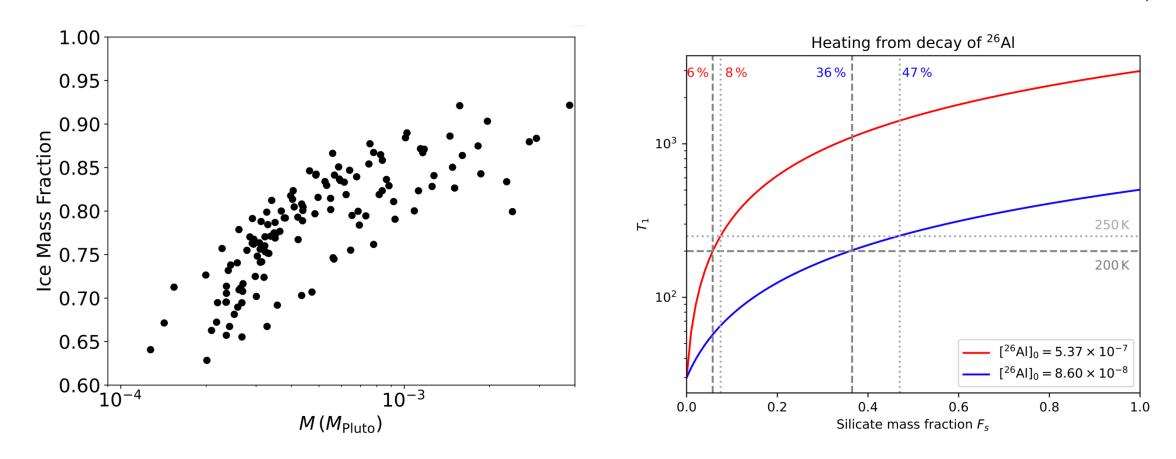
$$\mathcal{H} = \rho F_s [^{26} \text{Al}]_0 \mathcal{H}_0 e^{-\lambda t}$$

$$Q(t) = V \int_0^t \mathcal{H}(t') dt'$$

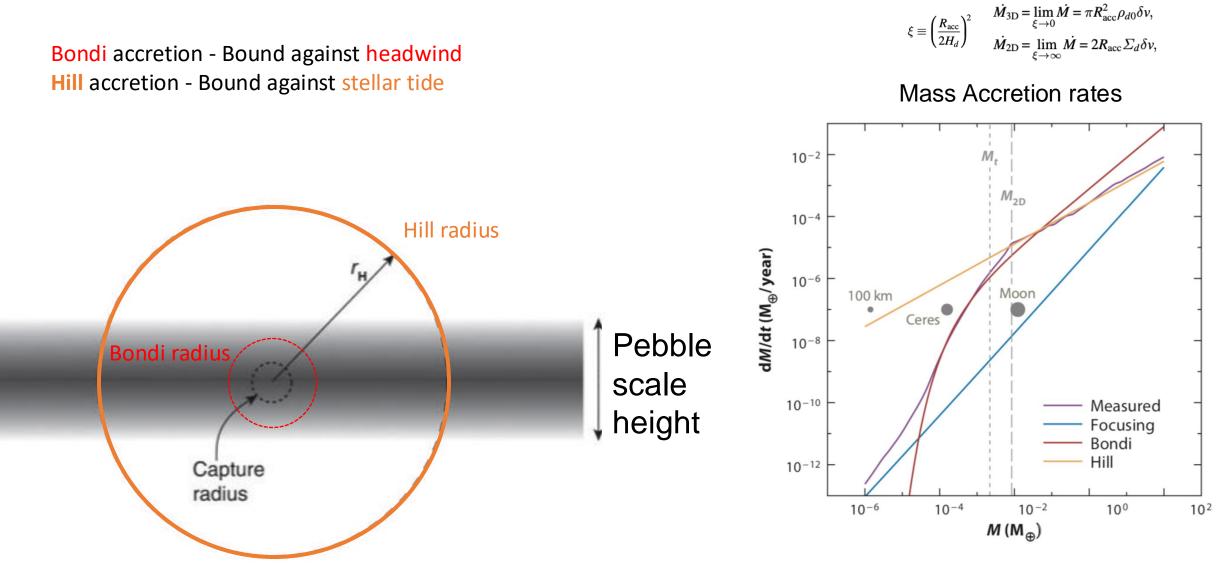
$$= M_p F_s [^{26} \text{Al}]_0 \mathcal{H}_0 \lambda^{-1} \left(1 - e^{-\lambda t}\right)$$

$$Q = M_p c_p \Delta T$$

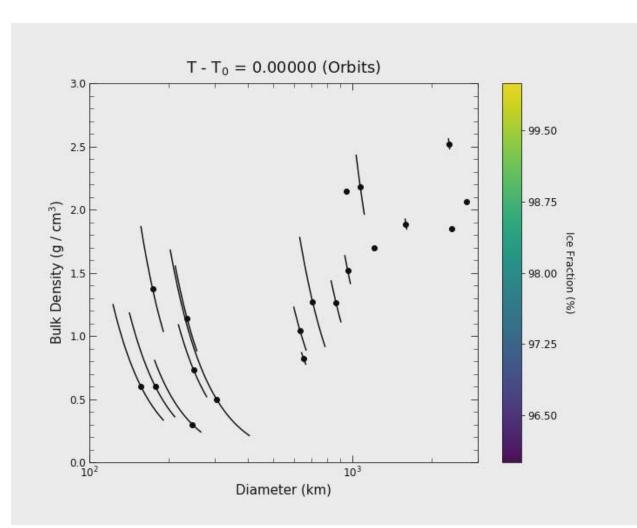
$$\Delta T = F_s [^{26} \text{Al}]_0 \mathcal{H}_0 \lambda^{-1} c_p^{-1}$$

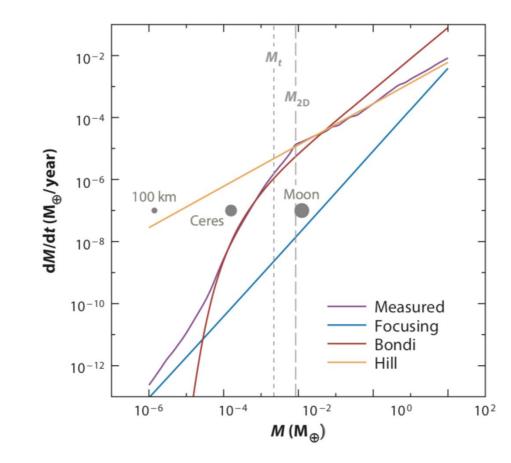


Pebble Accretion: Geometric, Bondi, and Hill regime

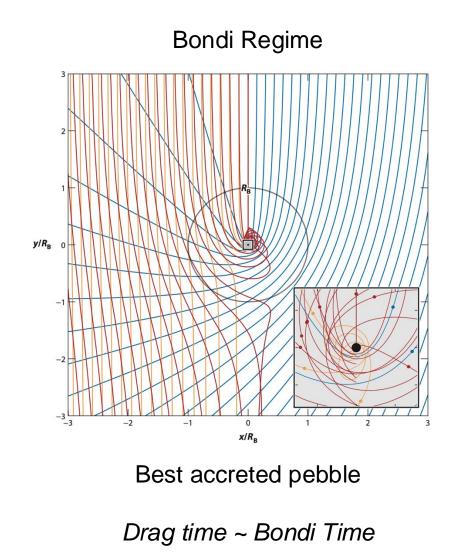


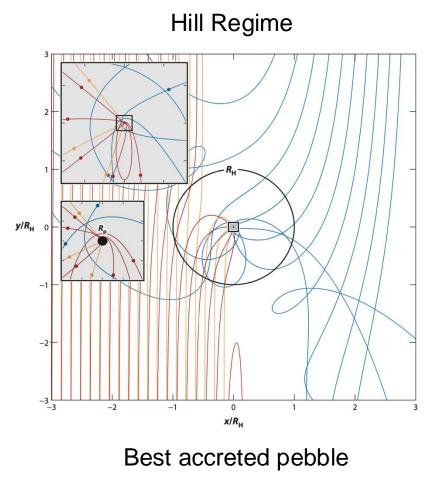
Integrate pebble accretion





Pebble Accretion: Pebbles of different size accrete differently





Drag time ~ Orbital Time

$$\rho_d(a, z) = \int_0^a m(a') F(a', z) \, da'.$$

$$F(a, z) \equiv f(a) \, e^{-z^2/2H_d^2},$$

$$f(a) = \frac{3(1-p)Z\Sigma_g}{2^{5/2}\pi^{3/2}H_g\rho_{\bullet}^{(0)}a_{\max}^{4-k}} \sqrt{1 + a\frac{\pi}{2}\frac{\rho_{\bullet}(a)}{\Sigma_g\alpha}} \, a^{-k}$$

$$\begin{split} S &\equiv \frac{1}{\pi R_{\rm acc}^2} \int_{-R_{\rm acc}}^{R_{\rm acc}} 2\sqrt{R_{\rm acc}^2 - z^2} \, \exp\left(-\frac{z^2}{2H_d^2}\right) dz, \\ W(a) &= \frac{3(1-p)Z\Sigma_g}{4\pi \rho_{\bullet}^{(0)} a_{\rm max}^{4-k}} \, a^{-k}, \\ \delta v &\equiv \Delta v + \Omega R_{\rm acc}, \\ R_{\rm acc} &\equiv \hat{R}_{\rm acc} \exp\left[-\chi(\tau_f/t_p)^{\gamma}\right], \end{split} \qquad \hat{R}_{\rm acc}^{({\rm Bondi})} &= \left(\frac{4\tau_f}{t_{\rm B}}\right)^{1/2} R_{\rm B}, \\ \frac{\partial \Sigma_d(a)}{\partial a} \propto a^{-p}; \\ \rho_{\bullet} \propto a^{-q}; \qquad t_p \equiv \frac{GM_p}{(\Delta v + \Omega R_{\rm H})^3} \end{split}$$

$$\dot{M}(a) = \int_0^a \frac{\partial \dot{M}(a')}{\partial a'} da',$$
$$\frac{\partial \dot{M}(a)}{\partial a} = \pi R_{\rm acc}^2(a) \,\delta v(a) S(a) m(a) f(a).$$

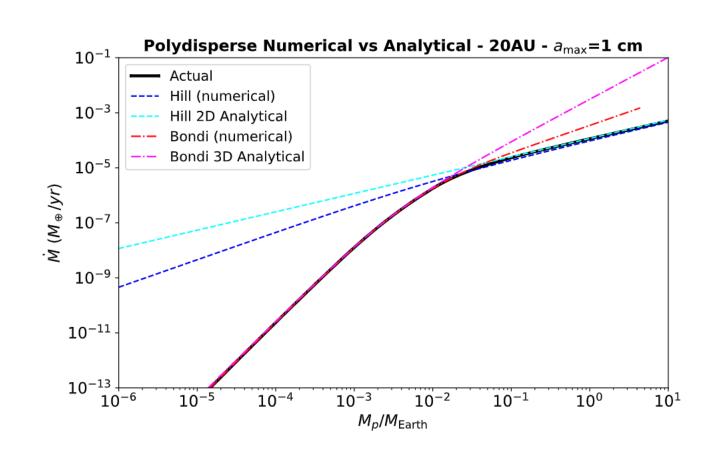
$$\dot{M}_{2D, \text{ Hill}} = 2 \times 10^{2/3} \Omega R_H^2 \int_0^{a_{\text{max}}} \operatorname{St}(a)^{2/3} m(a) W(a) \, da.$$

$$\dot{M}_{3D, \text{ Bondi}} = \frac{4\pi R_B \Delta v^2}{\Omega} \times \int_0^{a_{\text{max}}} \operatorname{St} e^{-2\psi} m(a) f(a) \times \left[1 + 2 \left(\operatorname{St} \frac{\Omega R_B}{\Delta v} \right)^{1/2} e^{-\psi} \right] da, \qquad \psi \equiv \chi [\operatorname{St}/(\Omega t_p)]^{\gamma}.$$

Lyra et al. 2023

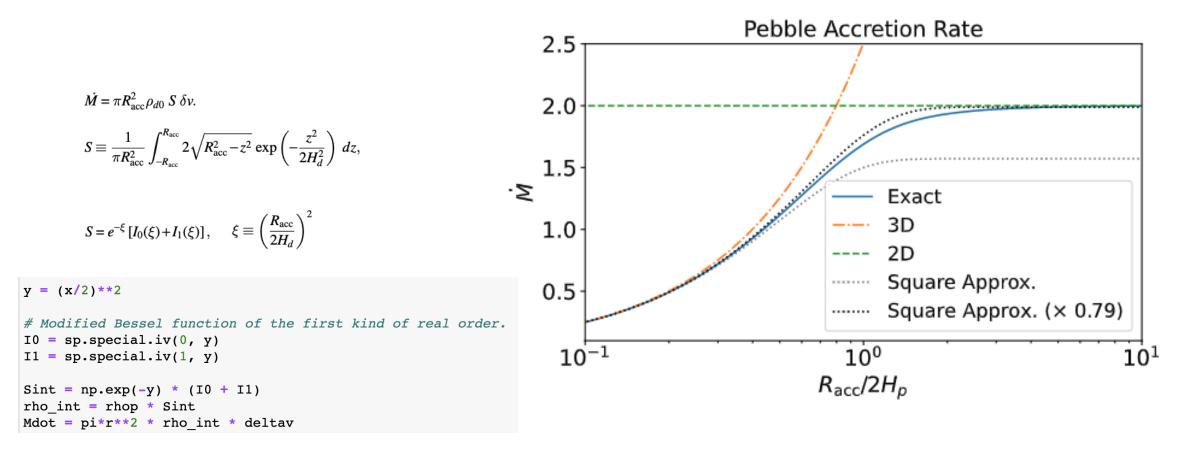
$$\begin{split} \varepsilon &= \left(\frac{R_{\rm acc}}{2H_d}\right)^2 \quad \begin{array}{l} \dot{M}_{\rm 3D} = \lim_{\substack{\xi \to 0}} \dot{M} = \pi R_{\rm acc}^2 \rho_{d0} \delta v, \\ \dot{M}_{\rm 2D} &= \lim_{\substack{\xi \to \infty}} \dot{M} = 2R_{\rm acc} \Sigma_d \delta v, \\ \dot{M}_{\rm 2D} = \lim_{\substack{\xi \to \infty}} \dot{M} = 2R_{\rm acc} \Sigma_d \delta v, \\ &\text{Lambrechts & Johansen (2012)} \end{split} \\ \hline \\ \begin{array}{l} \dot{P} Olydisperse \ (multiple \ species) \\ \dot{M}_{\rm 2D,Hill} = \frac{6(1-p)}{14-5q-3k} \left(\frac{St_{\rm max}}{0.1}\right)^{2/3} \Omega R_H^2 Z \Sigma_g. \\ \dot{M}_{\rm 3D,Bondi} \approx C_1 \frac{\gamma_l \left(\frac{b_l+1}{s}, j_l a_{\rm max}^s\right)}{s j_1^{(b_l+1)/s}} + C_2 \frac{\gamma_l \left(\frac{b_2+1}{s}, j_2 a_{\rm max}^s\right)}{s j_2^{(b_2+1)/s}} + \\ &C_3 \frac{\gamma_l \left(\frac{b_3+1}{s}, j_3 a_{\rm max}^s\right)}{s j_3^{(b_3+1)/s}} + C_4 \frac{\gamma_l \left(\frac{b_4+1}{s}, j_4 a_{\rm max}^s\right)}{s j_4^{(b_4+1)/s}}, \end{split}$$

Lyra et al. (2023)

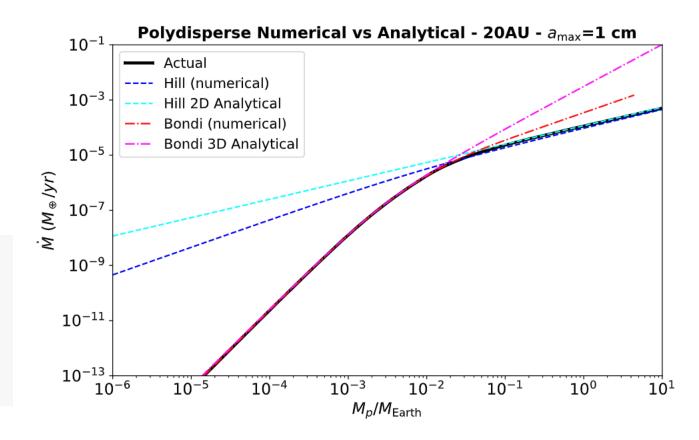


Lyra et al. 2023

Analytical Solution for General Monodisperse (single species) Pebble Accretion



Analytical Solutions for 2D and 3D Polydisperse (multi-species) Pebble Accretion



$$\dot{M}_{\rm 2D,Hill} = \frac{6(1-p)}{14-5q-3k} \left(\frac{\mathrm{St}_{\mathrm{max}}}{0.1}\right)^{2/3} \Omega R_H^2 Z \Sigma_g.$$

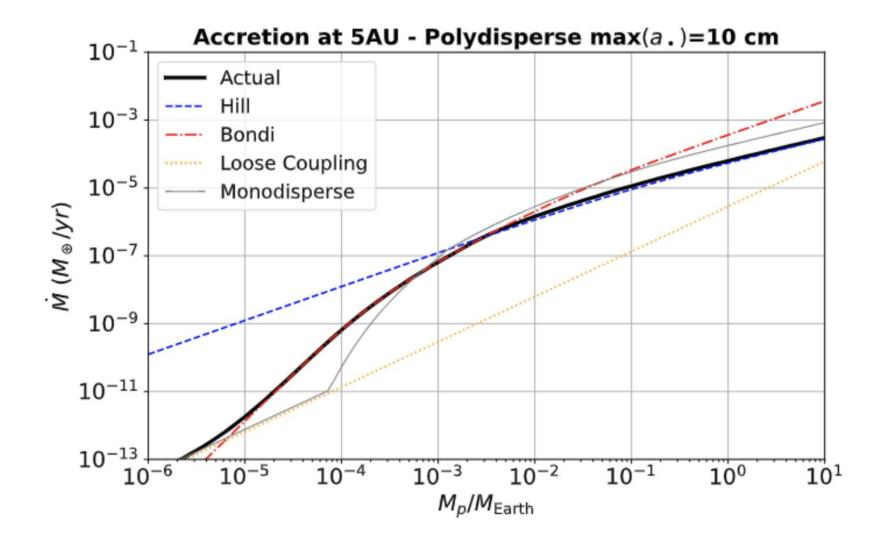
$$\dot{M}_{3\mathrm{D,Bondi}} \approx C_1 \frac{\gamma_l \left(\frac{b_1+1}{s}, j_1 a_{\mathrm{max}}^s\right)}{s j_1^{(b_1+1)/s}} + C_2 \frac{\gamma_l \left(\frac{b_2+1}{s}, j_2 a_{\mathrm{max}}^s\right)}{s j_2^{(b_2+1)/s}} + C_3 \frac{\gamma_l \left(\frac{b_3+1}{s}, j_3 a_{\mathrm{max}}^s\right)}{s j_3^{(b_3+1)/s}} + C_4 \frac{\gamma_l \left(\frac{b_4+1}{s}, j_4 a_{\mathrm{max}}^s\right)}{s j_4^{(b_4+1)/s}}$$

gammal1 = sp.special.gammainc((b1+1)/s,j1*a**s)*sp.special.gamma((b1+1)/s)
gammal2 = sp.special.gammainc((b2+1)/s,j2*a**s)*sp.special.gamma((b2+1)/s)
gammal3 = sp.special.gammainc((b3+1)/s,j3*a**s)*sp.special.gamma((b3+1)/s)
gammal4 = sp.special.gammainc((b4+1)/s,j4*a**s)*sp.special.gamma((b4+1)/s)

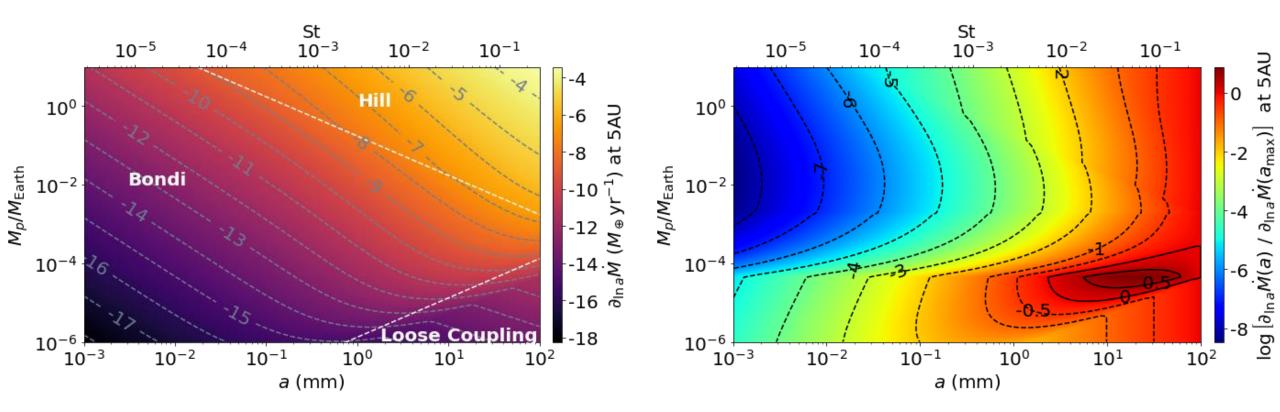
G1 = C1*gammal1/s/j1**((b1+1)/s) G2 = C2*gammal2/s/j2**((b2+1)/s) G3 = C3*gammal3/s/j3**((b3+1)/s) G4 = C4*gammal4/s/j4**((b4+1)/s)

Mbondi3d = G1 + G2 + G3 + G4

Accretion Rates

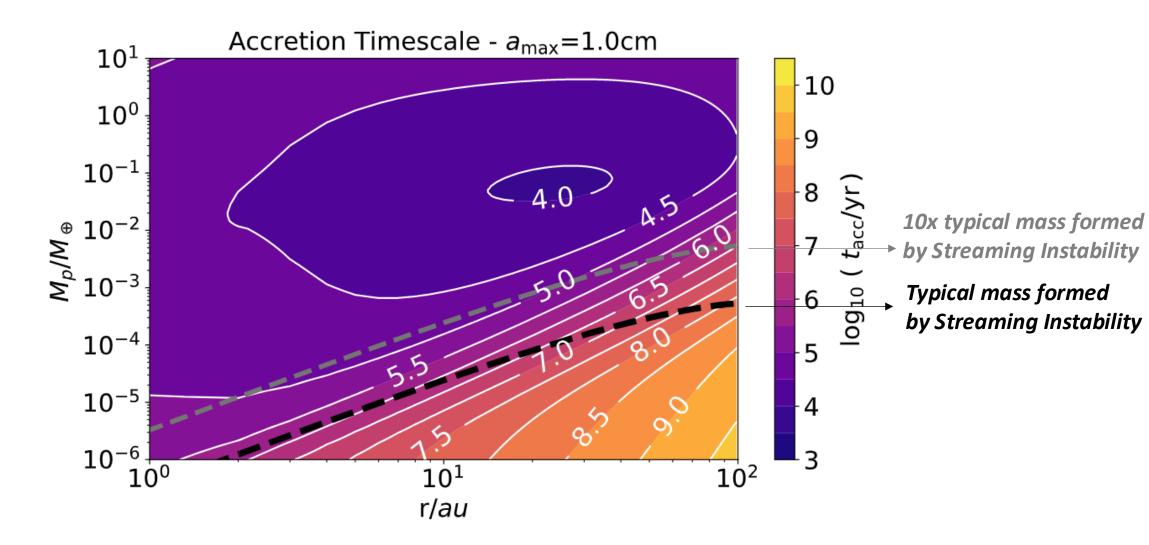


Accretion Rates

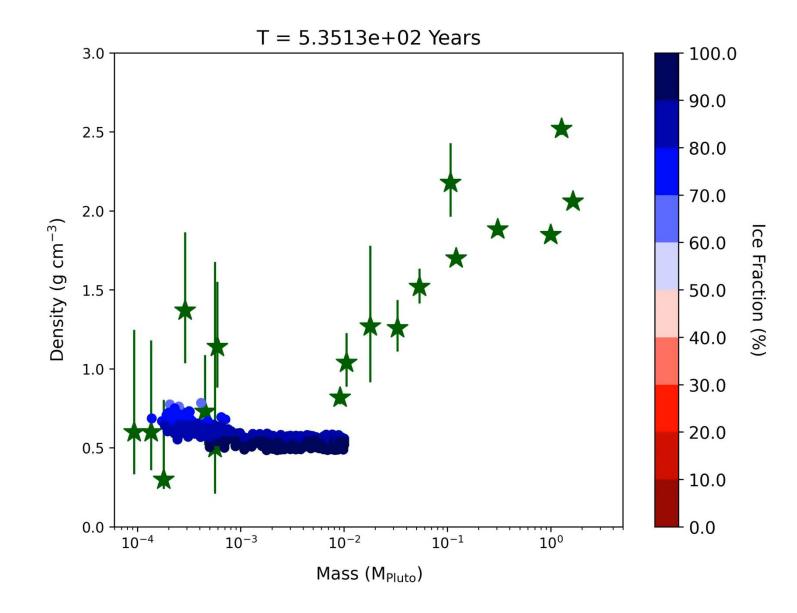


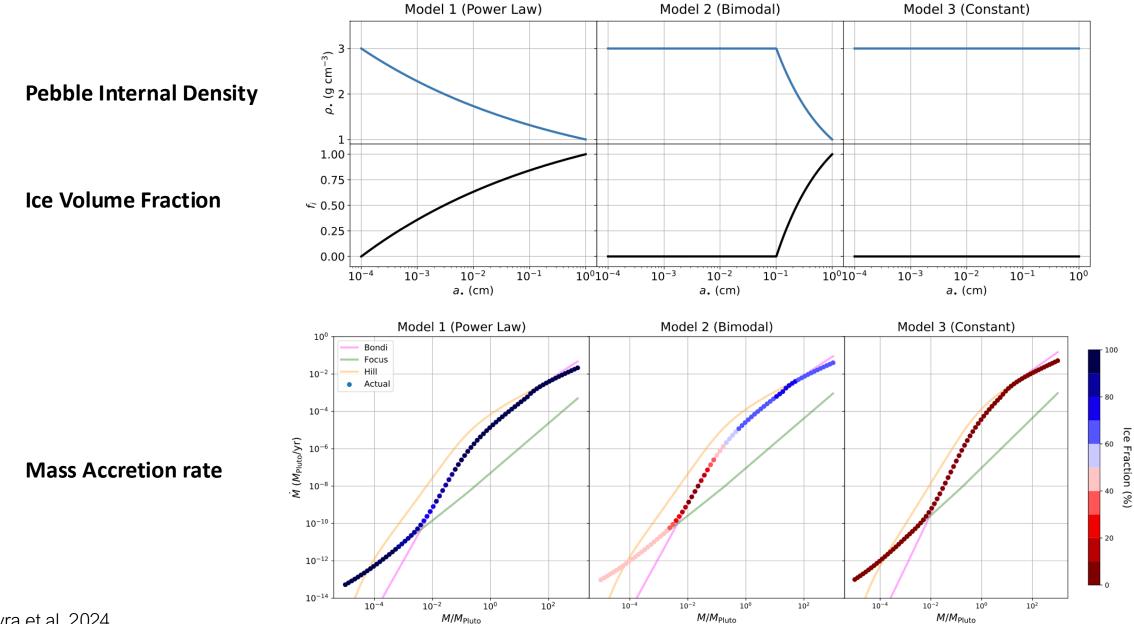
Accretion Timescales

Myr accretion timescales possible on top of planetesimals produces by Streaming Instability

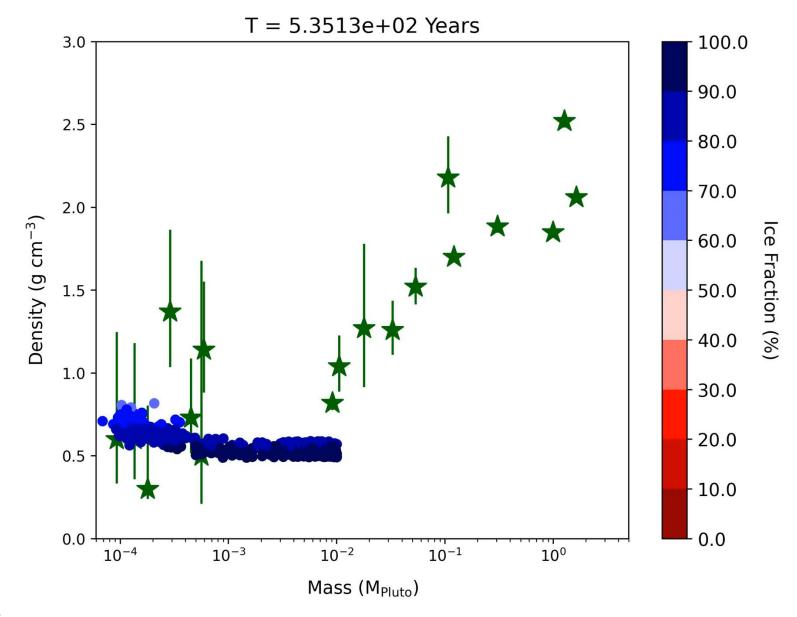


Growing Pluto by silicate pebble accretion

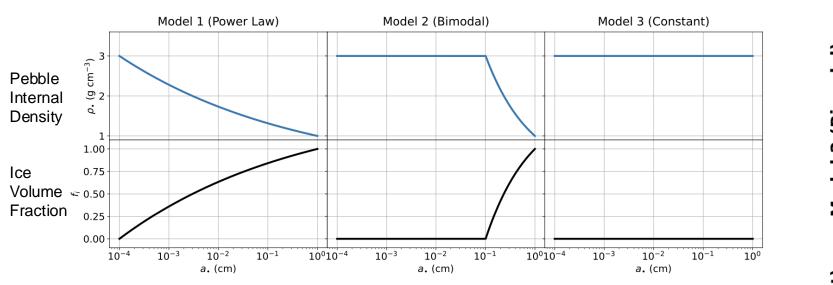


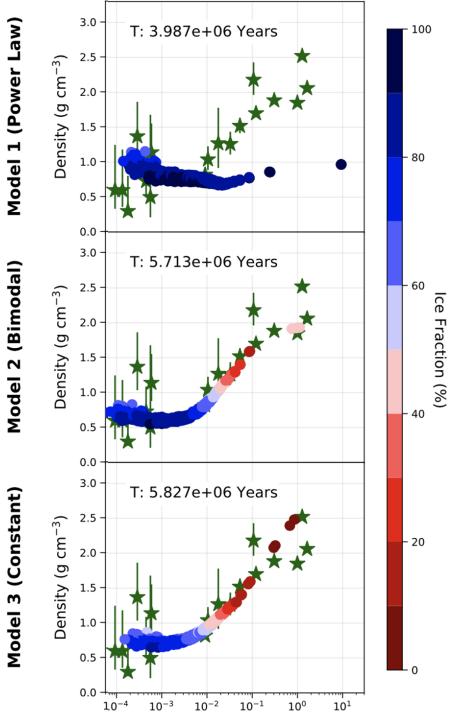


Growing Pluto by silicate pebble accretion

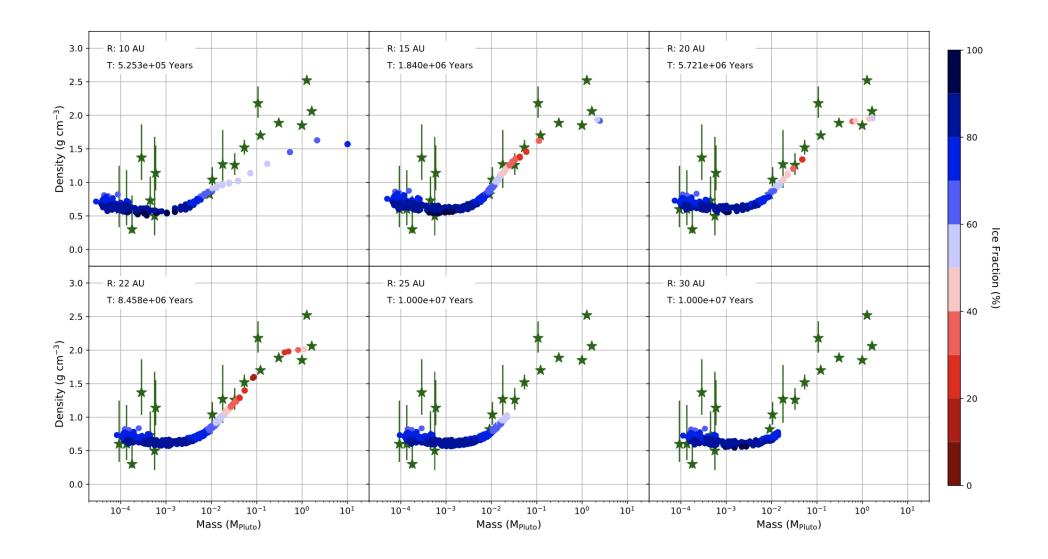


Resulting Densities vs Mass relations

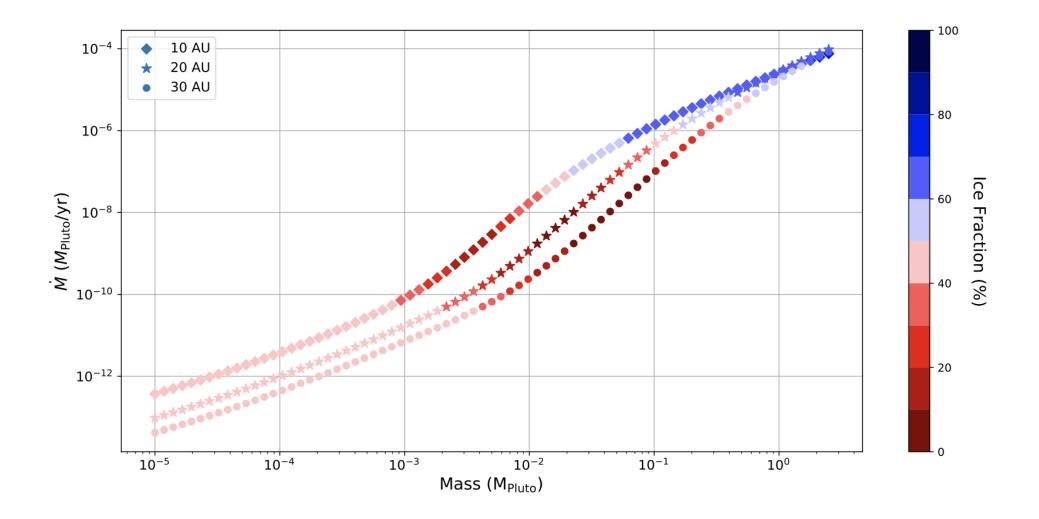




Distance Range 15 - 25AU



The window of silicate accretion



Conclusions

- Polydisperse Bondi accretion 1-2 orders of magnitude more efficient than monodisperse
 - Best accreted pebbles are those of drag time ~ Bondi time, not the largest ones
 - The largest ones dominate the mass budget, but accrete poorly
- Onset of Bondi accretion 1-2 orders of magnitude lower in mass compared to monodisperse
 - Bondi accretion possible on top of Streaming Instability planetary embryos within disk lifetime
 - Reaches 100-350km objects within Myr timescales
- Analytical solution to
 - Monodisperse general case
 - Polydisperse 2D Hill and 3D Bondi
- KBO density problem:
 - Two different pebble populations, maintained by ice desorption off small grains
 - Streaming instability: icy-rich small objects; nearly uniform composition
 - Polydisperse pebble accretion: silicate-rich larger objects; varied composition
 - Melting avoided by
 - ice-rich formation
 - ²⁶Al incorporated mostly in long (>Myr) phase of silicate accretion
 - KBOs best reproduced between 15-25 AU

